

1 Title: Experimental investigation on the controls of clumped isotopologue and hydrogen isotope  
2 ratios in microbial methane

3

4 Authors and affiliations:

5 Danielle S. Gruen<sup>1,2,\*</sup>, David T. Wang<sup>1,2,+</sup>, Martin Könneke<sup>3</sup>, Begüm D. Topçuoğlu<sup>4</sup>, Lucy C.  
6 Stewart<sup>4,^</sup>, Tobias Goldhammer<sup>3,5</sup>, James F. Holden<sup>4</sup>, Kai-Uwe Hinrichs<sup>3</sup>, and Shuhei Ono<sup>1,\*</sup>

7

8 <sup>1</sup>Department of Earth, Atmospheric and Planetary Sciences, Massachusetts Institute of  
9 Technology, Cambridge, Massachusetts 02139, USA.

10 <sup>2</sup>Marine Chemistry and Geochemistry Department, Woods Hole Oceanographic Institution,  
11 Woods Hole, Massachusetts 02543, USA.

12 <sup>3</sup>MARUM Center for Marine Environmental Sciences and Department of Geosciences,  
13 University of Bremen, Bremen D-28359, Germany.

14 <sup>4</sup>Department of Microbiology, University of Massachusetts, Amherst, Massachusetts 01003,  
15 USA.

16 <sup>5</sup>Leibniz-Institute for Freshwater Ecology and Inland Fisheries, Department of Chemical  
17 Analytics and Biogeochemistry, Mueggelseedamm 301, 12587 Berlin, Germany

18 <sup>+</sup>Present address: ExxonMobil Upstream Research Company, Spring, TX 77389, USA.

19 <sup>^</sup>Present address: GNS Science, Lower Hutt 5010, New Zealand.

20

21

22 \*To whom correspondence should be addressed:

23 Danielle Gruen (dgruen@mit.edu),

24 Shuhei Ono (sono@mit.edu)

25 Department of Earth, Atmospheric, and Planetary Sciences

26 Massachusetts Institute of Technology

27 77 Massachusetts Ave, Cambridge, MA 02139

28

29

30

31

32

33

### Abstract

34       The abundance of methane isotopologues with two rare isotopes (e.g.,  $^{13}\text{CH}_3\text{D}$ ) has been  
35 proposed as a tool to estimate the temperature at which methane is formed or thermally  
36 equilibrated. It has been shown, however, that microbial methane from surface environments and  
37 from laboratory cultures is characterized by low  $^{13}\text{CH}_3\text{D}$  abundance, corresponding to  
38 anomalously high apparent  $^{13}\text{CH}_3\text{D}$  equilibrium temperatures. We carried out a series of batch  
39 culture experiments to investigate the origin of the non-equilibrium signals in microbial methane  
40 by exploring a range of metabolic pathways, growth temperatures, and hydrogen isotope  
41 compositions of the media. We found that thermophilic methanogens  
42 (*Methanocaldococcus jannaschii*, *Methanothermococcus thermolithotrophicus*, and  
43 *Methanocaldococcus bathoardescens*) grown on  $\text{H}_2+\text{CO}_2$  at temperatures between 60 and 80°C  
44 produced methane with  $\Delta^{13}\text{CH}_3\text{D}$  values (defined as the deviation from stochastic abundance) of  
45 0.5 to 2.5‰, corresponding to apparent  $^{13}\text{CH}_3\text{D}$  equilibrium temperatures of 200 to 600°C.  
46 Mesophilic methanogens (*Methanosarcina barkeri* and *Methanosarcina mazei*) grown on  
47  $\text{H}_2+\text{CO}_2$ , acetate, or methanol produced methane with consistently low  $\Delta^{13}\text{CH}_3\text{D}$  values, down to  
48 -5.2‰. Closed system effects can explain part of the non-equilibrium signals for methane from  
49 thermophilic methanogens. Experiments with *M. barkeri* using D-spiked water or D-labeled  
50 acetate ( $\text{CD}_3\text{COO}^-$ ) indicate that 1.6 to 1.9 out of four H atoms in methane originate from water,  
51 but  $\Delta^{13}\text{CH}_3\text{D}$  values of product methane only weakly correlate with the D/H ratio of medium water.  
52 Our experimental results demonstrate that low  $\Delta^{13}\text{CH}_3\text{D}$  values are not specific to the metabolic  
53 pathways of methanogenesis, suggesting that they could be produced during enzymatic reactions  
54 common in the three methanogenic pathways, such as the reduction of methyl-coenzyme M.  
55 Nonetheless C-H bonds inherited from precursor methyl groups may also carry part of non-  
56 equilibrium signals.

57

## 58 1. Introduction

59 Methane is significant to the global carbon cycle (e.g., Alperin and Blair, 1992), a potent  
60 greenhouse gas (e.g., Wecht and Jacob, 2014), a source of energy (e.g., Whiticar, 1990), and a  
61 potential biosignature both for the deep biosphere (e.g., Inagaki et al., 2015) and planetary missions  
62 (e.g., Webster et al., 2014). The greatest natural source of methane to the atmosphere is produced  
63 by microbes in anoxic environments such as swamps, sediments, rice paddies, and ruminant tracts  
64 (Klapp et al., 2010). Microbial methanogenesis also contributes the majority of methane to oceanic  
65 gas hydrates, the largest reservoir of methane on Earth (Kvenvolden, 1993; Thauer et al., 2008).

