

# Effect of continuous light on leaf wax isotope ratios in *Betula nana* and *Eriophorum vaginatum*: Implications for Arctic paleoclimate reconstructions

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## Highlights

- *E. vaginatum* (sedge) and *B. nana* (shrub) grown under continuous or diurnal light
- The effect of continuous light on  $\delta D_{\text{wax}}$  varies between species
- $\delta D_{\text{wax}}$  is more negative for *E. vaginatum* than *B. nana* regardless of light
- $\delta^{13}C_{\text{wax}}$  is more negative in continuous light than in diurnal light

## Abstract

Reconstructions of climate using leaf wax D/H ratios ( $\delta D_{\text{wax}}$ ) require accounting for the apparent isotopic fractionation ( $\epsilon_{\text{app}}$ ) between plant source water and waxes. There have been conflicting publications on whether plants in the Arctic growing under 24-hour continuous light fractionate less than temperate and tropical plants. In this study, we examine the effect of diurnal light (DL) versus 24-hour continuous light (CL) on the isotopic composition of leaf *n*-alkanes and *n*-acids in greenhouse experiments using two common Arctic plants (*Eriophorum vaginatum*, or tussock cottongrass, and *Betula nana*, or dwarf birch). For *E. vaginatum*, the  $\delta D_{\text{wax}}$  values of various wax homologues were 5-11‰ more positive for CL plants relative to their DL counterparts, whereas for *B. nana*, CL waxes were 3-24‰ more negative, suggesting that daylight length is not a unifying control on leaf wax D/H ratios of Arctic plants. The  $\delta^{13}C_{\text{wax}}$  of *B. nana* was more negative for plants grown in continuous light compared to diurnal light, reflecting lower water-use efficiency associated with prolonged stomatal opening in the CL treatment. We modeled the impact of increasing stomatal conductance and effective flow path lengths (mimicking variable leaf morphologies) on the isotopic composition of leaf waters ( $\delta D_{\text{lw}}$ ) and find that variations in leaf-water enrichment may explain the variable  $\delta D_{\text{wax}}$  responses seen between *E. vaginatum* and *B. nana*. We suggest that between-species differences in the  $\delta D_{\text{lw}}$  response to light, and differences in the utilization of stored carbohydrates were important for governing  $\delta D_{\text{wax}}$ . Our greenhouse results suggest that Arctic plant leaf waxes do not consistently display reduced  $\epsilon_{\text{app}}$  values as a result of 24-hour day light, providing additional support for field observations.

Keywords: Leaf waxes, hydrogen isotopes, carbon isotopes, growth experiment, Arctic, continuous light

## 1.1 Introduction

Hydrogen isotope ratios of terrestrial leaf waxes ( $\delta D_{\text{wax}}$ ) are a powerful paleoclimate proxy, as waxes are abundant in sedimentary archives, geochemically stable, and  $\delta D_{\text{wax}}$  can be used to infer rainfall D/H ratios ( $\delta D_{\text{precipitation}}$ ) (Huang et al., 2004; Sachse et al., 2004; Hou et al., 2008; Garcin et al., 2012). Yet,  $\delta D_{\text{wax}}$  not only depends on the isotopic composition of source water, but also on the net apparent D/H fractionation ( $\epsilon_{\text{app}}$ ) between meteoric water and waxes. Apparent fractionation varies with changes in evaporative enrichment of soil water ( $\delta D_{\text{soil}}$ ) and leaf water ( $\delta D_{\text{lw}}$ ), as well as changes in biosynthetic isotope fractionations associated with different plant species, phenological stage, and environmental conditions (Roden and Ehleringer, 1999; Polissar and Freeman, 2010; Gao et al., 2014; Tipple et al., 2015; Freimuth et al., 2017). As  $\epsilon_{\text{app}}$  can be sensitive to environmental factors and vegetation types, variations in  $\epsilon_{\text{app}}$  must be considered and, if possible, corrected in paleoclimate reconstructions (Polissar and Freeman, 2010; Feakins, 2013; Konecky et al., 2016).

To this end, there are numerous surveys of  $\delta D_{\text{wax}}$  in modern vegetation and sediments in tropical and temperate regions that quantify  $\epsilon_{\text{app}}$ . Leaf wax samples from living plants exhibit a large range in  $\epsilon_{\text{app}}$ , from -34 to -202‰ (Gao et al., 2014), with monocotyledonous plants generally exhibiting greater biosynthetic fractionation than dicotyledons (Gao et al., 2014; Liu et al., 2016). D/H measurements of sedimentary waxes demonstrate that  $\epsilon_{\text{app}}$  generally ranges from -100 to -150‰ at the integrated watershed scale (Garcin et al., 2012; Sachse et al., 2012), although some studies from the Arctic suggest a much weaker fractionation (ie. less negative  $\epsilon_{\text{app}}$ ). Waxes preserved in lake surface sediments from Baffin Island (Shanahan et al., 2013) and in ancient (ca. 30 ka) paleosols in western Canada (Porter et al., 2016) indicate  $\epsilon_{\text{app}}$  of -60‰. Similarly, both grasses and woody plants from Baffin Island and Greenland appear to exhibit 20-50‰ less fractionation compared to temperate plants (Yang et al., 2011). These studies have led to suggestions of widespread weak apparent fractionation in the Arctic; however, they have relied on either modeled or reconstructed precipitation isotopic ratios for calculating fractionation factors, which, if incorrect, could bias estimates of  $\epsilon_{\text{app}}$ . Indeed, a recent global synthesis of plant leaf wax isotopes shows no such latitude effect on  $\epsilon_{\text{app}}$  (Liu et al., 2016). Furthermore, studies of lake sediments in Siberia (Wilkie et al., 2012), Alaska (Daniels et al., 2017), and northern Europe (Sachse et al., 2004) have demonstrated  $\epsilon_{\text{app}}$  values that are similar to those of the mid-latitudes. Thus, it is uncertain if and why Arctic plant D/H fractionation differs from mid- and low-latitude plants, and, if so, what values of  $\epsilon_{\text{app}}$  are appropriate for estimating ancient  $\delta D_{\text{precipitation}}$  based on measurements of  $\delta D_{\text{wax}}$ .

The principal mechanism to explain small  $\epsilon_{\text{app}}$  values in the Arctic posits that continuous daylight leads to increased/prolonged transpiration (E), and thereby enhanced evaporative D-enrichment of leaf water and correspondingly small apparent fractionation (Yang et al., 2009; Yang et al., 2011). An alternative leaf water theory, however, suggests that higher rates of E should result in less D-enrichment of leaf waters by increasing the advection of unfractionated xylem water to the site of biosynthesis (Barbour and Farquhar, 2004; Barbour et al., 2004; Song et al., 2013). Continuous transpiration, as is sometimes observed in plants transferred into continuous-light environments (Van Gestel et al., 2005), could therefore result in more D-depleted leaf waters, and by extension more negative  $\delta D_{\text{wax}}$ , relative to plants exhibiting diurnal

patterns in transpiration. A greenhouse experiment testing the effect of continuous light found that  $\epsilon_{\text{app}}$  was, in contrast, 40‰ more positive for plants receiving continuous-light compared to diurnal-light (Yang et al., 2009). The experimental conditions of this prior study, however, may not have been ideal for testing the effect of light because two of the species studied, *Taxodium sp.* and *Metasequoia sp.*, have not been found in the Arctic since the Eocene and because treatment differences in temperature and humidity (Equiza et al., 2006) could have confounded the effects of continuous light. Furthermore, multiple studies have demonstrated that metabolic hydrogen isotope fractionation and the contribution of stored carbohydrates affect lipid D/H ratios, varying with plant type (Kahmen et al., 2013b), position within a leaf (Gao et al., 2015), environmental conditions (Cormier et al., 2018), and seasonality (Sessions, 2006; Newberry et al., 2015), although little is known about how this factor would be affected by the diel light cycle. An experimental growth study using modern tundra plants and identical environmental conditions could constrain the sensitivity of modern tundra vegetation  $\delta D_{\text{wax}}$  to a 24-hour light effect.

Carbon isotope ratios of leaf waxes also record vegetation and climatic information (Diefendorf and Freimuth, 2017), and so are useful in paleoclimate research (Aichner et al., 2015; Konecky et al., 2016). As in the case of hydrogen isotopes, little is known about how  $\delta^{13}\text{C}_{\text{wax}}$  of plants grown in continuous light differs from plants grown in diurnal light. The  $\delta^{13}\text{C}$  of leaf biomass depends on interfoliar  $\text{CO}_2$  concentrations (Farquhar et al., 1982), and so an effect of light cycle on *n*-alkyl lipid  $\delta^{13}\text{C}$  could emerge due to changes in photosynthetic rates or stomatal conductance. At a global scale, there is no correlation between latitude and  $^{13}\text{C}/^{12}\text{C}$  discrimination (Diefendorf et al., 2010); however, the greenhouse study of Yang et al. (2009) revealed a decrease in  $\delta^{13}\text{C}$  values of bulk leaf tissue from plants grown under continuous light. Thus, it remains uncertain if compound-specific measurements of  $\delta^{13}\text{C}_{\text{wax}}$  are sensitive to continuous versus diurnal light.

In summary, empirical results from Arctic field studies are inconclusive as to whether the Arctic light environment affects the isotopic composition of leaf waxes. Using a controlled growth chamber experiment, we test the effect of 24-hour sunlight on leaf wax isotope ratios in two Arctic taxa with distinct growth forms. We evaluate the results in the context of a model examining how evaporative enrichment and leaf water circulation could influence  $\delta D_{\text{wax}}$  and  $\epsilon_{\text{app}}$  under different light environments.

## 2. Methods

### 2.1 Growth experiment

To assess the effect of the diel light cycle on  $\epsilon_{\text{app}}$  values,  $\delta D_{\text{wax}}$ , and  $\delta^{13}\text{C}_{\text{wax}}$ , we performed a growth chamber experiment at Brown University. We examined two common arctic tundra plants, the graminoid, *Eriophorum vaginatum var. spissum* (common name: tussock cottongrass) and the deciduous shrub, *Betula nana* (common name: dwarf birch). These two species together represent approximately 40% of above-ground biomass in tussock tundra of Northern Alaska (Chapin III et al., 1995), and are important components of tundra ecosystems across northern Eurasia and Canada (Walker et al., 2005). Eighteen specimens of *E. vaginatum* and ten specimens of *B. nana* were obtained in individual pots from the Welker's Grove Nursery in Connecticut, USA. Plants were divided equally into two E7/2 Conviron™ Plant Growth Chambers (height: 58 cm, area: 0.76 m<sup>2</sup>), allowing control of environmental parameters. One chamber was configured with a diurnal light cycle (DL), emulating temperate photoperiod, and

the other with continuous light (CL) such as can be found in Arctic summers.

A light sensor was used to measure photosynthetically active radiation (PAR) levels at hourly intervals. In our study, PAR varied between 0 and 350  $\mu\text{moles photons m}^{-2} \text{ s}^{-1}$  in the DL chamber and between 215 and 250  $\mu\text{moles photons m}^{-2} \text{ s}^{-1}$  in the CL chamber (Fig. 1). The maximum PAR in our greenhouse experiment was approximately 50% of the maximum natural light in the Arctic (e.g. Williams et al., 2014) due to space and resource constraints in the growth chamber. Temperatures varied between 10 and 15 °C on a diel cycle (Fig. 1) and humidity was held at 75%. Plants were grown in loamy peat and were watered on average every 5 days during the experiment, to minimize water stress while keeping the soil from becoming anoxic. Irrigation water (degassed tapwater) was stored in sealed carboys over the course of the experiment, thereby maintaining constant isotope ratios throughout the experiment. The isotope values, measured on a Picarro cavity ring-down analyzer at Brown University, had a  $\delta\text{D}$  value of  $-39.0 \pm 0.2\text{‰}$  and  $\delta^{18}\text{O}$  value of  $-6.1 \pm 0.04\text{‰}$ , relative to VSMOW. To ensure we measured waxes produced only during the growth period, we trimmed the *E. vaginatum* specimens to approximately 2 cm length at the onset of the experiment. For *B. Nana*, we pruned branches and leaves, and marked remaining leaves so that they could be avoided in later harvesting. At the end of this growth period, we compared leaf wax hydrogen and carbon isotopes between the two treatments to determine the magnitude of the continuous light effect on  $\delta\text{D}_{\text{wax}}$  and  $\delta^{13}\text{C}_{\text{wax}}$ .

