

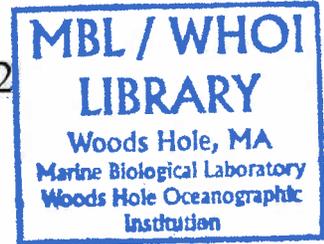
8/2003

5
GC
71
M32
2003

Mixing Processes and Hydraulic Control in a Highly Stratified Estuary

By

Daniel George MacDonald
B.S.C.E., University of New Hampshire, 1992
M.S., Cornell University, 1996



Submitted in partial fulfillment of the requirements for the degree of

DOCTOR OF PHILOSOPHY

at the

MASSACHUSETTS INSTITUTE OF TECHNOLOGY

and the

WOODS HOLE OCEANOGRAPHIC INSTITUTION

February 2003

© 2003, Daniel George MacDonald
All rights reserved.

The author hereby grants to MIT and WHOI permission to reproduce paper and electronic copies of this thesis in whole or in part and to distribute them publicly

Signature of Author

Joint Program in Oceanography/Applied Ocean Science and Engineering
Massachusetts Institute of Technology
and Woods Hole Oceanographic Institution
September 18, 2002

Certified by

W. Rockwell Geyer
Senior Scientist
Thesis Supervisor

Accepted by

Mark A. Grosenbaugh
Chair, Joint Committee for Applied Ocean Science and Engineering
Massachusetts Institute of Technology/
Woods Hole Oceanographic Institution

7/2/03

wam



Mixing Processes and Hydraulic Control in a Highly Stratified Estuary

by

DANIEL G. MACDONALD

MIT/WHOI Joint Program in Oceanography and Oceanographic Engineering

Submitted to the Massachusetts Institute of Technology/Woods Hole Oceanographic Institution Joint Program in Oceanography and Oceanographic Engineering in September 2002 in partial fulfillment of the requirements for the Degree of Doctor of Philosophy

ABSTRACT

This thesis utilizes field data from the Fraser River Estuary, a highly stratified system located in southwestern British Columbia, Canada, to investigate the nature of mixing processes in a highly stratified environment, and to extend two-dimensional hydraulic theory to a three dimensional environment.

During the late ebb, a stationary front exists at the Fraser mouth. Although densimetric Froude numbers in the vicinity of the front are supercritical in a frame of reference parallel to the local streamlines, the front itself is oriented such that the value of the Froude number is equal to the critical value of unity when taken in a frame of reference perpendicular to the front. This observation presents a robust extension of established two-dimensional, two-layer hydraulic theory to three dimensions, and implies similarity with trans-sonic flows, in that a Froude angle can be used to identify critical conditions in a manner similar to the Mach angle.

Mixing processes were evaluated at the mouth during the late ebb using a control volume approach to isolate mean vertical entrainment processes from turbulent processes, and quantify the vertical turbulent salt and momentum fluxes. Observed turbulent dissipation rates are high, on the order of $10^{-3} \text{ m}^2 \text{ s}^{-3}$, with vertical entrainment velocities on the order of $2 \times 10^{-3} \text{ m} \cdot \text{s}^{-1}$. Mixing efficiencies, expressed as flux Richardson numbers, are confined within a range from 0.15 to 0.2, at gradient Richardson number values between 0.2 and 0.25. These results are consistent with previous laboratory studies, but represent energetic conditions that are several orders of magnitude higher.

In the estuarine channel, the variability of mixing processes was investigated through the tidal cycle using control volume and overturn scale methods. Spatially, mixing was observed to be more intense near a width constriction on the order of 25%. Temporally, more dominant mixing was observed during ebbs, due to increases in both vertical shear and stratification. Mixing is active and important throughout the tidal cycle, and was found to be the dominant process responsible for removing salt from the estuarine channel during the ebb.

Thesis supervisor: Dr. W. Rockwell Geyer

Title: Senior Scientist, Woods Hole Oceanographic Institution

Acknowledgements

This work could never have been accomplished without assistance and encouragement from a large number of people. It has truly been a pleasure to work with my advisor, Rocky Geyer, over the last four years. His enthusiasm for his work has been inspiring and I have learned a lot from him. I also appreciate the timely advice and input I have received from the other members of my committee, John Trowbridge, Eric Adams, and Ole Madsen, and I would like to thank Jim Ledwell for chairing my defense.