66 Three major pathways are known for microbial methanogenesis (Thauer, 1998):



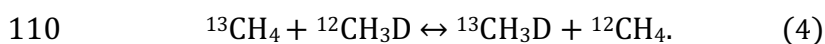
70 Hydrogenotrophic methanogenesis (1) is the reduction of  $\text{CO}_2$  with  $\text{H}_2$  as an electron donor,  
71 which is employed by at least six of the seven known orders of methanogens (Baptiste et al.,  
72 2005). Although not as common, methanogens capable of hydrogenotrophic methanogenesis can  
73 use formate as both a source of carbon and as electron donor (Baptiste et al., 2005). Acetoclastic  
74 methanogenesis (2), the disproportionation of acetate to  $\text{CH}_4$  and  $\text{CO}_2$ , is restricted to the genera  
75 *Methanosarcina* and *Methanosaeta* within the order Methanosarcinales. Hydrogenotrophic and  
76 acetoclastic methanogenesis are the two most common forms of microbial methanogenesis today  
77 (Conrad, 2005). Methylotrophic methanogenesis (3), used by Methanosarcinales and  
78 Methanomassiliicoccales, uses methyl compounds (e.g., methanol, methyl amines,  
79 dimethylsulfide, or methylthiols) as substrates (Baptiste et al., 2005; Penger et al., 2012).  
80 Nonetheless, all methanogenic archaea possess the enzyme methyl-coenzyme M reductase (MCR)  
81 which catalyzes the final step of methanogenesis (e.g., Ermler et al., 1997; Grabarse et al., 2000;  
82 Scheller et al., 2013; Wagner et al., 2016).

83 Carbon ( $^{13}\text{C}/^{12}\text{C}$ ) and hydrogen (D/H) isotope ratios have been widely used to identify the  
84 origin of methane in the environment (Blair and Carter, 1992; Whiticar, 1999; Conrad et al., 2009;  
85 McCalley et al., 2014; Blaser and Conrad, 2016). However, it is often challenging to accurately  
86 determine the methane source since the isotopic composition of methane depends upon carbon and  
87 hydrogen sources as well as isotope fractionation during formation processes (Waldron et al.,  
88 1999; Valentine et al., 2004; Yoshioka et al., 2008; Kawagucci et al., 2014).

89 Previous culture experiments using D-labeled water (Daniels et al., 1980) or D-spiked water  
90 (Kawagucci et al., 2014) indicate that hydrogen in hydrogenotrophic methane is primarily derived  
91 from the hydrogen in water with only minor contribution from hydrogen in hydrogen gas ( $\text{H}_2$ ). The  
92 contribution from  $\text{H}_2$  can be explained by the production of metabolic water (Sugimoto and Wada

93 1995) since the production of one mole of methane yields two moles of water  
94 ( $\text{CO}_2+4\text{H}_2\rightarrow\text{CH}_4+2\text{H}_2\text{O}$ ), and the high specific rate of methanogenesis results in rapid turnover of  
95 intracellular water. The residence time of intracellular water is estimated to be as short as a few  
96 seconds (much lower than the doubling time of cells) during exponential growth (Kawagucci et  
97 al., 2014). The  $\delta\text{D}$  value of intracellular water can also be influenced by exchange between  $\text{H}_2$   
98 and  $\text{H}_3\text{O}^+$ , which can be catalyzed by hydrogenase enzymes (Burke, 1993; Valentine et al., 2011).  
99 Alternatively, the direct transfer of hydrogen in  $\text{H}_2$  into  $\text{CH}_4$  can be mediated by the enzyme  
100 methylenetetrahydromethanopterin dehydrogenase (Schworer et al., 1993; Schleucher et al., 1994;  
101 Klein et al., 1995a,b; Hartmann et al., 1996). In contrast, experiments with acetoclastic  
102 methanogens indicate that up to three out of four hydrogen atoms in methane are derived from the  
103 methyl group of acetate, as implied in the stoichiometry of Reaction (2) (Pine and Barker, 1956).

104 In addition to the ratios of  $^{13}\text{C}/^{12}\text{C}$  and  $\text{D}/\text{H}$  of methane, measurements of the doubly isotope  
105 substituted isotopologue,  $^{13}\text{CH}_3\text{D}$  and/or  $^{12}\text{CH}_2\text{D}_2$ , have recently been applied as tools to constrain  
106 the source of methane in a variety of environments (Stolper et al., 2013; Stolper et al., 2014; Wang  
107 et al., 2015; Douglas et al., 2016; Wang et al., 2016; Young et al., 2016; Douglas et al., 2017;  
108 Whitehill et al., 2017; Young et al., 2017). The following reaction shows the equilibrium among  
109 four methane isotopologues, including  $^{13}\text{CH}_3\text{D}$ :



111 Here, its equilibrium constant ( $K_T$ ) can be written as:

$$112 \quad K_T = \frac{[^{13}\text{CH}_3\text{D}][^{12}\text{CH}_4]}{[^{12}\text{CH}_3\text{D}][^{13}\text{CH}_4]}. \quad (5)$$

113  
114 The value of  $K_T$  primarily depends on temperature, and it approaches unity at high temperatures  
115 (1.0002 at 1,000°C), but is about 1.0057 at 25°C as estimated by molecular simulations (Ma et al.,  
116 2008; Stolper et al., 2015; Wang et al., 2015; Liu and Liu, 2016). Thus, the precise measurements  
117 of four isotopologues' abundance ( $^{12}\text{CH}_4$ ,  $^{13}\text{CH}_4$ ,  $^{12}\text{CH}_3\text{D}$ , and  $^{13}\text{CH}_3\text{D}$ ) were thought to provide  
118 an estimate of the temperature at which the methane gas was formed or thermally equilibrated.  
119 Initial studies using high-resolution mass-spectrometry demonstrated that this new isotopologue  
120 thermometer provides a range of temperatures that are consistent with formation temperatures for  
121 methane samples from geologic environments, such as natural gas reservoirs (Stolper et al., 2014).  
122 Later studies, however, showed that methane sampled from surface environments (e.g., ruminants,  
123 lakes, and swamps) is characterized by clear non-equilibrium signals that yield apparent clumped  
124 isotopologue temperatures higher than environmental methane generation temperatures (Stolper et  
125 al., 2015; Wang et al., 2015; Douglas et al., 2016; Young et al., 2017).