## 2.2 Leaf wax analysis

After 3 months of growth, new *E. vaginatum* growth (entire tillers) and new *B. Nana* leaves were harvested, composited for each individual specimen, and freeze-dried for processing and isotopic analysis. Leaf waxes from each plant were extracted and purified using standard procedures (Gao et al., 2014). The *n*-acids were derivatized to fatty acid methyl esters (FAMES) using acidified methanol of known isotopic composition (Yang and Huang, 2003). The molecular composition of straight chain *n*-acids and *n*-alkanes were measured using a gas chromatograph (GC; 6890 gas chromatograph; Agilent Technologies) equipped with a split/splitless injector held at 320 °C and a flame ionization detector (FID) held at 325 °C. A 30 m fused silica column (HP-1MS, 0.32 mm i.d., 0.25  $\mu\text{m}$  film thickness) was used with a helium carrier gas with a flow rate of 1.7 mL/min. The oven temperature program was: 60°C (1 min), ramping to 220 °C at 20 °C/min, then to 315 °C at 6 °C/min, where it was held for 15 min. Identification was done by comparison with a standard containing a suite of even-chain-length FAMES ranging from *n*-C<sub>16</sub> to *n*-C<sub>28</sub>, and a second standard containing C<sub>25</sub>, C<sub>27</sub>, C<sub>29</sub>, C<sub>30</sub>, and C<sub>32</sub> *n*-alkanes.

Carbon isotope ratios were determined using a Finnigan Delta V isotope-ratio mass spectrometer (IRMS) with the same GC configuration as in GC-FID analysis. Samples were run in duplicate and an external standard containing either FAMES or alkanes of known isotopic composition was run in between every 6 injections. Not all *n*-acid and *n*-alkane homologues were measured for isotopic ratios due to prohibitively small abundances; as such, the maximum number of samples for a given compound and treatment was 9 for *E. vaginatum* and 5 for *B. nana*, but for some compounds, *n* was lower. The  $\delta^{13}\text{C}$  values of FAMES were mathematically corrected for the methyl groups added during derivatization using the following equation:

$$1) \delta\text{D}_{\text{real}} = [(n + 1)\delta^{13}\text{C}_{\text{measured}} + 36.52]/n,$$

where *n* is carbon chain length of the compound and -36.52‰ is the  $\delta^{13}\text{C}$  value of the added methanol. Compound specific hydrogen isotope ratios were determined in an identical fashion, in

triplicate, using a Finnigan Delta Plus IRMS and the  $\delta D$  of FAMES were likewise corrected for the hydrogen added during methylation using:

$$2) \delta D_{corrected} = [(2n + 2)\delta D_{measured} + 123.7 * 3]/(2n - 1),$$

where  $n$  is carbon chain length of the compound and  $-123.7\%$  is the  $\delta D$  value of the added methanol. For  $\delta D$  measurements, only chromatogram peaks with areas between 2000 mV and 6000 mV are included in the analysis. Uncertainty was estimated from the results of replicate plant specimen.

Statistical comparisons were made using a 2-way ANOVA in R, with species and light treatment as independent factors. Statistical analyses were conducted independently for each lipid homologue. Data met the assumptions of normality (Shapiro-Wilk test,  $\alpha=0.05$ ) and homogeneity of variance (Levene's test,  $\alpha=0.05$ ) and no data required transformations. Where the 2-way ANOVA found statistical differences at the  $p<0.05$  level, pairwise comparisons were performed using a Student's t-test to evaluate the effect of light for each test species, and to test the effect of species within each light treatment. To avoid type I error among the multiple comparisons, we only consider pairwise differences significant if the p-value was below a critical value of 0.01.

## 2.2 Model sensitivity test

We also modeled the isotopic composition of leaf waters to evaluate the potential effect of 24-hour daylight on the  $\delta D$  of leaf water. The model calculates steady state values of  $\delta D_{w}$ , and is applicable for understanding the potential effect of diurnal light cycles on  $\delta D_{w}$ . We adapt a commonly used leaf water  $\delta^{18}O$  model to predict  $\delta D_{w}$  across a range of stomatal conductance ( $g_s$ ) values (Flanagan et al., 1991; Roden and Ehleringer, 1999; Barbour et al., 2004; Tipple et al., 2015; Cernusak et al., 2016). The basis of the model is a modified version of the Craig-Gordon evaporation model (Flanagan et al., 1991) that accounts for kinetic fractionation through both stomatal pores and across the leaf boundary layer. The model approximates the enrichment of evaporated water over source water ( $\Delta_e$ ) using Eq. 3 :

$$3) \Delta_e = \varepsilon^* + \varepsilon_k + (\Delta_v - \varepsilon_k) \frac{e_a}{e_i}$$

where  $\varepsilon^*$  is the equilibrium fractionation between  $HD^{16}O$  and  $H_2^{16}O$ ,  $\varepsilon_k$  is the diffusive fractionation through the leaf boundary layer,  $\Delta_v$  is the isotopic difference between external water vapor and source water, and  $e_a$  and  $e_i$  are the external and internal vapor pressures, respectively. We modified the values of the relevant parameters from Barbour et al. (2004) to describe the fractionation and diffusivities of  $HD^{16}O$  rather than  $H_2^{18}O$ .

Equilibrium fractionation ( $\varepsilon^*$ ) is calculated as a function of temperature (Eq. 4;  $T$  is in kelvin) (Majoube, 1971).

$$4) \varepsilon^* = \exp\left(-0.052612 + \frac{76.248}{T} - \frac{24.844}{T^2}\right)$$

Diffusive fractionation ( $\varepsilon_k$ ) is comprised of fractionation through the stomata and through the boundary layer (Eq. 3) and depends on the kinetic fractionation ( $\varepsilon_s$  and  $\varepsilon_b$ ) and diffusive

conductance ( $g_s$  and  $g_b$ ) through each layer.  $\epsilon_s$  has an H/D fractionation of 1.025, while  $\epsilon_b$  is calculated as  $\epsilon_s$  to the two-thirds power, or 1.017 (Roden and Ehleringer, 1999). The boundary layer conductance is set to  $1 \text{ mol m}^{-2} \text{ s}^{-1}$  following Barbour et al. (2004), although we note that slightly higher values have also been reported (Song et al., 2013). Values of  $g_s$  are allowed to vary from 0 to  $1 \text{ mol m}^{-2} \text{ s}^{-1}$  in order to evaluate the sensitivity of  $\delta D_{wax}$  to this variable.

$$5) \epsilon_k = \frac{25/g_s + 17/g_b}{1/g_s + 1/g_b}$$

The isotopic composition of mesophyll water from which waxes are biosynthesized, is a mixture of leaf water that has experienced evaporative enrichment at the leaf edge and xylem water that is isotopically similar to soil water and not yet undergone transpiration-related D-enrichment (Roden and Ehleringer, 1999; Barbour et al., 2004). The relative fraction of each component can be estimated empirically; however, advances in understanding water movement through leaves indicates that a Péclet-based model of the relative contribution of each pool can improve estimates of leaf water  $\delta^{18}\text{O}$  and leaf cellulose  $\delta^{18}\text{O}$  (Barbour et al., 2004). The Péclet number describes the ratio of advected xylem water to evaporated water diffused from the leaf surface at the site of biosynthesis, and is described by Eq. 6,

$$6) \wp = \frac{L \cdot E}{C \cdot D}$$

where  $L$  is the effective path length (m),  $E$  is the rate of transpiration ( $\text{mmol m}^{-2} \text{ s}^{-1}$ ),  $C$  is the molar density of water ( $55.6 \times 10^3 \text{ mol m}^{-3}$ ), and  $D$  is the diffusivity of  $\text{DH}^{18}\text{O}$  in water.  $D$  is set to  $1.78 \times 10^{-9} \text{ m}^2 \text{ s}^{-1}$  (Mills, 1973), a slight decrease compared to the parameter for  $\delta^{18}\text{O}$  ( $3.4 \times 10^{-9} \text{ m}^2 \text{ s}^{-1}$ ). The resulting offset ( $\Delta_{lw}$ ) between the isotopic composition of the average leaf water and the xylem water is thus described as,

$$7) \Delta_{lw} = \frac{\Delta_e(1 - e^{-\wp})}{\wp}$$

We use this model to explore how leaf physiology (described by  $L$ ) influences the sensitivity of  $\delta D_{wax}$  to different illumination regimes. In this exercise,  $g_s$  serves as a proxy for illumination conditions, as it is thought to be the main mechanism connecting light cycles to  $\delta D_{wax}$  (Yang et al., 2009).

### 3. Results

#### 3.1 Greenhouse experiment

We observed production of new foliar biomass over the three-month growth period in both study species and treatments. The cut *E. vaginatum* grew to approximately 20 cm in length. *B. Nana* showed less growth, but new leaves appeared and were sampled for analysis. No quantitative information on growth is available.

#### 3.2 Leaf wax distributions

For both species, *n*-acids show a strong even over odd predominance, whereas *n*-alkanes show strong odd over even predominance (Fig. 2; Table 1). *B. Nana* has a higher concentration of waxes, except for the C<sub>30</sub> and C<sub>32</sub> *n*-acid homologues, of which *E. vaginatum* has more. Both species also produce a small amount of short-chain (C<sub>16</sub> and C<sub>18</sub>) *n*-acids, but no corresponding short-chain *n*-alkanes. The average chain length (ACL) was determined for the C<sub>20</sub>-C<sub>32</sub> *n*-acids and C<sub>21</sub>-C<sub>32</sub> *n*-alkanes. ACL for *B. nana* is 25.9 ± 0.4 for *n*-acids and 28.8 ± 0.4 for *n*-alkanes. ACL for *E. vaginatum* is 26.5 ± 0.5 for *n*-acids and 28.5 ± 0.2 for *n*-alkanes. No difference is observed in the wax distributions of either plant in response to light treatments.

### 3.3 Hydrogen isotope ratios

$\delta D_{wax}$  ranges from -207‰ to -135‰ across all species, treatments and leaf wax homologues. The pooled analytical standard deviation of replicate standards and samples averages 2.6‰ across all homologues (Table 2). Calculated against the irrigation water  $\delta D$  value of -39‰,  $\epsilon_{app}$  varies from -100‰ to -176‰ (Fig. 3, Table 3).

We observe an inconsistent effect of light treatment on  $\delta D_{wax}$ . For *E. vaginatum*, the C<sub>24</sub> *n*-acid and C<sub>27</sub> and C<sub>31</sub> *n*-alkanes are 6-11‰ D-enriched in the CL treatment relative to the DL treatment ( $p < 0.01$ ). For other lipid homologues the effect of light was insignificant. In contrast, the CL waxes of *B. nana* are 3-24‰ more D-depleted than DL waxes ( $p < 0.01$ ), with the exception of C<sub>28</sub> *n*-acid that showed no difference between light treatments. A species effect is apparent wherein *n*-acids and *n*-alkanes of *E. vaginatum* are D-depleted by, on average, 36 ± 16‰ relative to *B. nana* ( $p < 0.001$  for all homologues). The species effect is present in both the diurnal and continuous light treatments.