The Fraser River cruises would not have been possible without help from David Jay, Philip Orton and Doug Wilson. The crew of the R/V Clifford A. Barnes, Ray McQuin and Nikki Hix, was fantastic to work with. Alex Horner was a tremendous help during both cruises, particularly with the collection of the data described in Chapter 4, and its subsequent analysis. Here at WHOI, Terry Donoghue and Jim Irish provided a tremendous amount of assistance in preparation for the two cruises.

The Coastal and Estuarine GFD course at Friday Harbor Labs during the summer of 1999 was a great experience and provided me with a good foundation for the research presented here. I'd like to thank Peter Rhines and Steve Monismith for setting up and running the course, and all the visitors, teaching assistants, and other students for making it such a productive summer. In particular, Jody Klymak and Derek Fong both provided great insight, ideas, and advice as I began to scratch the surface of the Fraser River data.

The other students and staff here at Woods Hole are what make it such a great place to work and study. I'd like to thank the education office staff, Marsha Bissonette, Julia Westwater, and Marcey Simon, for making the administrative side of the Joint Program such a pleasant experience. Other Joint Program students, especially the incoming class of 1998 and all of the AOPE students, have been a great resource for both science and friendship. Many thanks to Charlie Stock, Ben Reeder, Fernanda Hoefel, Joe Warren, Ben Evans, Oscar Pizarro, Jon Woodruff, and many others that have made these years in the Joint Program so worthwhile. Outside of the Joint Program, I have to thank Eric Amundsen for convincing me to head back to graduate school, and Paul Arnold, Dave Corso, and Bruce Eaton for many great hikes over the years.

My most consistent supporters on the long road towards this Ph.D. have always been my family: my parents, Ken and Claire, and my brothers, Rob and Jon. Their encouragement over the past 32 years has been irreplaceable. My wife Lisa has always provided me with unqualified support in my graduate school and other endeavors since we first met. Her support has been limitless, and so is my appreciation. Thank you Lisa.

This research was funded by Office of Naval Research grants N000-14-97-10134 and N000-14-97-10566, National Science Foundation grant OCE-9906787, a National Science Foundation graduate fellowship, and the WHOI Academic Programs Office.

Contents

Abstract	3
Acknowledgements	5
1 Introduction	15
1.1 The Physical Significance of Estuaries.....	16
1.2 The Physics of Mixing in Stratified Fluids.....	18
1.3 Two-Layer Hydraulic Theory.....	21
1.4 The Fraser River Estuary.....	24
1.5 The 1999 and 2000 Field Efforts.....	29
1.6 Structure of the Thesis.....	31
2 The Three-Dimensional Structure and Hydraulic Control of a Highly Stratified Estuarine Front	35
2.1 The Near-Field Plume.....	37
2.1.1 Hydraulic Control.....	39
2.1.2 Structure of River Plumes.....	40
2.2 The Fraser Estuary and the 2000 Field Effort.....	43
2.3 Structure of the Lift-Off Zone.....	45
2.3.1 Streamlines of the Upper Layer.....	46
2.3.2 Lift-Off Cross Sections.....	48
2.3.3 Front Location and Temporal Variability.....	51
2.4 Control of the Front.....	55
2.4.1 Interface Height Determination.....	58