126           These studies also showed that the degree of methane isotopologue disequilibrium is  
127 correlated with D/H-isotope disequilibrium between H<sub>2</sub>O and CH<sub>4</sub> (i.e., CH<sub>3</sub>D + H<sub>2</sub>O ↔ CH<sub>4</sub> +  
128 HDO). To explain this observed relationship, Wang et al. (2015), and Stolper et al. (2015) both  
129 presented a mathematical model that considered metabolic reversibility, which is defined as the  
130 ratio of backward to forward fluxes through an enzymatically-mediated reaction sequence. These  
131 models were based on earlier models for sulfur isotope effects of sulfate reducers (Rees, 1973;  
132 Farquhar et al., 2007; Sim et al., 2011). By choosing the appropriate fractionation factors, these  
133 models can describe isotopologue compositions of microbial methane between kinetic and  
134 equilibrium end-members, corresponding to unidirectional and reversible reactions, respectively.  
135 Accordingly, both studies attributed the origin of kinetic clumped isotope signals intrinsic to one  
136 or more enzymatic reactions in the methanogenic pathways. The application of transition state  
137 theory (Bigeleisen, 1949) can explain <sup>13</sup>CH<sub>3</sub>D abundance between equilibrium and stochastic  
138 ( $\Delta^{13}\text{CH}_3\text{D} > 0\text{‰}$ ), but anti-clumped <sup>13</sup>CH<sub>3</sub>D abundance ( $\Delta^{13}\text{CH}_3\text{D} < 0\text{‰}$ ) requires the mixing of  
139 methane reservoirs with an often unreasonably large range of bulk  $\delta\text{D}$  and  $\delta^{13}\text{C}$  values or a physical  
140 mechanism, including quantum mechanical tunneling (Wang et al., 2015; Whitehill et al., 2017;  
141 Young et al., 2017) ( $\Delta^{13}\text{CH}_3\text{D}$  is a measure of excess <sup>13</sup>CH<sub>3</sub>D as defined later in equation (8)). For  
142 the doubly deuterated isotopologue CH<sub>2</sub>D<sub>2</sub>, purely statistical combinational-effects can also  
143 produce large apparent depletions in CH<sub>2</sub>D<sub>2</sub> (Yeung et al., 2016, Young et al., 2017).

144           Recent work cultivating methanogens produced isotopologue compositions consistently  
145 out of isotopic equilibrium (Douglas et al., 2016; Young et al., 2017). These results were in  
146 agreement with previous culture studies (Stolper et al., 2015; Wang et al., 2015), but highlighted  
147 the need for further assessment of the mechanisms that control microbial <sup>13</sup>CH<sub>3</sub>D compositions. In  
148 particular, the source of H in CH<sub>4</sub> for acetoclastic and methylotrophic methanogenesis remained  
149 uncertain (Douglas et al., 2016, 2017). The goal of this work is to better characterize the kinetic  
150 <sup>13</sup>CH<sub>3</sub>D effects that lead to these generally low  $\Delta^{13}\text{CH}_3\text{D}$  values, specifically during microbial  
151 methanogenesis, using a comprehensive set of metabolic pathways and temperatures. We  
152 investigated this with a series of batch culture experiments to test the effect of 1) species  
153 (*Methanothermococcus thermolithotrophicus*, *Methanocaldococcus jannaschii*,  
154 *Methanocaldococcus bathoardescens*, *Methanosarcina barkeri*, and *Methanosarcina mazei*), 2)  
155 temperature (from 30 to 85°C), and 3) substrate (H<sub>2</sub>+CO<sub>2</sub>, acetate, and methanol). We also  
156 investigated the effect of closed-system processes as well as D/H ratios of medium water to test if  
157 apparent high-temperature signals are produced by mixing of two or more pools of methane (or its  
158 precursors), as mixing has been shown to produce a bias in the clumped isotopologue temperature  
159 estimate (Stolper et al., 2015; Wang et al., 2015; Douglas et al., 2016).

## 160 2. Methods

### 161 2.1 Laboratory Culture Experiments

162 Table 1 summarizes all culture experiments conducted in this study as well as results from  
163 our earlier experiments presented in Wang et al., (2015). Descriptions of specific experimental  
164 conditions are provided below.

#### 166 2.1.1 Temperature Series Experiments

167 Pure cultures of methanogens were grown in duplicate in batch cultures at a range of  
168 temperatures (30 to 85°C). Three different hydrogenotrophic methanogens were selected based on  
169 their growth kinetics and optimum growth temperatures:  
170 *Methanothermococcus thermolithotrophicus*, *Methanocaldococcus jannaschii*, and  
171 *Methanocaldococcus bathoardescens*. Cultures of *M. thermolithotrophicus* and *M. jannaschii*  
172 were purchased from the German Collection of Microorganisms and Cell Cultures (DSMZ,  
173 Braunschweig, Germany). *M. bathoardescens* was originally isolated from vent fluid at Axial  
174 Volcano, Juan de Fuca Ridge, and maintained in culture at the University of Massachusetts,  
175 Amherst (Stewart et al., 2015).