### 3.4 Carbon isotope ratios

$\delta^{13}C_{wax}$  ranges from -42.3 to -32.0‰ across all species, treatments, and leaf wax homologues (Fig. 3, Table 4), and the pooled analytical standard deviation averages 0.22‰ across all homologues (Table 2). For *B. nana*,  $\delta^{13}C_{wax}$  averages 2.8‰ lower in the CL treatment than in the DL treatment, although no difference is observed for the C<sub>22</sub> *n*-acid ( $p = 0.038$ ) and C<sub>28</sub> *n*-acid ( $p = 0.021$ ). For *E. vaginatum*,  $\delta^{13}C$  of wax homologues average 0.8‰ lower under CL than DL, although the differences are not statistically significant for any of the homologues ( $p > 0.01$  for all compounds). The difference in  $\delta^{13}C_{wax}$  between the two species is insignificant for most wax homologues, with the exception that the C<sub>22</sub>-C<sub>26</sub> *n*-acids were <sup>13</sup>C-enriched in the DL *B. nana* compared to the DL *E. vaginatum*.

Upon combining data from both light treatments and various wax homologues, we find significant ( $p < 0.001$ ) correlations between  $\delta^{13}C_{wax}$  and  $\delta D_{wax}$  for the *n*-alkanes for both taxa (Fig. 4). The correlation is positive for *Betula nana* ( $r = 0.85$ ) and negative for *Eriophorum vaginatum* ( $r = 0.50$ ). For *n*-acids, the  $\delta^{13}C_{wax}$ - $\delta D_{wax}$  correlation is weakly positive for *B. nana* ( $r = 0.28$ ,  $p = 0.002$ ) and absent entirely for *E. vaginatum* ( $r = 0.03$ ,  $p = 0.200$ ), possibly reflecting slower regeneration rates of the *n*-acids (Gao and Huang, 2013) or more variable hydrogen exchange for that compound class.

### 3.5 Leaf water model results

We investigate the sensitivity of  $\delta D_w$  to changing stomatal conductance using a modified

version of the Craig-Gordon model (Equations 1-5) in which the environmental parameters are similar to our growth chamber environments with  $\delta D_{\text{sourceWater}} = -39\text{‰}$ ,  $rH = 80\%$ , and leaf temperature =  $13\text{ °C}$ . The enrichment over source water is also a function of  $\delta D_{\text{vapor}}$ , which is unconstrained in our study, and was prescribed to  $-121\text{‰}$  assuming equilibrium fractionation between the  $\delta D_{\text{sourceWater}}$  ( $-39\text{‰}$ ) and  $\delta D$  of vapor at the average incubation temperature ( $13\text{ °C}$ ) (Majoube, 1971). The patterns of sensitivity are robust across different values of  $\delta D_{\text{vapor}}$ . It is important to note that in the model,  $E$  and  $g_s$  are positively correlated, which is only the case if vapor pressure is held constant. In field studies from the Arctic diel changes in temperature and  $rH$  can be more important factors controlling transpiration than stomatal conductance (Gebauer et al., 1998). In the current experiment, the diel changes in temperature and  $rH$  were identical between light treatments and so should not greatly interfere with a test of stomatal conductance.

Using the modified Craig-Gordon leaf water model, we estimate that at the site of evaporation, leaf water isotopes are  $24\text{‰}$  D-enriched over irrigation water under the conditions of our growth experiment, resulting in a  $\delta D_{\text{w}}$  value of  $-15\text{‰}$ . The  $\delta D_{\text{w}}$  at the site of evaporation is not strongly dependent on  $g_s$ , but after accounting for the mixing of evaporated and unevaporated water within the leaf (i.e. the Péclet number), we find that  $\delta D_{\text{w}}$  of the bulk leaf water decreases with increasing  $g_s$ , especially for leaf morphologies with high effective path lengths ( $L$ ) (Fig. 5). At high  $g_s$  and high Péclet number (i.e. large leaves or mixing path length), modeled  $\delta D_{\text{w}}$  is only slightly D-enriched relative to the irrigation water because relatively D-depleted xylem water dominates sites of wax synthesis within leaf mesophyll. In low  $L$  cases (small leaves),  $\delta D_{\text{w}}$  is insensitive to  $g_s$ , but shows consistently large D-enrichment relative to source water. As such, we might expect the largest leaves (high  $L$ ) to be most responsive to a 24-hour light effect on  $\delta D_{\text{w}}$  or  $\delta D_{\text{wax}}$ , if  $g_s$  varies as a function of the light cycle.

Song et al. (2013) showed that the effective path length can also vary as a function of gas exchange rates across the stomata. As  $E$  increases,  $L$  decreases due to a shift to more direct, extracellular, flowpaths of water through the leaf. In the model case where  $L$  varies with transpiration rate (dark red lines in Fig. 5), the Péclet number is small and relatively stable (Eq. 4), and  $\delta D_{\text{w}}$  increases slightly ( $< 3\text{‰}$ ) with increasing  $g_s$ .

## 4. Discussion

### 4.1 *The theoretical basis for $\delta D_{\text{wax}}$ response to light duration*

Plants regulate their evapotranspiration rates to maximize the ratio of  $\text{CO}_2$  uptake relative to water loss by controlling their stomatal conductance (Farquhar et al., 1989). There is also a rich body of evidence that stomatal opening and closure occurs directly in response to light availability (Darwin, 1898; Sharkey and Ogawa, 1987). Illustrating this point are multiple studies documenting that both angiosperms and gymnosperms, when transferred into a continuous light environments, will begin continuously photosynthesizing and transpiring over a 24-hour cycle (Van Gestel et al., 2005; Yang et al., 2009). Genetic differences between species, however, can occasionally maintain circadian rhythms in  $g_s$  despite continuous light or dark environments (Kerr et al., 1985; Hennessey and Field, 1991), and so we do not necessarily expect all species to respond in the same manner to alterations in daily light duration. Furthermore, the stomatal response to light availability depends on water availability – if soil water is limited, plants may be forced to close their stomata even in the presence of sunlight to avoid desiccation (Farquhar et al., 1989). It is not immediately obvious whether the integrated stomatal conductance ( $g_s$ ) of



Arctic plants is in fact higher or lower than the same plants grown in low- or mid-latitudes (Gebauer et al., 1998; Llorens et al., 2009), especially given that the intensity of light in the Arctic during mid-summer can vary by over 1000  $\mu\text{mol photons m}^{-2} \text{s}^{-1}$  on a diel cycle (Williams et al., 2014) and given that there are large variations in moisture across the Arctic (Shanahan et al., 2013).

The presence or absence of a diel light cycle could affect  $\delta D_{\text{wax}}$  by influencing the magnitude of evaporative enrichment at the leaf surface, by influencing the mixing ratio between unfractionated xylem water and fractionated leaf water at the site of biosynthesis, or by influencing whether the hydrogen atoms used in lipid synthesis are derived from stored or freshly produced organic substrates (Sessions, 2006). If  $\delta D_{\text{iw}}$  is the main control on  $\delta D_{\text{wax}}$ , then the specific stomatal response to the light cycle is a critical consideration. Indeed, the leaf water model results suggest that increased  $g_s$  in a CL environment would likely lead to more negative  $\delta D_{\text{iw}}$  values, similar to predictions by (Sullivan and Welker, 2007), but at odds with the experimental results of Yang et al. (2009).

We might expect that in our growth experiment, CL plants express more negative  $\delta D_{\text{wax}}$  than DL plants because the daily integrated  $\delta D_{\text{iw}}$  values decrease with prolonged periods of transpiration. Alternatively, CL plants might express more positive  $\delta D_{\text{wax}}$  than DL plants because the daily peak light intensity (Fig. 1) and, correspondingly, the daily peak instantaneous values of  $g_s$ , are lower in the CL treatment. That is, the daily duration of stomatal opening and photosynthesis may be less important than the instantaneous values of  $g_s$  during the period of photosynthesis, particularly given that  $\delta D_{\text{iw}}$  reaches steady state values rapidly (<2 hours) upon stomatal opening or closure (Roden and Ehleringer, 1999). The magnitude of leaf water enrichment during the dark hours are irrelevant because photosynthesis is not occurring. Variations in  $\delta D_{\text{wax}}$  should thus only be discernible where there are non-linear relationships between photosynthesis and  $g_s$ , or where there are variations in biosynthetic strategies.

#### 4.2 D/H fractionation under varying light in growth experiments

Few studies on the diurnal patterns of transpiration in continuous light have focused on changes in the hydrogen isotopic composition of leaf biomass. Previous observations (Yang et al., 2009; 2011) have suggested a relatively large (15-40‰) effect of CL on plant wax hydrogen isotopic compositions (Fig. 6). Our greenhouse experiment demonstrates that the effect of CL versus DL can vary substantially between Arctic-dwelling species, pointing to a need to better constrain spatial and species differences in  $\epsilon_{\text{app}}$  across light regimes. The opposing  $\delta D_{\text{wax}}$  responses between our two study species to the light treatment is surprising and indicates that day length is not a unifying control on the  $\delta D_{\text{wax}}$  of plants.

The magnitude of D/H fractionation observed for *B. nana* and *E. vaginatum* in our greenhouse study agrees almost perfectly with field studies of the same species conducted in the Alaskan Arctic (Daniels et al., 2017). Comprehensive analyses of the isotopic composition of soil water, leaf water, and waxes showed that, in the field, *B. nana* has an  $\epsilon_{\text{app}}$  of -100 to -110‰ for various *n*-acid and *n*-alkane homologues, compared to -100 to -147‰ for our greenhouse specimens. The  $\epsilon_{\text{app}}$  values for *n*-alkanes ranges from -100 to -121‰, almost identical to the range of -102 to -118‰ reported for *n*-alkanes from a related species, *B. pubescens*, sampled across a wide range of latitudes, including latitudes that receive 24-hour daylight in summer (Sachse et al., 2006). Likewise,  $\epsilon_{\text{app}}$  of waxes from *E. vaginatum* in Alaska range from -155 to -170‰

(Daniels et al., 2017) and from -145 to -176‰ in this growth experiment. Additionally, our greenhouse experiments yield  $\epsilon_{app}$  values that are similar to those from studies of plants within the same families and growth forms grown in DL environments (Gao et al., 2014), indicating that our study is broadly representative for these two species in a range of environmental conditions. This is noteworthy, as the light intensity in the growth chambers is somewhat less than occurs in nature. The consistency of  $\epsilon_{app}$  values observed for these plants, in the field and in experiments, and across very different light cycles and intensities strongly suggests that light is not the principle control on  $\epsilon_{app}$ .

Previously, Liu et al. (2016) showed that the  $\epsilon_{app}$  values of dicots and monocots exhibit similar latitudinal gradients (a proxy for day length, among other variables). Our data suggest these two plant groups might display differing sensitivities to day length. For *E. vaginatum*, CL waxes are the same or D-enriched (6-11‰) relative to DL waxes, but CL waxes are D-depleted (3-24‰) for *B. nana*. The model results imply that either the two species altered their  $g_s$  in opposite directions in response to continuous light, that leaf morphology modulates the impact of  $g_s$  variability in governing  $\delta D_w$ , or alternatively that the two species exhibited contrasting biosynthetic responses to light. There is no *a priori* reason to assume the stomatal response to light treatment differ for *E. vaginatum* and *B. nana*. Thus, the contrasting responses likely originate from morphologic or biosynthetic differences in these plants.

A possible explanation for why the two species responded differently is that  $\delta D_{wax}$  is less sensitive to environmental changes in monocotyledonous plants, such as *E. vaginatum*, than it is in dicotyledonous plants such as *B. nana* (Kahmen et al., 2013b). These authors demonstrate that dicots express the full magnitude of leaf water enrichment in their waxes, whereas only 18-68% of leaf water variability is expressed in the waxes of monocots. In this framework, *B. nana* specimens in the CL treatment may have experienced a decrease in  $\Delta_w$  relative to the DL treatment, associated with an increase in stomatal conductance (Fig. 5). This shift in  $\Delta_w$  was then expressed as more negative  $\epsilon_{app}$ . A similar decrease in  $\Delta_w$  in *E. vaginatum* would not have been completely expressed in the lipid  $\delta D$  values.