2.4.2	Froude Angle	60
2.4.3	Expansion Control vs. Bottom Control	64
2.5	Kinematics of the Salt Supply	66
2.5.1	Distribution of Salt Flux in Salinity and Vertical Space	70
2.5.2	Potential Energy Changes Within the Lift-Off Zone	75
2.6	Cross Channel Dynamics	77
2.7	Implications of Froude Angle Control and Local Kinematics	81
2.7.1	Importance of Lateral Flow	85
2.8	Concluding Remarks on Lift-Off Zone Structure	86
3	Turbulent Energy Production and Entrainment at a Highly Stratified Estuarine Front	89
3.1	Mixing in Shear-Stratified Flows	91
3.1.1	Turbulent Kinetic Energy and Mixing Efficiency	91
3.1.2	Entrainment Velocity	95
3.1.3	Turbulent Length Scales	100
3.2	The Fraser River and the 1999 Field Effort	102
3.3	Hydrography Seaward of the Front	105
3.4	Estimation of Turbulent Fluxes	109
3.4.1	Measured Quantities	113
3.4.2	Volume Conservation and Vertical Velocity	117
3.4.3	Salt Conservation	119
3.4.4	Momentum Conservation	121
3.4.5	Flux Richardson Numbers	126
3.5	Lateral Effects	128
3.6	Turbulence, Entrainment and Closure	133
3.6.1	Local Production vs. Advected TKE	134
3.6.2	Mean Vertical Velocities	136

3.6.3	Mechanisms of turbulence generation	137
3.7	Concluding Remarks on Mixing in the Lift-Off Zone	139
4	The Variability of Vertical Salt Flux in a Highly Stratified Estuarine Channel	141
4.1	Introduction	143
4.2	The Mechanics of Shear-Induced Mixing	146
4.2.1	Gradient Richardson Number	146
4.2.2	Froude Numbers and Hydraulic Theory	147
4.2.3	Turbulent Salt Transport and Buoyancy Flux	147
4.2.4	Overturn Scales	148
4.2.5	Stratification	150
4.3	The Fraser River Estuary	151
4.4	Data Collection During Freshet, Spring Tide Conditions	155
4.4.1	Integrated Observations From a Down-Stream Time Series	155
4.4.2	Buoyancy Flux Estimates Using a Control Volume Approach	160
4.4.3	Overturn Scale Estimates	166
4.4.4	Stratification Profiles	170
4.4.5	Richardson Number Profiles	173
4.4.6	Froude Number Profiles	178
4.5	Temporal Variability of Mixing Processes	181
4.5.1	Initial Flood	181
4.5.2	First Ebb	183
4.5.3	Second Flood	183
4.5.4	Final Ebb	184
4.6	Spatial Variability of Mixing Processes During Ebb Tide	185
4.7	A Summary of Mixing Through the Tidal Cycle	186
4.7.1	Mixing in the Fraser	186

4.7.2	Comparisons With Other Estuaries	191
4.8	Concluding Remarks on Mixing Variability	192
5	Conclusions	193
5.1	The Dynamic Cycle of the Fraser River	194
5.2	Contributions of the Thesis	197
5.2.1	Three-Dimensional Extension of Two-Layer Hydraulic Theory	197
5.2.2	Development of Control Volume Method For Turbulent Flux Calculations	198
5.2.3	Mixing Efficiencies (R_{ij}) in Energetic Shear-Stratified Flows	199
5.2.4	Simple Turbulent Closure Scheme for Pure Shear-Stratified Flows ..	199
5.2.5	Importance of Mixing Throughout the Tidal Cycle	200
5.3	Unresolved Issues	201
	Appendix: A New Stratification Parameter	203
	Bibliography	207

List of Figures

1.1	Location and plan view of the Fraser River Estuary	25
1.2	Plots of tidal stage from the two field collection efforts	28
1.3	Fraser River discharge	31
2.1	Definition sketch for front and near-field regions	38
2.2	Estuarine channel, and mouth regions of the Fraser River Estuary	44
2.3	Mouth of the Fraser, with upper layer streamlines	47
2.4	Along channel cross-sections through the lift-off zone	49
2.5	Transverse channel cross-sections through the lift-off zone	50
2.6	Observed width of frontal zone	52
2.7	Maps of the intersection of the 20 psu isohaline with the bottom	54
2.8	Upper layer Froude number across the lift-off zone	57
2.9	Streamwise cross section, velocity profiles, and interface height	59
2.10	Propagation of wave fronts for critical and supercritical flow	61
2.11	Froude angle estimates for lift-off region	63
2.12	Froude Angle Path and Prandtl Meyer Expansion Fan at Sand Heads	65
2.13	Zones for control volume salt balance analysis	67
2.14	Definition sketch for control-volume salt balance analysis	68
2.15	Distribution of salt flux relative to salinity	71
2.16	Distribution of salt flux relative to depth	72
3.1	Cartoons of entrainment	98
3.2	Plan view of estuary mouth, with ship tracks	104
3.3	Cross sections through the front	106
3.4	Profiles of the squared composite Froude number across the front	110
3.5	Two-dimensional schematic of the control volume method	114