176 Culture medium was prepared following the recipe for DSMZ medium 282 according to  
177 Stewart et al. (2015). The headspace was filled with H<sub>2</sub>:CO<sub>2</sub> (in a ratio of 80:20 by volume) at 2  
178 bar absolute pressure. For each experiment, 5 mL of inoculum from a culture in the exponential  
179 growth phase was added to a sample vial containing 50 mL media. *M. thermolithotrophicus* was  
180 grown at 30, 40, 50, and 60°C, *M. jannaschii* was grown at 70 and 80°C, and *M. bathoardescens*  
181 was grown at 85°C. All cultures were incubated in 140 mL rubber-stoppered glass serum vials in  
182 forced-air convection ovens. Cell concentrations were monitored by cell counts with a Petroff-  
183 Hauser counting chamber and phase-contrast light microscope to determine the growth kinetics as  
184 a function of temperature (Stewart et al., 2016). Experiments for isotope measurement were  
185 stopped at a time when stationary phase was reached (5 to 64 hours, as measured in prior studies  
186 and replicated in our laboratory (Huber et al., 1982, Jones et al., 1983, Ver Eecke et al., 2013);  
187 Table 2). Most CO<sub>2</sub> (>95%) was converted to CH<sub>4</sub> as indicated by gas chromatography (GC)  
188 measurements of carbon dioxide and methane in the headspace gas.

#### 189 2.1.2 Time Series Experiments

190 *M. bathoardescens* was grown under an H<sub>2</sub>:CO<sub>2</sub> (80:20) atmosphere in replicate batch  
191 cultures at 80°C to study the effects of growth phase and closed system on <sup>13</sup>CH<sub>3</sub>D and to estimate  
192 instantaneous clumped isotopologue fractionation factors. Culture medium was prepared as above  
193 (Sec. 2.1.1). Methane was sampled and analyzed by GC from batch cultures at time points

194 corresponding to fractional conversion of 6, 10, 68, and 77% of the initial carbon dioxide to  
195 methane (Table 1). The fractional conversion was calculated by dividing the volume of methane  
196 produced at the conclusion of the experiment by the volume of methane expected.

### 197 **2.1.3 Substrate Series Experiments**

198 To determine the effect of different metabolic pathways, established batch culture  
199 incubations of *Methanosarcina barkeri* were grown on three different substrates: H<sub>2</sub>+CO<sub>2</sub>,  
200 (referred to as hydrogenotrophic cultures hereafter), methanol (methylotrophic cultures), and  
201 acetate (acetoclastic cultures). Cultures of *M. barkeri* (strain DSM-800) were purchased from the  
202 DSMZ (Braunschweig, Germany). The growth medium was prepared according to the recipe for  
203 DSMZ medium 120 (Balch et al., 1979). For hydrogenotrophic cultures, the headspace was filled  
204 with 1.5 absolute bar of H<sub>2</sub>:CO<sub>2</sub> (80:20) gas mix. For acetoclastic and methylotrophic cultures, the  
205 headspace was filled with 1.5 absolute bar of N<sub>2</sub>:CO<sub>2</sub> (70:30) gas mix and the medium was  
206 amended with 30 mM of Na-acetate or 250 mM of methanol, respectively. Cultures were incubated  
207 in duplicate near room temperature. Two sets of experiments were carried out. The first set of  
208 experiments (Set 1) was intended to provide preliminary data, and thus temperature was not strictly  
209 controlled over the course of the experiment (cultures exposed to ambient temperatures between  
210 21 and 38°C), and the medium contained yeast extract. Nonetheless, all bottles in this series were  
211 subjected to identical environmental conditions. A second set of cultures (Set 2) was prepared and  
212 incubated under close monitoring at constant temperature (38°C). Yeast extract (YE) and casitone  
213 were omitted from medium unless otherwise noted. At the end of the experiment, cultures were  
214 killed with 1M NaOH to prevent any additional methanogen activity as described in Methods  
215 Section 2.2.

### 216 **2.1.4. D-label and d-spike Experiments**

217 To constrain the source of hydrogen in the hydrogenotrophic, methylotrophic, and  
218 acetoclastic pathways, a subset of *M. barkeri* cultures was also spiked with either 15 or 30 μL of  
219 D<sub>2</sub>O per one liter of media. Additionally, acetoclastic cultures were prepared containing 10, 50, or  
220 100% (molar fraction) deuterated acetic acid (CD<sub>3</sub>COOD, 99% purity, Sigma-Aldrich, St. Louis,  
221 MO).

## 223 **2.2 Sample Preparation and Isotopologue measurements**

224 At the completion of an experiment, 1 M NaOH was injected in each culture bottle (at a  
225 ratio of 1 mL per 10 mL of medium) to sacrifice the culture and to draw down CO<sub>2</sub> pressures in  
226 the headspace. Methane samples from culture experiments were measured within one year of the  
227 completion of culture experiments. Repeated measurements of NaOH-treated samples did not

228 show measureable changes in isotope or isotopologue ratios during storage. The headspace was  
229 sampled by flushing with helium via two needles. Methane gas was purified from culture gas  
230 mixtures (mostly methane, hydrogen, and nitrogen) using an automated preparative gas  
231 chromatography system as previously described (Wang et al., 2015). For most analyses,  
232 approximately 10 mL STP of methane was used.

233 The abundance of isotopologues in methane samples was measured by a tunable infrared  
234 laser direct absorption spectrometer (TILDAS) that measures absorption in the infrared region of  
235 the electromagnetic spectrum corresponding to bending vibrations of C–H and C–D bonds (Ono  
236 et al., 2014; Wang et al., 2015). A typical measurement consists of eight to ten cycles of alternating  
237 measurements of reference and sample methane. Measured isotopologue ratios were averaged and  
238 95% confidence intervals were calculated according to Student’s *t*-distribution as previously  
239 described (Wang et al., 2015).