The slight increase in  $\epsilon_{app}$  for some of the *E. vaginatum* wax homologues in response to continuous light suggests that specific plant traits might affect not only the sensitivity, but the direction, of the  $\delta D_{wax}$  response to light. Our isotope modeling results suggest that leaf morphology, reflected in  $L$ , could contribute to the species-specific differences (Barbour and Farquhar, 2004; Cernusak et al., 2016). In particular, the model predicts that for plants in which  $L=f(E)$  (Song et al., 2013), changing leafwater flowpaths could potentially alter the direction of the  $\delta D_w$  response to continuous light. Unfortunately, we do not have sufficient species-specific information on  $L$  to assess the importance of this factor directly in our experiments, and future work could assess leaf water dynamics, particularly within grasses, under different light regimes. In general, both monocots and dicots exhibit a gradient of increasing  $\delta D_w$  and  $\delta D_{wax}$  from the base or center of the leaf to the leaf edge (Gao and Huang, 2013; Liu et al., 2016). While our leaf water model does not account for variations in  $\Delta_w$  along the leaf blade, the positive shift in  $\epsilon_{app}$  for *E. vaginatum* is a possible indication that back-diffusion of D-enriched water from the terminus of the leaf may increase as stomatal conductance increases under CL, i.e.  $L=f(E)$ , as in the dark red line of Fig. 5. Likewise, the large  $\delta D_{wax}$  response to CL observed by Yang et al. (2009) could derive from morphological traits of the conifer needles they studied, wherein the low-

intensity/long-duration light gave rise to greater D-enrichment. In conifers needles, there is a resistant physical barrier between the xylem and mesophyll that limits water transport between these two components (Roden et al., 2015; Cernusak et al., 2016) which could make the leaf water used in biosynthesis more sensitive to evaporative enrichment and a 24-hour light effect. Analogous morphological factors could be important for *E. vaginatum*, which showed the same direction of  $\delta D_{\text{wax}}$  change as the conifers, albeit to a lesser degree.

Experimental differences in temperature regimes and watering protocols may have also contributed to the contrasting results between our results and those of Yang et al. (2009). For comparison, in order to mimic Eocene Arctic conditions, Yang et al. (2009) kept soils saturated throughout the growth phase in both treatments such that water stress did not limit transpiration (Equiza et al., 2006; Yang et al., 2009), whereas our intermittent watering protocol may have resulted in occasional stomatal closure as might occur in naturally dry conditions. That said, none of the plants shows outward signs of water stress in our experiment, and so other factors most likely cause the results of our experiment to differ from those of Yang et al. (2009). In general, the contrasting  $\delta D_{\text{wax}}$  responses between *E. vaginatum* (a graminoid), *B. nana* (a broadleaf angiosperm), and conifers (Yang et al., 2009) (Fig. 6) suggest that leaf morphology could modulate the daylength response of  $\delta D_{\text{wax}}$ . Future study across a broader array of species could further elucidate a relationship between plant morphology and a light effect on  $\delta D_{\text{wax}}$ .

In addition to the role of  $\delta D_{\text{iv}}$  as a control on long-chain *n*-alkane and *n*-acid  $\delta D$  values within leaves and across environmental gradients (Sachse et al., 2006; Gao et al., 2015; Tipple et al., 2015), there is abundant evidence that biosynthetic fractionation ( $\epsilon_{\text{bio}}$ ) is an important control on  $\delta D_{\text{wax}}$ .  $\epsilon_{\text{bio}}$  is determined by the D/H of wax precursors and the D/H fractionation in various biosynthetic pathways, and it varies between monocots and dicots (Gao et al., 2014; Liu et al., 2016), throughout the season (Sessions et al., 1999; Newberry et al., 2015; Freimuth et al., 2017), and as a function of light availability and  $p\text{CO}_2$  (Cormier et al., 2018). It remains difficult to disentangle the joint effects of  $\epsilon_{\text{bio}}$  and  $\Delta_{\text{iv}}$  in our study without direct measurements of photosynthesis rates or leaf water isotope measurements. In general, waxes that are produced from recent photosynthate tend to be D-depleted compared to waxes derived from stored sugars (Sessions, 2006; Cormier et al., 2018). As such, a relatively greater utilization of newly produced carbohydrates for *B. nana* wax production under CL relative to DL could explain the more negative  $\delta D_{\text{wax}}$  values in the CL treatment. Greater utilization of fresh photosynthate under CL is also supported by the lower  $\delta^{13}\text{C}$  values (Helle and Schleser, 2004). In *E. vaginatum*, on the other hand, no such effect is apparent. The lower concentrations of waxes in the *E. vaginatum* specimens (Fig. 2) indicate lower requirements for wax production such that lipid precursors may be present in sufficient quantity in both the CL and DL treatment and no shift in metabolic pathway is required. If differences in photosynthesis rates or wax formation pathways differed between light treatments, it will be important to perform perennial growth experiments for determining if continuous light could eventually affect  $\delta D$  values of the stored carbohydrates and NADPH pools, thereby resulting in a different effect of light than that observed over the 3 months of this experiment.

Lastly, we highlight the important role that vegetation plays as a determinant of ecosystem-scale  $\epsilon_{\text{app}}$ . The considerable difference in  $\epsilon_{\text{app}}$  between *E. vaginatum* and *B. nana* agrees with findings showing that grasses fractionate between D and H more strongly than shrubs (Hou et al., 2007; Gao et al., 2014), and that species more basal in their evolutionary histories tend to

show smaller  $\epsilon_{\text{app}}$  than later-derived plants (Gao et al., 2014). The three conifers grown by Yang et al. (2009), *Taxodium distichum*, *Larix laricina*, and *Metasequoia glyptostroboides*, had  $\epsilon_{\text{app}}$  values from -95 to -106‰ in the DL treatment, somewhat smaller than the angiosperms *B. nana* (-110‰) and the *E. vaginatum* (-150‰) grown here. While these differences do not directly explain the response of *Larix* and *Metasequoia* to changes in light conditions, the differences in  $\epsilon_{\text{app}}$  for DL-grown plants between our study and that of Yang et al. (2009) are not unexpected. The between-species difference in  $\epsilon_{\text{app}}$  eclipses leaf water enrichment effects, pointing to a need to further quantify  $\epsilon_{\text{app}}$  for a broader range of species in order better constrain vegetation effects on paleoclimate reconstructions using  $\delta D_{\text{wax}}$ .

#### 4.3 $^{13}\text{C}/^{12}\text{C}$ fractionation under varying light in growth experiments

Variations in light regime have previously been documented to influence the  $\delta^{13}\text{C}$  of bulk leaf biomass (Smith et al., 1976; Ehleringer et al., 1986; Pearcy and Pfitsch, 1991; Yang et al., 2009) but changes of  $\delta^{13}\text{C}$  of leaf wax compounds in response to day length have not been described. We observe an average  $\delta^{13}\text{C}_{\text{wax}}$  depletion of 1.8‰ for plants in the CL treatment relative to the DL treatment, although the effect is smaller and non-significant in the *E. vaginatum* specimens. The direction and magnitude of the observed  $\delta^{13}\text{C}_{\text{wax}}$  shift is similar to those observed in several other studies that examined bulk leaf  $\delta^{13}\text{C}$  and found that higher light intensity induces less carbon isotope discrimination. Yang et al. (2009) document a negative shift of 1.8-4.6‰ for plants grown under low-intensity/continuous light compared to high-intensity/diurnal light; both Pearcy and Pfitsch (1991) and Lockheart et al. (1997) document a negative shift of 1-2‰ for shade leaves relative to sun leaves; and, Ehleringer et al. (1986) report  $\delta^{13}\text{C}$  shifts on the order of 4‰ across a light intensity gradient.

A likely explanation for the shift in  $\delta^{13}\text{C}_{\text{wax}}$  is that the light treatment gives rise to changes in the intercellular concentration of  $\text{CO}_2$  ( $C_i$ ), and thereby  $^{13}\text{C}/^{12}\text{C}$  discrimination, by impacting photosynthesis rates and/or stomatal conductance (Farquhar et al., 1982; Diefendorf et al., 2010).  $\delta^{13}\text{C}_{\text{wax}}$  generally decreases as water-use efficiency decreases, reflecting for example, an increase in  $g_s$  in response to greater water availability or light, or a reduction in photosynthesis,  $A$ . It is difficult to partition the effects of varying  $A$  versus varying  $g_s$  (Sullivan and Welker, 2007) without direct measurements of these variables. Yang et al. (2009) observed sustained transpiration by plants in CL, and, given that plants were not water stressed in either experiment, we propose that increased/sustained  $g_s$  can explain lower  $\delta^{13}\text{C}_{\text{wax}}$  in our CL plants. This carbon isotope-inferred change in  $g_s$  further suggests that stomatal regulation likely plays some role in governing leaf water D-enrichment and  $\delta D_{\text{wax}}$  in response to light availability, particularly for *B. nana*. The correlations between  $\delta^{13}\text{C}_{\text{wax}}$  and  $\delta D_{\text{wax}}$  (Fig. 4) have not previously been noted for our study taxa. The correlation is positive for *B. nana* ( $n$ -acid  $R^2=0.28$ ,  $n$ -alkane  $R^2<0.85$ ) but negative for *E. vaginatum* ( $n$ -acid  $R^2 = 0.03$ ,  $n$ -alkane  $R^2=0.50$ ). A similar contrast in the  $\delta^{13}\text{C}$ - $\delta D$  relationships was previously reported between  $C_3$  trees and  $C_4$  grasses (Bi et al., 2005). Our results, from two  $C_3$  species, suggest that growth form (i.e. woody vs herbaceous), in addition to photosynthetic pathway, may be an important determinant of whether the slope is positive or negative. Tandem measurements of the two isotopes could provide a useful means of identifying wax sources in Arctic lakes.

Given the experimental results showing that  $\delta^{13}\text{C}$  of *B. nana* leaf wax is sensitive to the

intensity and diurnal cycle of light, we might expect there should be a latitude effect on  $\delta^{13}\text{C}$  of foliar biomass at a global scale. The apparent absence of correlation between  $\delta^{13}\text{C}$  and latitude (Diefendorf et al., 2010), however, suggests that other environmental factors mask the effect of latitude, or that the gradient in light intensity and day length is weaker across latitudes than it was between our experimental treatments. Thus, caution is necessary in applying these results to field studies.

#### 4.4 Application to Field Studies

How applicable are our growth experiments to understanding the isotopic composition of leaf waxes in the Arctic? Although light is available continuously during polar summer, daily rhythms in temperature, plant photosynthesis, soil moisture, and even light availability exist at high latitude sites and may lead to diel patterns of gas exchange in plants. For example, at Toolik Field Station in Alaska (69°N latitude), hourly PAR measurements during the summer solstice from 1999 to 2013 show a strong daily cycle from 30  $\mu\text{mol photons m}^{-2} \text{s}^{-1}$  at ‘night’ up to 980  $\mu\text{mol photons m}^{-2} \text{s}^{-1}$  at midday.