3.6	Vertical profiles of velocity and salinity in the control volume region	115
3.7	Vertical profiles of the gradient Richardson number	116
3.8	Vertical profiles of the mean vertical velocities, w_h and w_j	118
3.9	Buoyancy flux, and eddy diffusivity profiles from control volume analysis ...	120
3.10	Estimates of surface elevation across the lift-off region	122
3.11	Turbulent momentum flux profiles from control volume analysis	124
3.12	TKE production, and eddy viscosity profiles from control volume analyses	125
3.13	Flux Richardson number as a function of the gradient Richardson number	127
3.14	Profiles of cross-stream velocity and lateral volume flux ratio	130
3.15	Lateral influx ratios for both salt and momentum	132
4.1	The main arm of the Fraser River	153
4.2	Composite profiles of salt wedge salinity structure	154
4.3	Tidal height, surface velocity, and bottom velocity at the anchor station	156
4.4	Contours of salinity through a tidal cycle at the anchor station	157
4.5	Integrated landward and seaward salt fluxes at anchor station	159
4.6	Profile of average buoyancy flux with respect to salinity	162
4.7	Spatial variability of buoyancy flux during the ebb (control volume analyses).	165
4.8	Temporal variability of overturn scale and buoyancy flux at anchor station	167
4.9	Salinity profile at the anchor station, two hours after second high tide	169
4.10	Spatial variability of overturn scale and buoyancy flux during ebb	171
4.11	Stratification at anchor station	172
4.12	Subcritical Richardson numbers through the tidal cycle at anchor station	174
4.13	Bulk Richardson number through the tidal cycle at anchor station	176
4.14	Subcritical Richardson numbers across the channel during ebb	177
4.15	Froude number time series at anchor station	179
4.16	Froude number profiles during ebb	180
4.17	Temperature salinity relationships for time series CTD casts	182
4.18	Schematic representation of boundary mixing due to secondary circulation	187

List of Tables

2.1	Contributions to the 3-D salt balance	68
2.2	Salt input/output distribution statistics	73
4.1	Summary of buoyancy flux estimates through the tidal cycle	190

Chapter 1

Introduction

1.1 The Physical Significance of Estuaries

The discharge of fresh river water represents a significant source of nutrients, sediments, pollutants and other terrestrially derived material to the world's oceans (Yin et al., 1995; Caspers, 1967; Sommerfield and Nittrouer, 1999). During its journey to the sea, the fresh river discharge first interacts with ocean water within an estuary situated at or near the river mouth. Because very different water masses are colliding within an estuary, estuaries can often be characterized by strong gradients, of which the most dynamically important are density and velocity. Strong gradients of dissolved and suspended constituents such as nutrients, pollutants, and suspended sediment are also typically present, and it is the ultimate disposition of these components that provides much of the motivation for understanding the local physics.

The nature and intensity of the mixing processes within an estuary are driven, to a large extent, by the dynamically important gradients of density and velocity. Shear instabilities resulting from perturbations in the velocity profile large enough to overcome the ambient density stratification provide the primary mechanism for turbulence generation and mixing (Thorpe, 1973). At the seaward end of the estuary, the density difference between the discharge and the ambient coastal water has been greatly reduced, but the brackish residual that remains can continue to persist in the coastal ocean and be carried great distances, sometimes on the order of 10^2 to 10^3 kilometers, as the plume forms a coastal current (Yankovsky and Chapman, 1997). Eventually, through the action of external forcing mechanisms, such as wind and wave action, the signal of the plume is diluted by mixing with ocean water (Fong, 1998; Lentz, 2001).