240 The hydrogen isotope composition of culture medium water ( $\delta D_{H_2O}$ ), except for that of  
241 *Methanosarcina* Set 2 experiments, was measured using a cavity ring-down spectrometer (CRDS,  
242 Picarro Inc., Santa Clara, California, USA) at the University of Massachusetts, Amherst. The  
243  $\delta D_{H_2O}$  values of the *Methanosarcina* Set 2 cultures were measured at the University of Bremen  
244 also by CRDS (Picarro L2130-*i* Analyzer (Picarro Inc., Santa Clara, CA, USA). The hydrogen  
245 isotope composition of H<sub>2</sub> was not measured. Bulk  $\delta^{13}C$  of methanol and acetate were measured  
246 via LC-IRMS at the University of Bremen (Heuer et al., 2006). CO<sub>2</sub> in the N<sub>2</sub>:CO<sub>2</sub> and H<sub>2</sub>:CO<sub>2</sub>  
247 gas mixes, bicarbonate solution, and culture media were measured via isotope ratio infrared  
248 spectrometry (IRIS) at the University of Bremen. The D/H ratio of sodium acetate (CH<sub>3</sub>COONa)  
249 was measured by high temperature conversion elemental analyzer interfaced with isotope ratio  
250 mass-spectrometer (IRMS) at University of Chicago. Typical uncertainties were 0.2 to 0.4‰ and  
251 2 to 5‰ for  $\delta^{13}C$  and  $\delta D$ , respectively.

252

### 253 2.3. Isotope Notation and Calibration

254 In this work, stable isotopic ratios of carbon and hydrogen are reported in conventional  $\delta$   
255 notation, defined as:

$$256 \quad \delta^{13}C = \frac{(^{13}C/^{12}C)_{\text{sample}}}{(^{13}C/^{12}C)_{\text{PDB}}} - 1 \quad (6)$$

$$257 \quad \delta D = \frac{(D/H)_{\text{sample}}}{(D/H)_{\text{SMOW}}} - 1 \quad (7)$$

258 where, PDB and SMOW are Pee Dee Belemnite and Standard Mean Ocean Water, respectively.  
259 The factor of 1000, which commonly appears in definitions of  $\delta$  values in the geochemical  
260 literature, has been omitted from Equations (6) and (7), as it is implied by the permil (‰) symbol,  
261 in accordance with IUPAC recommendations (Coplen, 2011). Values for  $\delta^{13}C$  and  $\delta D$  of methane



262 analyzed via TILDAS at MIT have been calibrated against PDB and SMOW via measurements of  
 263 natural gas standards NGS-1 and NGS-3 (Wang et al., 2015). Reference values for  $\delta^{13}\text{C}$  and  $\delta\text{D}$   
 264 were taken to be  $-29.0\text{‰}$  and  $-138\text{‰}$  for NGS-1, and  $-72.8\text{‰}$  and  $-176\text{‰}$  for NGS-3,  
 265 respectively (Hut, 1987).

266 Because the TILDAS measures ratios of methane isotopologues, bulk  $\delta^{13}\text{C}$  and  $\delta\text{D}$  values  
 267 reported in this paper are necessarily derived quantities. For samples of methane containing a mix  
 268 of isotopologues at or sufficiently close to their naturally-occurring abundances, including all  
 269 samples analyzed via TILDAS in this study, ratios of isotopologues are interchangeable with ratios  
 270 of isotopes (i.e.,  $^{13}\text{C}/^{12}\text{C}$  and  $\text{D}/\text{H}$ ) when calculating  $\delta$  values, with no difference within achievable  
 271 uncertainties of isotope ratio measurements:  $^{13}\text{C}/^{12}\text{C} \approx [^{13}\text{CH}_4]/[^{12}\text{CH}_4]$  and  $\text{D}/\text{H} \approx \frac{1}{4}$   
 272  $[^{12}\text{CH}_3\text{D}]/[^{12}\text{CH}_4]$ . Note that the symmetry factor of  $\frac{1}{4}$  cancels out when  $\delta\text{D}$  values are calculated  
 273 via equation (7).

274 The abundance of the clumped isotopologue  $^{13}\text{CH}_3\text{D}$  is reported as  $\Delta^{13}\text{CH}_3\text{D}$ , which  
 275 represents the deviation (excess) of the abundance of  $^{13}\text{CH}_3\text{D}$  from a stochastic distribution (i.e.,  
 276 one in which all carbon and hydrogen isotopes are randomly distributed amongst the isotopologues  
 277  $^{12}\text{CH}_4$ ,  $^{13}\text{CH}_4$ ,  $^{12}\text{CH}_3\text{D}$ , and  $^{13}\text{CH}_3\text{D}$ ) (Ono et al., 2014):

$$\begin{aligned}
 278 \quad \Delta^{13}\text{CH}_3\text{D} &= \frac{[^{13}\text{CH}_3\text{D}][^{12}\text{CH}_4]}{[^{13}\text{CH}_4][^{12}\text{CH}_3\text{D}]} - 1 \\
 279 \quad &\cong \ln \frac{^{13}\text{CH}_3\text{D}}{^{12}\text{CH}_4} - \ln \frac{^{13}\text{CH}_4}{^{12}\text{CH}_4} - \ln \frac{^{12}\text{CH}_3\text{D}}{^{12}\text{CH}_4}. \quad (8)
 \end{aligned}$$

280 We used the following equation to derive apparent  $\Delta^{13}\text{CH}_3\text{D}$  temperatures:

$$\Delta^{13}\text{CH}_3\text{D}(T) = -0.1101 \left( \frac{1000}{T} \right)^3 + 1.0415 \left( \frac{1000}{T} \right)^2 - 0.5223 \left( \frac{1000}{T} \right) \quad (9)$$

281 where  $T$  is in Kelvin. Density function theory (B3LYP) with 6-31G(d) basis set was used to  
 282 estimate harmonic vibrational frequencies, and isotope fractionation factors were calculated  
 283 following conventional theory by Urey (1947). A sample of methane with stochastically-  
 284 distributed abundances of isotopologues has a  $\Delta^{13}\text{CH}_3\text{D}$  value of zero, corresponding to an  
 285 apparent equilibrium temperature of infinity for Reaction (4). Negative  $\Delta^{13}\text{CH}_3\text{D}$  values represent  
 286 “anti-clumped” signals, where the abundance of  $^{13}\text{CH}_3\text{D}$  is more depleted than that expected for  
 287 stochastic isotopologue distribution.