For *Betula sp.* and *Eriophorum sp.*, the relationship between light and  $g_s$  is ill-defined (Gebauer et al., 1998). Jagels and Day (2000) indicate that  $g_s$  and E can be as much a function of temperature, water availability, and species as it can be of light. Moreover, diurnal studies from the Arctic show day/night differences in E despite continuously available PAR. Gebauer et al. (1998) found that E decrease during “night” in the arctic sedge *Eriophorum*, similar to results from Bliss (1960) for *Betula* and *Salix* shrub specimens in northern Alaska. Diel patterns in E have also been observed in eddy flux measurements in the tundra and are attributed to a stomatal conductance control (Vourlitis and Oechel, 1999). These measurements demonstrate that  $g_s$  is most sensitive to vapor pressure deficit and water stress in the tundra and show that in the field, under continuous light, many Arctic plants experience diurnal cycles in key physiologic processes that would influence the isotopic compositions of their leaf waters. Given these results, it is not surprising that continuous light does not have a consistent effect on the hydrogen isotopic composition of Arctic plant waxes.

In the Arctic, day length changes substantially over the course of the year. Given the differences we observe between CL and DL treatments, the annual cycle in day length could possibly impart a small seasonal signal in  $\epsilon_{\text{app}}$ , species dependent. Over long time scales, however, the annual cycle of day length at a particular location is approximately unchanging. As such, the interpretation of sedimentary  $\delta\text{D}_{\text{wax}}$  records at a given location is not contingent on the effect of day length. More important is the identification of accurate fractionation values in the Arctic and information on the latitude and light dependencies of  $\epsilon_{\text{app}}$  at ecosystem or plant scales.

Our data, based on leaf-level measurements, do not conclusively support a 24-hour light effect to explain small  $\epsilon_{\text{app}}$  values ( $\sim -60\text{‰}$ ) observed in some studies in the Arctic. Furthermore, Arctic sites, which experience similar day lengths, exhibit highly variable  $\epsilon_{\text{app}}$  signatures indicating that day length likely plays a secondary role to vegetation and climate effects (Sachse et al., 2006; Yang et al., 2011; Wilkie et al., 2012; Shanahan et al., 2013; Porter et al., 2016; Thomas et al., 2016). These other factors, or possibly inaccuracies in  $\epsilon_{\text{app}}$  estimates, are important for explaining the  $\epsilon_{\text{app}}$  differences across Arctic sites, particularly with regards to the small fractionation inferred at Baffin Island (Shanahan et al., 2013) and central Canada (Porter et al., 2016) which were previously attributed to a 24-hour continuous light effect.

Species differences in our greenhouse and field experiments could implicate vegetation type as an important control on  $\epsilon_{app}$  variations across the Arctic. Likewise, relative humidity, which we did not test in this experiment, has been frequently identified as a critical influence on  $\epsilon_{app}$  through its control on leaf evaporation (Kahmen et al., 2013a; Tipple et al., 2015). At present, however, Baffin Island, Lake El'gygytgyn, and the North Slope of Alaska all have growing season rH values similar to the 75% rH used in our growth experiment (average 74%, 78%, and 75%, respectively), yet field studies at those sites (Wilkie et al., 2012; Shanahan et al., 2013; Daniels et al., 2017) document difference in  $\epsilon_{app}$  of approximately 40‰ between the sites. Likewise, variable soil conditions may contribute to pan-Actic  $\epsilon_{app}$  variations by influencing the magnitude of evaporative enrichment of soil waters, but at present, there is insufficient information on the D/H or  $^{18}\text{O}/^{16}\text{O}$  ratios of soil waters at sites with leaf wax measurements to quantitatively assess the importance of soil evaporation. We cannot rule out inaccuracies in the estimates of  $\delta\text{D}$  of source water in the Arctic studies, which are difficult to constrain and strongly influence the field estimates of  $\epsilon_{app}$ . Shanahan et al. (2013) relied upon modeled, rather than measured, source water isotopic composition and used modeled mean annual  $\delta\text{D}_{precipitation}$ . There is evidence, however, that Arctic plants derive much of their biosynthetic water from precipitation received during the growing season (Wilkie et al., 2012; Daniels et al., 2017). The precipitation-weighted mean annual  $\delta\text{D}_{precipitation}$  at Hall Beach, Baffin Island is -161‰ and the growing season (JJA)  $\delta\text{D}_{precipitation}$  is -134‰ (IAEA/WMO, 2017), so the assumed seasonality of source water could contribute as much as 27‰ toward the approximately 40‰ difference in  $\epsilon_{app}$  estimated at Baffin Island relative to Alaska or Siberia. Likewise, Porter et al. (2016) compared the D/H of fossil waxes to fossil (permafrost) water to estimate  $\epsilon_{app}$ , but the seasonality and the age of the fossil water is not strictly known, creating uncertainty in the calculated  $\epsilon_{app}$ . Continued work constraining  $\epsilon_{app}$  at both the whole-ecosystem scale for a range of arctic bioclimatic subregions, as well as at the scale of individual plants, will help guide interpretations of Arctic sedimentary  $\delta\text{D}_{wax}$  records, particularly where  $\delta\text{D}_{wax}$  is paired with pollen or biomarker-based vegetation/bioclimatic information (Feakins, 2013).

## 5. Conclusion

Our greenhouse experiments and modeling studies establish that day length has an inconsistent and relatively small influence on apparent hydrogen isotopic fractionations between leaf wax lipids and source water for two prominent plant species found in high-latitude northern locations. Variable responses of  $\delta\text{D}_{wax}$  to light regime suggest that the effect of increased stomatal conductance in a 24-hour light environment may be modulated by leaf water dynamics or changing biosynthetic pathways, possibly explaining divergent results from previous growth experiments. Furthermore, the ~40‰ difference in  $\epsilon_{app}$  between the monocotyledon, *E. vaginatum*, and the dicotyledon, *B. nana*, supports previous research suggesting that vegetation differences are an important factor governing  $\delta\text{D}_{wax}$  across the modern Arctic (Wilkie, 2013; Gao et al., 2014; Daniels et al., 2017). In contrast to hydrogen isotopes, carbon stable isotope ratios are sensitive to the light regime, with more negative  $\delta^{13}\text{C}_{wax}$  values apparent in a continuous light environment. Overall, these results will guide reconstructions of  $\delta\text{D}_{precipitation}$  across vast areas of tundra where *B. nana* and *E. vaginatum* have been prominent in the plant communities for

millennia.

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## 7. References

Aichner, B., Feakins, S.J., Lee, J., Herzsuh, U. and Liu, X. (2015) High-resolution leaf wax carbon and hydrogen isotopic record of the late Holocene paleoclimate in arid Central Asia. *Climate of the Past* 11, 619.

Barbour, M. and Farquhar, G. (2004) Do pathways of water movement and leaf anatomical dimensions allow development of gradients in  $H_2^{18}O$  between veins and the sites of evaporation within leaves? *Plant, Cell Environ.* 27, 107-121.

Barbour, M.M., Roden, J.S., Farquhar, G.D. and Ehleringer, J.R. (2004) Expressing leaf water and cellulose oxygen isotope ratios as enrichment above source water reveals evidence of a Péclet effect. *Oecologia* 138, 426-435.

Bi, X., Sheng, G., Liu, X., Li, C. and Fu, J. (2005) Molecular and carbon and hydrogen isotopic composition of n-alkanes in plant leaf waxes. *Org. Geochem.* 36, 1405-1417.

Bliss, L. (1960) Transpiration rates of arctic and alpine shrubs. *Ecology* 41, 386-389.

Cernusak, L.A., Barbour, M.M., Arndt, S.K., Cheesman, A.W., English, N.B., Feild, T.S., Helliker, B.R., Holloway-Phillips, M.M., Holtum, J.A. and Kahmen, A. (2016) Stable isotopes in leaf water of terrestrial plants. *Plant, Cell Environ.* 39, 1087-1102.

Chapin III, F.S., Shaver, G.R., Giblin, A.E., Nadelhoffer, K.J. and Laundre, J.A. (1995) Responses of arctic tundra to experimental and observed changes in climate. *Ecology* 76, 694-711.

Cormier, M.A., Werner, R.A., Sauer, P.E., Gröcke, D.R., Leuenberger, M.C., Wieloch, T., Schleucher, J. and Kahmen, A. (2018)  $^3H$ -fractionations during the biosynthesis of carbohydrates

and lipids imprint a metabolic signal on the  $\delta^2\text{H}$  values of plant organic compounds. *New Phytol.*

Daniels, W.C., Russell, J.M., Giblin, A.E., Welker, J.M., Klein, E.S. and Huang, Y. (2017) Hydrogen isotope fractionation in leaf waxes in the Alaskan Arctic tundra. *Geochim. Cosmochim. Acta* 213, 216-236.

Darwin, F. (1898) Observations on Stomata. *Proceedings of the Royal Society of London* 63, 413-417.

Diefendorf, A.F. and Freimuth, E.J. (2017) Extracting the most from terrestrial plant-derived n-alkyl lipids and their carbon isotopes from the sedimentary record: A review. *Org. Geochem.* 103, 1-21.

Diefendorf, A.F., Mueller, K.E., Wing, S.L., Koch, P.L. and Freeman, K.H. (2010) Global patterns in leaf  $^{13}\text{C}$  discrimination and implications for studies of past and future climate. *Proceedings of the National Academy of Sciences* 107, 5738-5743.

Ehleringer, J., Field, C., Lin, Z.-f. and Kuo, C.-y. (1986) Leaf carbon isotope and mineral composition in subtropical plants along an irradiance cline. *Oecologia* 70, 520-526.

Equiza, M.A., Day, M.E. and Jagels, R. (2006) Physiological responses of three deciduous conifers (*Metasequoia glyptostroboides*, *Taxodium distichum* and *Larix laricina*) to continuous light: adaptive implications for the early Tertiary polar summer. *Tree Physiology* 26, 353-364.

Farquhar, G., Hubick, K., Condon, A. and Richards, R. (1989) Carbon isotope fractionation and plant water-use efficiency, *Stable isotopes in ecological research*. Springer, pp. 21-40.

Farquhar, G.D., O'Leary, M.H. and Berry, J.A. (1982) On the relationship between carbon isotope discrimination and the intercellular carbon dioxide concentration in leaves. *Funct. Plant Biol.* 9, 121-137.

Feakins, S.J. (2013) Pollen-corrected leaf wax D/H reconstructions of northeast African hydrological changes during the late Miocene. *Palaeogeogr., Palaeoclimatol., Palaeoecol.* 374, 62-71.

Flanagan, L.B., Comstock, J.P. and Ehleringer, J.R. (1991) Comparison of modeled and observed environmental influences on the stable oxygen and hydrogen isotope composition of leaf water in *Phaseolus vulgaris* L. *Plant Physiol.* 96, 588-596.



Freimuth, E.J., Diefendorf, A.F. and Lowell, T.V. (2017) Hydrogen isotopes of n-alkanes and n-alkanoic acids as tracers of precipitation in a temperate forest and implications for paleorecords. *Geochim. Cosmochim. Acta* 206, 166-183.

Gao, L., Edwards, E.J., Zeng, Y. and Huang, Y. (2014) Major Evolutionary Trends in Hydrogen Isotope Fractionation of Vascular Plant Leaf Waxes. *PLoS ONE* 9, e112610.

Gao, L., Guimond, J., Thomas, E. and Huang, Y. (2015) Major trends in leaf wax abundance,  $\delta^2\text{H}$  and  $\delta^{13}\text{C}$  values along leaf venation in five species of  $\text{C}_3$  plants: physiological and geochemical implications. *Org. Geochem.* 78, 144-152.

Gao, L. and Huang, Y. (2013) Inverse gradients in leaf wax  $\delta\text{D}$  and  $\delta^{13}\text{C}$  values along grass blades of *Miscanthus sinensis*: Implications for leaf wax reproduction and plant physiology. *Oecologia* 172, 347-357.

Garcin, Y., Schwab, V.F., Gleixner, G., Kahmen, A., Todou, G., Séné, O., Onana, J.-M., Achoundong, G. and Sachse, D. (2012) Hydrogen isotope ratios of lacustrine sedimentary n-alkanes as proxies of tropical African hydrology: insights from a calibration transect across Cameroon. *Geochim. Cosmochim. Acta* 79, 106-126.