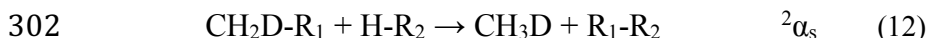
288 Measurements made via TILDAS give the abundances of methane isotopologues relative  
 289 to a reference gas against which the samples are measured (here, a commercially-sourced cylinder  
 290 of methane termed “AL1” was used as the reference gas). To express  $\Delta^{13}\text{CH}_3\text{D}$  values of samples  
 291 relative to the stochastic distribution requires determination of the  $\Delta^{13}\text{CH}_3\text{D}$  value of the reference  
 292 gas AL1. This was determined by heating AL1 in a flame-sealed glass tube in the presence of a

293 platinum catalyst between 150 and 400°C for several days to months (Ono et al., 2014; Wang et  
294 al., 2015).

295 The fractionation factor ( $\alpha$ ) quantifies the difference in the relative abundance of isotopes  
296 between the substrate and the instantaneous product of a reaction. For the reduction of carbon  
297 dioxide to methane, the fractionation factor is defined as:

$$298 \quad {}^{13}\alpha_{\text{CH}_4/\text{CO}_2} = \frac{\delta^{13}\text{C}_{\text{CH}_4} + 1}{\delta^{13}\text{C}_{\text{CO}_2} + 1}. \quad (10)$$

299 Two modes of D/H fractionation characterize each hydrogen addition step during the biosynthesis  
300 of methane. For example, for addition of H (or D) onto a methyl group:



303 Reaction (11) is accompanied by a *primary* D/H isotope effect (characterized by the fractionation  
304 factor  ${}^2\alpha_{\text{p}}$ ), where a D is substituted for H in the bond formed (or broken). Reaction (12) is  
305 accompanied by a *secondary* D/H isotope effect (with fractionation factor  ${}^2\alpha_{\text{s}}$ ), where the  
306 substitution of D for H occurs on the site *adjacent* to the C–H bond being formed (or broken) and  
307 the C–D bond are carried from reactant to product. Primary D/H isotope effects are typically much  
308 larger compared to secondary isotope effects. For the reduction of methyl-coenzyme M to methane  
309 above, the secondary isotope effect is 0.84 and the primary isotope effect of the backward reaction  
310 is 0.41 (Scheller et al., 2013).

311 According to the rule of geometric mean (Bigeleisen, 1955), the fractionation factor for the  
312 clumped isotopologue  ${}^{13}\text{CH}_3\text{D}$  is usually close, but not necessarily equal, to the product of carbon  
313 and hydrogen fractionation factors ( ${}^{13-2}\alpha \approx {}^{13}\alpha^2\alpha$ ). A significant departure from this rule has been  
314 observed for some *in vitro* enzyme assay experiments for doubly deuterated substrates, and  
315 attributed to quantum mechanical tunneling (e.g., Srinivasan et al., 1985; Amin et al., 1988; Huskey  
316 et al., 2007). We represent the departure from this relationship by the  $\gamma$  factor, which is a metric of  
317 the kinetic clumped isotope effect (Wang et al., 2015). There are two ways by which a  ${}^{13}\text{C}$ -  
318 containing methyl group can acquire an H (or D) to form  ${}^{13}\text{CH}_3\text{D}$  (analogous to Reactions 5 and  
319 6). Thus, there are two  $\gamma$  factors corresponding to primary ( $\gamma_{\text{p}}$ ) and secondary D/H isotope effects  
320 ( $\gamma_{\text{s}}$ ):

$$321 \quad {}^{13-2}\alpha_{\text{p}} = \gamma_{\text{p}} {}^{13}\alpha^2\alpha_{\text{p}} \quad \text{and} \quad (13)$$

$$322 \quad {}^{13-2}\alpha_{\text{s}} = \gamma_{\text{s}} {}^{13}\alpha^2\alpha_{\text{s}}. \quad (14)$$

323 For bond forming reactions, product methane could become anti-clumped ( $\Delta^{13}\text{CH}_3\text{D} < 0$ ) when the  
324 value of  $\gamma$  is less than unity.

## 325 **2.4 FTIR Analysis of Methane Isotopologues**

326 A Fourier transform infrared (FTIR) spectrometer (iS5, Thermo Scientific, Waltham,  
327 Massachusetts, USA) was used to quantify the mixing ratios of deuterated isotopologues of

328 methane ( $\text{CHD}_3$ ,  $\text{CH}_2\text{D}_2$  and  $\text{CH}_3\text{D}$ ) and non-deuterated methane ( $\text{CH}_4$ ), produced in acetoclastic  
329 cultures spiked with  $\text{CD}_3\text{COOD}$ . The FTIR spectrometer has a  $0.8\text{ cm}^{-1}$  spectral resolution, and is  
330 equipped with a gas cell that has a path length of 10 cm, volume of 70 mL, and windows of KBr.  
331 The cell was evacuated and filled with argon three times prior to injection of the sample or  
332 standard. For each measurement, 100  $\mu\text{L}$  to 1 mL standard temperature and pressure (STP) of the  
333 standard or sample (culture headspace, subsampled with a gas-tight syringe, Vici Valco, Houston,  
334 Texas, USA) was injected into the cell through a small inlet valve. Reference spectra were taken  
335 on samples of ordinary  $\text{CH}_4$  (containing D at natural abundance) and on pure (>98% purity)  
336 deuterated isotopologues ( $\text{CH}_3\text{D}$ ,  $\text{CH}_2\text{D}_2$ , and  $\text{CD}_3\text{H}$ ) purchased from Cambridge Isotope  
337 Laboratories (Cambridge, MA). The mixing ratio of methane isotopologues was determined by a  
338 least squares fit in the region of the absorption spectrum between 3200 and 2800  $\text{cm}^{-1}$ .

### 339 3. Results

340 Table 1 (Section 2) summarizes the tables and figures in which results for each set of  
341 methanogen culture experiments are displayed, Table 2 summarizes all data used in the figures  
342 that follow, and Table 3 summarizes the isotope composition of substrates and medium water.  
343

#### 344 3.1 Non-equilibrium $\Delta^{13}\text{CH}_3\text{D}$ signals of methane from microbial cultures

345 Microbial methane produced from pure culture experiments yielded non-equilibrium  
346 signals with corresponding apparent clumped isotope temperatures much higher than the  
347 temperatures at which the cultures were incubated (Figure 1). Overall, thermophilic methanogens  
348 (grown at  $>40^\circ\text{C}$ ) produced  $\Delta^{13}\text{CH}_3\text{D}$  values that are lower than those expected for equilibrium  
349 distribution (0.5 to 2.5‰), whereas mesophilic methanogens produced lower (mostly anti-  
350 clumped) signals (-5.2 to 1.6‰), consistent with limited measurements reported in Douglas et al.,  
351 (2016) and Young et al., (2017).

352 Thermophilic methanogens (*M. jannaschii*, *M. bathoardescens*, and *M.*  
353 *thermolithotrophicus*) grown on  $\text{H}_2+\text{CO}_2$  between 30 and  $80^\circ\text{C}$  produced methane with  $\Delta^{13}\text{CH}_3\text{D}$   
354 values ranging from 2.5 to 0.5‰, corresponding to apparent clumped isotopologue temperatures  
355 of 195 to  $603^\circ\text{C}$ , respectively. Methane produced by *Methanosarcina* (*M. barkeri* and *M. mazei*)  
356 grown on  $\text{H}_2+\text{CO}_2$ , acetate, and methanol was characterized by  $\Delta^{13}\text{CH}_3\text{D}$  values ranging from -5.2  
357 to -1.1‰, which are lower than those of methane produced by thermophilic methanogens (Figure  
358 1).

359 As described in the methods (Section 2.1.3), Set 1 cultures were exposed to ambient  
360 temperatures between 21 and 38°C and the media contained yeast extract. Set 2 cultures were  
361 prepared and incubated under close monitoring at constant temperature (38°C). Yeast extract (YE)  
362 and casitone were omitted from media unless otherwise noted (Table 2), and at the end of the  
363 experiment, cultures were killed with 1M NaOH. There are some notable differences in the  
364  $\Delta^{13}\text{CH}_3\text{D}$  values between Set 1 and Set 2 experiments for cultures with *M. barkeri* and *M. mazei*.  
365 Table 2 shows that Set 1 exhibits slightly higher  $\Delta^{13}\text{CH}_3\text{D}$  values for each substrate, most extreme  
366 for cultures grown on  $\text{H}_2+\text{CO}_2$  (as much as a 3‰ difference). *M. barkeri* cultures in Set 2 grown  
367 on  $\text{H}_2+\text{CO}_2$  at lower temperatures (21 vs. 38°C) but in the same conditions without any isotope  
368 spike exhibit higher  $\Delta^{13}\text{CH}_3\text{D}$  values (by 2.0‰).

369 Because methanogens grown on different substrates and at different temperatures exhibit  
370 different growth rates, methane was generated more quickly or slowly for some bottles. For the  
371 Time Series experiments (TS), the culture headspace reached <1%  $\text{CO}_2$  in as little as 5 hours for  
372 80°C cultures and as long as 3 days for 30°C cultures. Cultures in the Time Series experiments  
373 (CS) were stopped intervals between 2.75 and 4 hours, spanning 6-77% reaction completion (Table  
374 2). Methanogens from the Substrate, Spike, and Temperature experiments (1&2) took much longer  
375 to generate methane. Set 2 cultures required long as two months of incubation in order to produce  
376 enough methane to be sampled and analyzed.

377

### 378 **3.2 Effect of a closed system on $\delta^{13}\text{C}$ , $\delta\text{D}_{\text{CH}_4}$ , and $\Delta^{13}\text{CH}_3\text{D}$ systematics**

379 In order to test the potential bias in  $\Delta^{13}\text{CH}_3\text{D}$  values due to closed system isotope effects,  
380 methane was sampled from batch cultures in time series experiments (Table 2; Figure 2). The  $\delta^{13}\text{C}$   
381 value of methane increased from -18.0 to -3.8‰ over the course of experiments (Figure 2A). This  
382 increase is consistent with closed system isotope effects. In contrast,  $\delta\text{D}$  values of methane  
383 decreased (-350.3 to -402.3‰) over the course of the experiment (Figure 2B). Our results show  
384  $\Delta^{13}\text{CH}_3\text{D}$  values remain relatively constant over the course of the experiment between 2.1 and  
385 2.5‰ (Figure 2C).