Gebauer, R., Reynolds, J. and Tenhunen, J. (1998) Diurnal patterns of  $\text{CO}_2$  and  $\text{H}_2\text{O}$  exchange of the Arctic sedges *Eriophorum angustifolium* and *E. vaginatum* (Cyperaceae). *Am. J. Bot.* 85, 592-592.

Helle, G. and Schleser, G.H. (2004) Beyond  $\text{CO}_2$ -fixation by Rubisco – an interpretation of  $^{13}\text{C}/^{12}\text{C}$  variations in tree rings from novel intra-seasonal studies on broad-leaf trees. *Plant, Cell Environ.* 27, 367-380.

Hennessey, T.L. and Field, C.B. (1991) Circadian rhythms in photosynthesis oscillations in carbon assimilation and stomatal conductance under constant conditions. *Plant Physiol.* 96, 831-836.

Hou, J., D'Andrea, W.J. and Huang, Y. (2008) Can sedimentary leaf waxes record D/H ratios of continental precipitation? Field, model, and experimental assessments. *Geochim. Cosmochim. Acta* 72, 3503-3517.

Hou, J., D'Andrea, W.J., MacDonald, D. and Huang, Y. (2007) Hydrogen isotopic variability in leaf waxes among terrestrial and aquatic plants around Blood Pond, Massachusetts (USA). *Org. Geochem.* 38, 977-984.

Huang, Y., Shuman, B., Wang, Y. and Webb, T. (2004) Hydrogen isotope ratios of individual lipids in lake sediments as novel tracers of climatic and environmental change: a surface sediment test. *J. Paleolimnol.* 31, 363-375.

IAEA/WMO (2017) Global Network of isotopes in Precipitation. The GNIP Database.  
Accessible at: <http://www.iaea.org/water>.

Kahmen, A., Hoffmann, B., Schefuß, E., Arndt, S.K., Cernusak, L.A., West, J.B. and Sachse, D. (2013a) Leaf water deuterium enrichment shapes leaf wax n-alkane  $\delta D$  values of angiosperm plants II: Observational evidence and global implications. *Geochim. Cosmochim. Acta* 111, 50-63.

Kahmen, A., Schefuß, E. and Sachse, D. (2013b) Leaf water deuterium enrichment shapes leaf wax n-alkane  $\delta D$  values of angiosperm plants I: Experimental evidence and mechanistic insights. *Geochim. Cosmochim. Acta* 111, 39-49.

Kerr, P.S., Rufty, T.W. and Huber, S.C. (1985) Endogenous rhythms in photosynthesis, sucrose phosphate synthase activity, and stomatal resistance in leaves of soybean (*Glycine max* [L.] Merr.). *Plant Physiol.* 77, 275-280.

Konecky, B., Russell, J. and Bijaksana, S. (2016) Glacial aridity in central Indonesia coeval with intensified monsoon circulation. *Earth. Planet. Sci. Lett.* 437, 15-24.

Liu, J., Liu, W., An, Z. and Yang, H. (2016) Different hydrogen isotope fractionations during lipid formation in higher plants: Implications for paleohydrology reconstruction at a global scale. *Scientific Reports* 6, 19711.

Llorens, L., Osborne, C.P. and Beerling, D.J. (2009) Water-use responses of 'living fossil' conifers to CO<sub>2</sub> enrichment in a simulated Cretaceous polar environment. *Ann. Bot.* 104, 179-188.

Lockheart, M.J., Van Bergen, P.F. and Evershed, R.P. (1997) Variations in the stable carbon isotope compositions of individual lipids from the leaves of modern angiosperms: implications for the study of higher land plant-derived sedimentary organic matter. *Org. Geochem.* 26, 137-153.

Majoube, M. (1971) Oxygen-18 and deuterium fractionation between water and steam. *J. Chim. Phys. Phys.-Chim. Biol.* 68, 1423-+.

Mills, R. (1973) Self-diffusion in normal and heavy water in the range 1-45. deg. The Journal of Physical Chemistry 77, 685-688.

Newberry, S.L., Kahmen, A., Dennis, P. and Grant, A. (2015) n-Alkane biosynthetic hydrogen isotope fractionation is not constant throughout the growing season in the riparian tree *Salix viminalis*. Geochim. Cosmochim. Acta 165, 75-85.

Pearcy, R.W. and Pfitsch, W.A. (1991) Influence of sunflecks on the  $\delta^{13}\text{C}$  of *Adenocaulon bicolor* plants occurring in contrasting forest understory microsites. Oecologia 86, 457-462.

Polissar, P.J. and Freeman, K.H. (2010) Effects of aridity and vegetation on plant-wax  $\delta\text{D}$  in modern lake sediments. Geochim. Cosmochim. Acta 74, 5785-5797.

Porter, T.J., Froese, D.G., Feakins, S.J., Bindeman, I.N., Mahony, M.E., Pautler, B.G., Reichart, G.-J., Sanborn, P.T., Simpson, M.J. and Weijers, J.W.H. (2016) Multiple water isotope proxy reconstruction of extremely low last glacial temperatures in Eastern Beringia (Western Arctic). Quaternary Science Reviews 137, 113-125.

Roden, J., Kahmen, A., Buchmann, N. and Siegwolf, R. (2015) The enigma of effective path length for  $^{18}\text{O}$  enrichment in leaf water of conifers. Plant, Cell Environ. 38, 2551-2565.

Roden, J.S. and Ehleringer, J.R. (1999) Observations of hydrogen and oxygen isotopes in leaf water confirm the Craig-Gordon model under wide-ranging environmental conditions. Plant Physiol. 120, 1165-1174.

Sachse, D., Billault, I., Bowen, G.J., Chikaraishi, Y., Dawson, T.E., Feakins, S.J., Freeman, K.H., Magill, C.R., McInerney, F.A. and Van der Meer, M.T. (2012) Molecular paleohydrology: interpreting the hydrogen-isotopic composition of lipid biomarkers from photosynthesizing organisms. Annual Review of Earth and Planetary Sciences 40, 221-249.

Sachse, D., Radke, J. and Gleixner, G. (2004) Hydrogen isotope ratios of recent lacustrine sedimentary n-alkanes record modern climate variability. Geochim. Cosmochim. Acta 68, 4877-4889.

Sachse, D., Radke, J. and Gleixner, G. (2006)  $\delta\text{D}$  values of individual n-alkanes from terrestrial plants along a climatic gradient—Implications for the sedimentary biomarker record. Org. Geochem. 37, 469-483.

Sessions, A.L. (2006) Seasonal changes in D/H fractionation accompanying lipid biosynthesis in

*Spartina alterniflora*. *Geochim. Cosmochim. Acta* 70, 2153-2162.

Sessions, A.L., Burgoyne, T.W., Schimmelmann, A. and Hayes, J.M. (1999) Fractionation of hydrogen isotopes in lipid biosynthesis. *Org. Geochem.* 30, 1193-1200.

Shanahan, T., Hughen, K., Ampel, L., Sauer, P. and Fornace, K. (2013) Environmental controls on the  $^3\text{H}/\text{H}$  values of terrestrial leaf waxes in the Eastern Canadian Arctic. *Geochim. Cosmochim. Acta*.

Sharkey, T.D. and Ogawa, T. (1987) Stomatal Responses to Light, in: Zeiger, E., Farquhar, G.D., Cowan, I.R. (Eds.), *Stomatal Function*. Stanford University Press, Stanford, California.

Smith, B.N., Oliver, J. and Millan, C.M. (1976) Influence of carbon source, oxygen concentration, light intensity, and temperature on  $^{13}\text{C}/^{12}\text{C}$  ratios in plant tissues. *Botanical Gazette* 137, 99-104.

Song, X., Barbour, M.M., Farquhar, G.D., Vann, D.R. and Helliker, B.R. (2013) Transpiration rate relates to within-and across-species variations in effective path length in a leaf water model of oxygen isotope enrichment. *Plant, Cell Environ.* 36, 1338-1351.

Sullivan, P.F. and Welker, J.M. (2007) Variation in leaf physiology of *Salix arctica* within and across ecosystems in the High Arctic: test of a dual isotope ( $\Delta^{13}\text{C}$  and  $\Delta^{18}\text{O}$ ) conceptual model. *Oecologia* 151, 372-386.

Thomas, E.K., Briner, J.P., Ryan-Henry, J.J. and Huang, Y. (2016) A major increase in winter snowfall during the middle Holocene on western Greenland caused by reduced sea ice in Baffin Bay and the Labrador Sea.

Tipple, B.J., Berke, M.A., Hambach, B., Roden, J.S. and Ehleringer, J.R. (2015) Predicting leaf wax n-alkane  $^3\text{H}/\text{H}$  ratios: controlled water source and humidity experiments with hydroponically grown trees confirm predictions of Craig-Gordon model. *Plant, Cell Environ.* 38, 1035-1047.

Van Gestel, N.C., Nesbit, A.D., Gordon, E.P., Green, C., Paré, P.W., Thompson, L., Peffley, E.B. and Tissue, D.T. (2005) Continuous light may induce photosynthetic downregulation in onion—consequences for growth and biomass partitioning. *Physiol. Plant.* 125, 235-246.

Vourlitis, G.L. and Oechel, W.C. (1999) Eddy covariance measurements of  $\text{CO}_2$  and energy fluxes of an Alaskan tussock tundra ecosystem. *Ecology* 80, 686-701.

Walker, D.A., Raynolds, M.K., Daniëls, F.J., Einarsson, E., Elvebakk, A., Gould, W.A., Katenin, A.E., Kholod, S.S., Markon, C.J. and Melnikov, E.S. (2005) The circumpolar Arctic vegetation map. *Journal of Vegetation Science* 16, 267-282.

Wilkie, K. (2013) Compound-specific hydrogen isotopes of lipid biomarkers in Lake El'Gygytgyn, NE Russia. Thesis.

Wilkie, K., Chaplignin, B., Meyer, H., Burns, S., Petsch, S. and Brigham-Grette, J. (2012) Modern isotope hydrology and controls on  $\delta D$  of plant leaf waxes at Lake El'gygytgyn, NE Russia. *Climate of the Past Discussions* 8, 3719-3764.

Williams, M., Rastetter, E.B., Van der Pol, L. and Shaver, G.R. (2014) Arctic canopy photosynthetic efficiency enhanced under diffuse light, linked to a reduction in the fraction of the canopy in deep shade. *New Phytol.* 202, 1267-1276.

Yang, H. and Huang, Y. (2003) Preservation of lipid hydrogen isotope ratios in Miocene lacustrine sediments and plant fossils at Clarkia, northern Idaho, USA. *Org. Geochem.* 34, 413-423.

Yang, H., Liu, W., Leng, Q., Hren, M.T. and Pagani, M. (2011) Variation in n-alkane  $\delta D$  values from terrestrial plants at high latitude: Implications for paleoclimate reconstruction. *Org. Geochem.* 42, 283-288.

Yang, H., Pagani, M., Briggs, D.E., Equiza, M., Jagels, R., Leng, Q. and LePage, B.A. (2009) Carbon and hydrogen isotope fractionation under continuous light: implications for paleoenvironmental interpretations of the High Arctic during Paleogene warming. *Oecologia* 160, 461-470.

8. Tables:

	<i>n</i> -acids		<i>n</i> -alkanes	
	C <sub>20</sub> -C <sub>22</sub> CPI	C <sub>20</sub> -C <sub>22</sub> ACL	C <sub>21</sub> -C <sub>23</sub> CPI	C <sub>21</sub> -C <sub>23</sub> ACL
<i>B. nana</i>	15.8 ± 2.7 (9)	25.9 ± 0.4 (10)	11.9 ± 3.7 (10)	28.8 ± 0.4 (10)
<i>E. vaginatum</i>	23.7 ± 4.3 (18)	26.5 ± 0.5 (18)	11.7 ± 0.7 (18)	28.5 ± 0.2 (18)

Table 1. Chain length distributions for *E. vaginatum* and *B. nana*. Uncertainties represent 1 standard deviation of the replicate plant specimens, while the numbers in parentheses are the numbers of samples analyzed.

	<i>n</i> -acids				<i>n</i> -alkanes			
	C <sub>24</sub>	C <sub>26</sub>	C <sub>28</sub>	C <sub>30</sub>	C <sub>25</sub>	C <sub>27</sub>	C <sub>29</sub>	C <sub>31</sub>
Hydrogen								
Standards		2.8 (41)	2.7 (43)		1.9 (22)	1.2 (23)	1.5 (24)	
Samples	2.3 (97)	3.1 (98)	3.0 (107)	2.0 (62)	3.2 (52)	1.8 (57)	3.6 (80)	4.9 (81)
Carbon								
Standards		0.1 (20)	0.1 (20)		0.1 (25)	0.2 (25)	0.3 (25)	
Samples	0.5 (62)	0.3 (40)	0.4 (56)	0.2 (38)	0.1 (22)	0.3 (60)	0.1 (63)	0.2 (44)

Table 2. Pooled standard deviation (1σ) and number of injections for δD<sub>wax</sub> and δ<sup>13</sup>C<sub>wax</sub> analytical measurements. Error is calculated at daily time steps and the pooled standard deviation

s is calculated as  $\sigma = \sqrt{\frac{\sum((N_i-1)*s_i^2)}{\sum N_i}}$ , where N<sub>i</sub> is the number of replicate standards for each day i, or the number of replicate injections for each sample i.

	<i>B. nana</i> (n=5)		<i>E. vaginatum</i> (n=9)	
	DL	CL	DL	CL
<i>n</i> -acids				
C <sub>20</sub>	-124.6 ± 9.0	na	-170.8 ± 8.1	-161.6 ± 6.1
C <sub>22</sub>	-122.9 ± 9.3	-147.1 ± 1.8	-167.8 ± 7.3	-160.5 ± 8.3
C <sub>24</sub>	-116.4 ± 4.8	-129.7 ± 1.1	-171.2 ± 8.7	-159.8 ± 4.9
C <sub>26</sub>	-112.4 ± 4.2	-120.1 ± 2.7	-149.8 ± 7.2	-145 ± 5.3
C <sub>28</sub>	-103.7 ± 6.5	-106.5 ± 1.6	-170.7 ± 4.9	-165.8 ± 6.5
C <sub>30</sub>	na	na	-175.8 ± 6.4	-170.3 ± 5
<i>n</i> -alkanes				
C <sub>25</sub>	-110 ± 2.5	-116.5 ± 1.3	-131.6 ± 3.3	-123.3 ± 4.6
C <sub>27</sub>	-109.4 ± 2.3	-121.3 ± 1.6	-147.1 ± 3.9	-140.8 ± 3.8
C <sub>29</sub>	-100.3 ± 2.5	-110.6 ± 3.1	-157.4 ± 5.7	-148.9 ± 6.9
C <sub>31</sub>	-102.5 ± 3.5	-111.9 ± 2	-158.1 ± 4	-149.4 ± 5.4

Table 3. Net apparent fractionation ( $\epsilon_{app}$ ) for *B. nana* and *E. vaginatum* under different light conditions. DL = Diurnal Light; CL = Continuous Light. Fractionations are calculated relative to irrigation water with a  $\delta D$  value of -39‰, and the uncertainties represent 1 $\sigma$  based on replicate plant specimens.

	<i>B. nana</i> (n=5)		<i>E. vaginatum</i> (n=9)	
	DL	CL	DL	CL
<i>n</i> -acids				
C <sub>20</sub>	-37.5 ± 1.2	-39.7 ± 1.2	-38.2 ± 0.7	-38.9 ± 1.0
C <sub>22</sub>	-35.5 ± 1.2	-39.2 ± 0.8	-37.3 ± 0.7	-38.3 ± 1.1
C <sub>24</sub>	-36.6 ± 0.6	-38.5 ± 0.8	-38.4 ± 0.6	-39.2 ± 1.1
C <sub>26</sub>	-34.9 ± 0.8	-37.2 ± 0.5	-36.6 ± 0.6	-38.0 ± 1.0
C <sub>28</sub>	-35.8 ± 1.0	-37.4 ± 0.5	-36.6 ± 0.7	-37.3 ± 1.0
C <sub>30</sub>	-35.5	-37.1 ± 1.4	-36.6 ± 0.8	-37.2 ± 1.1
<i>n</i> -alkanes				
C <sub>25</sub>	-36.2 ± 0.5	-40.4 ± 0.6		-38.6 ± 0.7
C <sub>27</sub>	-39.4 ± 0.7	-42.3 ± 0.7	-41.1 ± 0.5	-42.3 ± 1.2
C <sub>29</sub>	-38.1 ± 1.1	-40.5 ± 0.7	-38.1 ± 0.6	-39.2 ± 1.0
C <sub>31</sub>	-37.5 ± 1.1	-40.5 ± 0.6	-38.1 ± 0.5	-39.1 ± 1.2

Table 4. Carbon isotope ratios ( $\delta^{13}C_{wax}$ ) for *B. nana* and *E. vaginatum* grown under different light conditions. DL = Diurnal Light; CL = Continuous Light. The uncertainties represent 1 $\sigma$  based on replicate plant specimens.

## 9. Figures

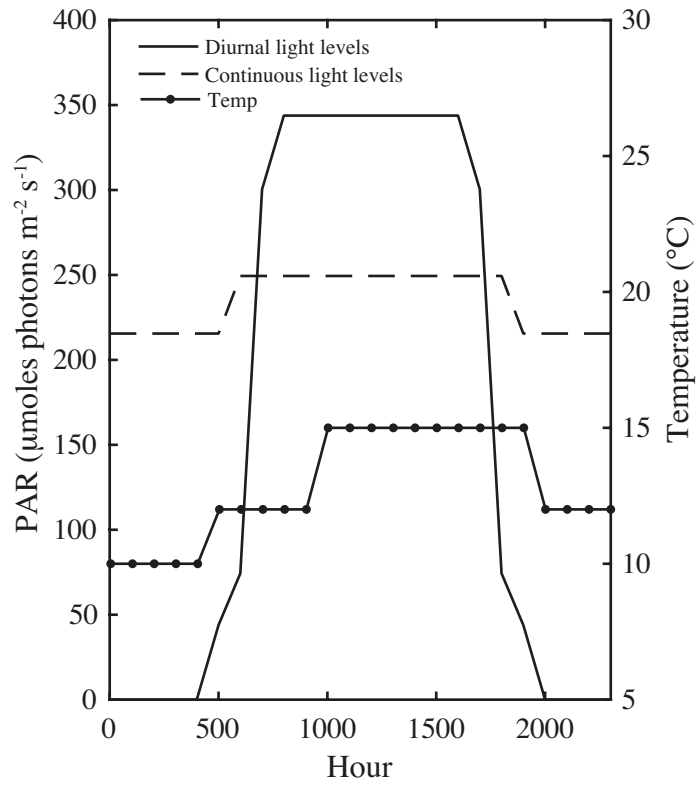


Figure 1. Growth chamber light and temperature conditions.



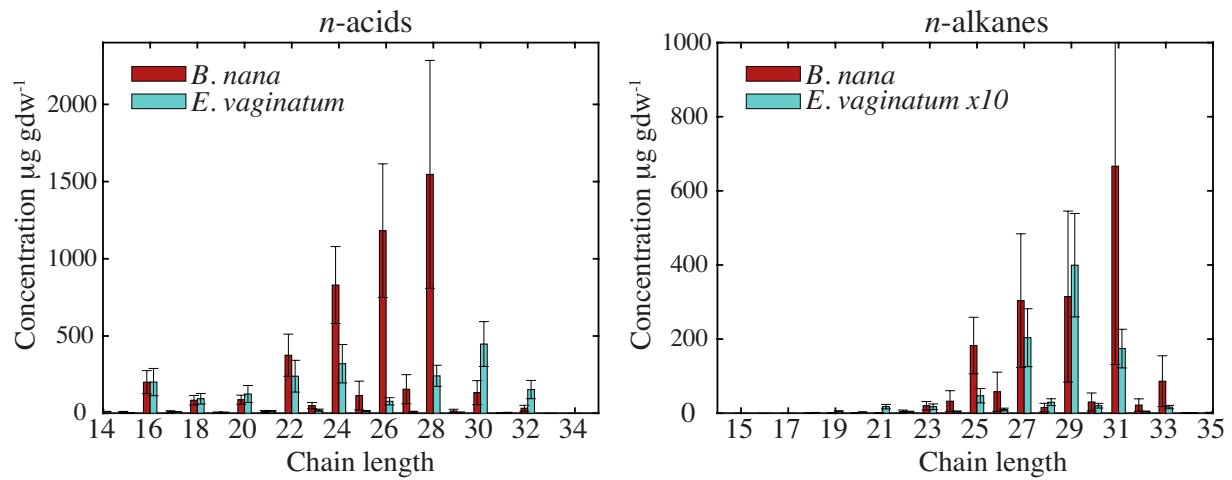


Figure 2. The *n*-alkane and *n*-acid distributions for *B. nana* and *E. vaginatum*. No differences in wax distributions were observed between light treatments, and so data represent an average of both treatments. Note scale difference between *n*-acids and *n*-alkanes. Error bars represent  $1\sigma$  uncertainty of the concentrations based on replicate plant specimens ( $n=5$  for *B. nana* and  $n=9$  for *E. vaginatum*).

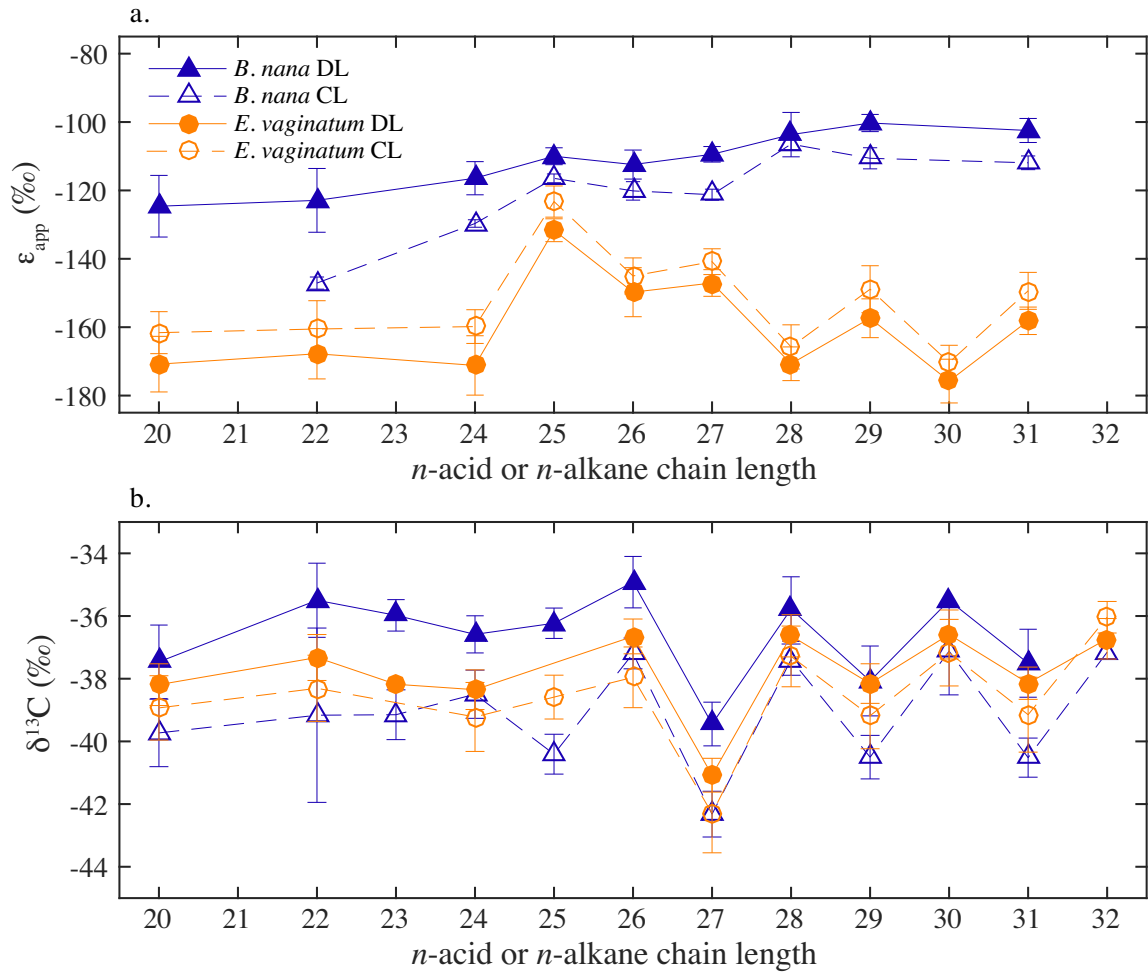


Figure 3. a) Net apparent fractionation between irrigation water ( $-39.0\text{‰}$ ) and leaf waxes under diurnal light (closed symbols) or continuous light (open symbols) for *Betula nana* (blue) and *Eriophorum vaginatum* (orange). Even-numbered chain lengths represent *n*-acids and odd-numbered chain lengths are *n*-alkanes. b) Carbon isotope values of leaf waxes. Symbols and colors are the same as in a. Error bars represent  $1\sigma$  uncertainty of  $\delta D_{wax}$  or  $\delta^{13}C$  based on replicate plant specimens ( $n=5$  for *B. nana* and  $n=9$  for *E. vaginatum*).

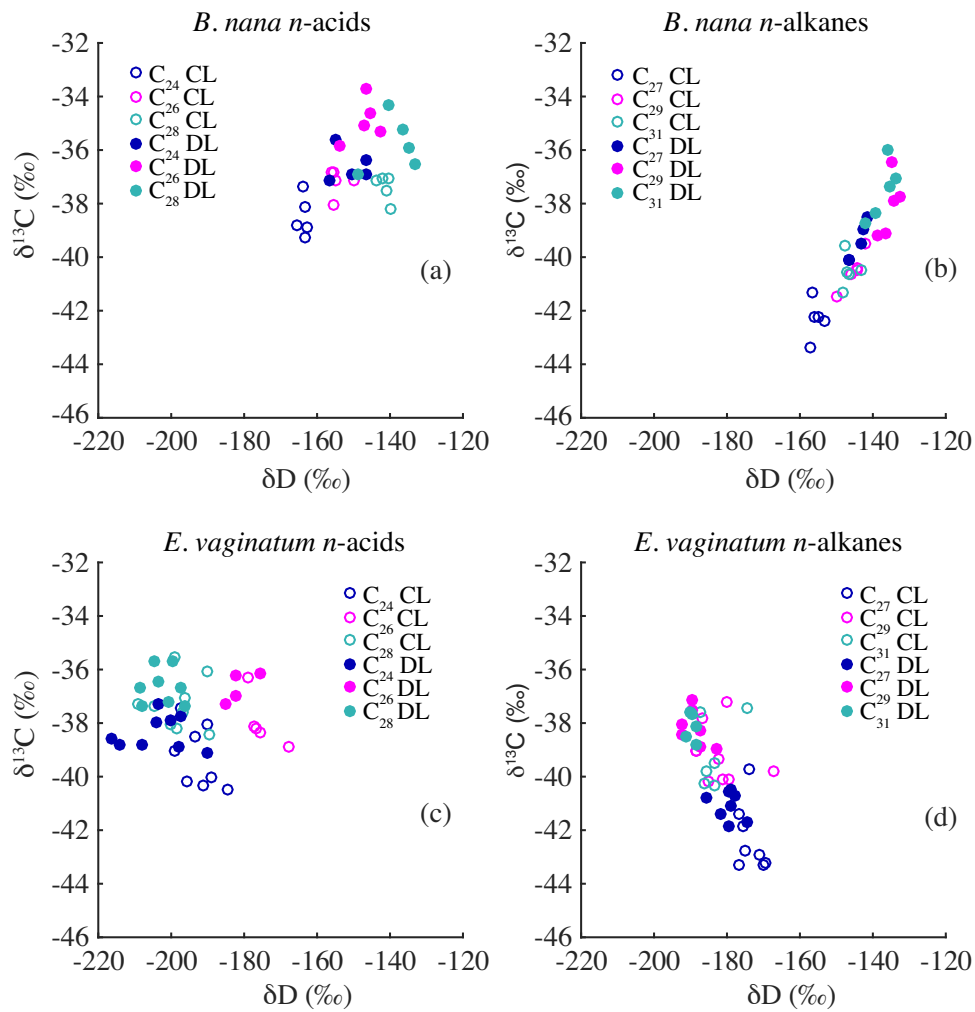


Figure 4.  $\delta^{13}\text{C}_{\text{wax}}$  vs  $\delta\text{D}_{\text{wax}}$  for *B. nana* (a, b) and *E. vaginatum* (c, d) and for *n*-acids (a, c) and *n*-alkanes (b, d). Filled circles are from the diurnal light treatment (DL), open circles are from the continuous light treatment (CL), and colors represent different chain length homologues.

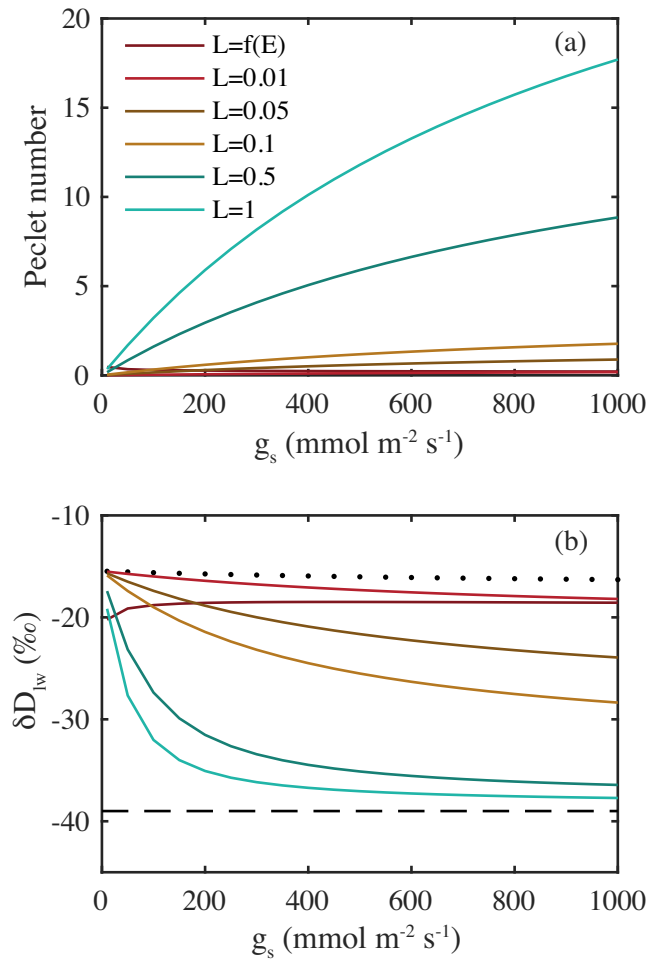


Figure 5. Modeled values of the Péclet number (a) and leaf water D/H ratios (b), as a function of stomatal conductance ( $g$ ) and leaf effective path length ( $L$ ).  $L$  is shown in units of meters, and in the first case,  $L$  is calculated as a function of transpiration rate ( $E$ ,  $\text{mmol m}^{-2} \text{s}^{-1}$ ) with the equation  $L = 2.36 \times 10^{-3} \cdot E^{-1.2}$  (Song et al. (2013)). Colors in panel b are as in panel a, with the dotted line representing the  $\delta D_w$  at the site of evaporation and the dashed line representing the  $\delta D$  of the irrigation water.

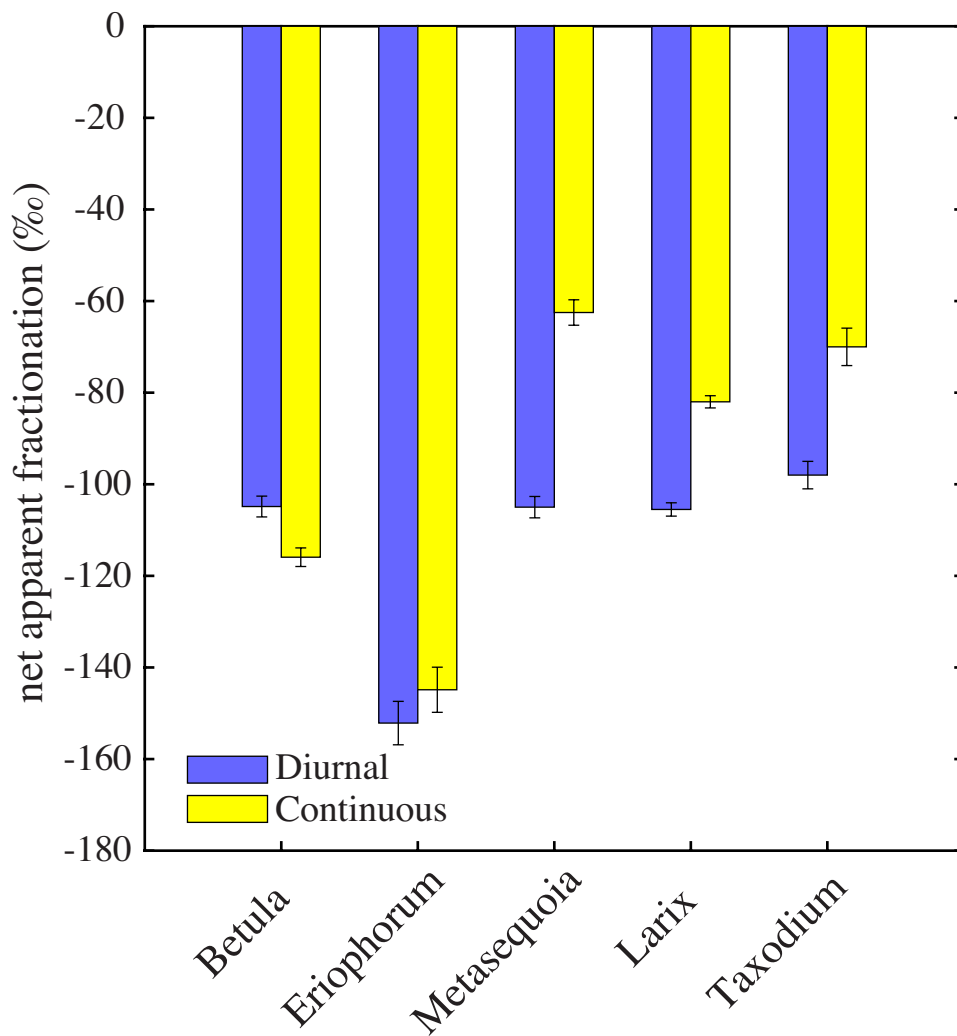


Figure 6. A comparison of net apparent fractionation ( $\epsilon_{app}$ ) of five plant species grown under diurnal and continuous light conditions. Data for *Metasequoia*, *Larix*, and *Taxodium* are from Yang et al. (2009). Data represent averages of  $C_{27}$  and  $C_{29}$  *n*-alkanes.