386

### 387 **3.3 D-spiked $\text{H}_2\text{O}$ experiments**

388 As the  $\delta\text{D}$  of water is increased by spiking the media water, the  $\delta\text{D}$  of product methane  
389 also increased (Figure 3). This illustrates the uptake of hydrogen from water to form methane,  
390 consistent with previous pure culture (Yoshioka et al., 2008; Kawagucci et al., 2014; Okumura et  
391 al., 2016) and incubation experiments (Schoell, 1980; Sugimoto and Wada, 1995). Linear

392 regression of the data for hydrogenotrophic, acetoclastic, and methylotrophic cultures yielded the  
393 following relationships, respectively:

394

$$395 \quad \delta D_{CH_4} = (0.571 \pm 0.011) \delta D_{H_2O} - (423.0 \pm 2.1), \quad (15)$$

$$396 \quad \delta D_{CH_4} = (0.212 \pm 0.004) \delta D_{H_2O} - (331.5 \pm 0.8), \text{ and} \quad (16)$$

$$397 \quad \delta D_{CH_4} = (0.269 \pm 0.005) \delta D_{H_2O} - (369.7 \pm 1.2). \quad (17)$$

398

399 where intercepts are in ‰. The linear fit and standard error for the slope and intercept was  
400 calculated following (York et al., 2004) by taking into account standard errors of 0.2‰ and 5‰  
401 for  $\delta D_{CH_4}$  and  $\delta D_{H_2O}$  (assuming errors are not correlated). Hydrogenotrophic cultures yielded a  
402 higher slope (0.571) compared to methylotrophic (0.269) and acetoclastic (0.211) cultures.

403 The values of  $\Delta^{13}CH_3D$  are weakly dependent on pathways: -2.9 to -3.8‰, -4.2 to -4.9‰,  
404 and -2.4 to -3.1‰ for hydrogenotrophic, methylotrophic, and acetoclastic cultures, respectively.  
405 For each pathway, lower  $\Delta^{13}CH_3D$  values tend to be associated with lower  $\delta D_{H_2O}$  values (Figures  
406 3 and 4). The linear fit and standard error for the slope and intercept was calculated as described  
407 above. Hydrogenotrophic cultures yielded a higher slope ( $0.0020 \pm 0.0010$ ) compared to  
408 methylotrophic ( $0.0019 \pm 0.0014$ ) and acetoclastic ( $0.0016 \pm 0.0012$ ) cultures.

409

### 410 **3.4 D-labeled acetate experiment**

411 In order to track the transfer of D (or H) from the methyl group of acetate to methane, *M.*  
412 *barkeri* was cultured with medium spiked with  $CD_3COO^-$ . As the amount of  $CD_3COO^-$  in an  
413 acetoclastic methanogen culture increased, not only the relative abundance of  $CHD_3$  but also that  
414 of  $CH_2D_2$  and  $CH_3D$  increased at the expense of  $CH_4$  (Figure 4). Cultures incubated with 100%  
415  $CD_3COO^-$  produced methane comprised of a majority of triply deuterated isotopologues (68%  
416  $CHD_3$ ) but also contained  $CH_2D_2$ , (13%) and  $CH_3D$  (5%).

417

## 418 **4. Discussion**

### 419 **4.1 Closed system isotope effect does not explain non-equilibrium $\Delta^{13}CH_3D$**

420 Methanogens in the temperature series experiment (Method 2.1) were incubated until  
421 nearly all the substrate ( $CO_2$ ) had been converted to product ( $CH_4$ ). As a result, the  $\delta^{13}C$  value of  
422 the product methane would have increased with reaction progress, eventually reaching the  $\delta^{13}C$   
423 value of the starting  $CO_2$ . In addition to changes in  $\delta$  values due to closed system effects, it has  
424 been shown that the apparent D/H fractionation factor between methane and water changes with

425 growth phase (Valentine et al., 2004; Kawagucci et al., 2014). This could be due to changes in the  
 426  $\delta D$  value of intracellular water via D/H exchange with  $H_2$  (Burke, 1996) or the contribution from  
 427 metabolic water (Kawagucci et al., 2014). Values of  $\Delta^{13}CH_3D$  do track non-linearly with  $\delta^{13}C$  and  
 428  $\delta D$  upon the mixing of two or more pools of methane (Stolper et al., 2015; Wang et al., 2015;  
 429 Douglas et al., 2016), such that mixing of two methane reservoirs would result in non-equilibrium  
 430  $\Delta^{13}CH_3D$  values even when  $\Delta^{13}CH_3D$  values of the source reservoirs carry equilibrium signals.  
 431 We sought to isolate any experimental effects introduced in the closed system and therefore tested  
 432 if changing the  $^{13}C/^{12}C$  and/or D/H ratios of bulk methane over the course of the reaction may also  
 433 affect the  $\Delta^{13}CH_3D$  value of the end product.

434 The production of four methane isotopologues from two isotopologues of  $CO_2$  can be  
 435 written as:



440 where,  $k$  is the pseudo-first order rate constant for  $^{12}CO_2$  to  $^{12}CH_4$ ,  $^{13}\alpha$  and  $^2\alpha$  are the carbon  
 441 and hydrogen isotope fractionation factors, respectively, and  $R_H$  is the D/H ratio of source hydrogen  
 442 (intracellular water,  $R_H = R_{SMOW}(\delta D_{H_2O} + 1)$ ). For the application of equation (20) and (21), we assume  
 443 that the source for H of  $CH_4$  is intracellular water. The value of  $\delta D$  of intercellular water can be  
 444 different from that of media due to exchange with  $H_2$  or production of metabolic water inside the  
 445 cytoplasm (Burke, 1993; Kawagucci et al., 2014). In addition, the direct transfer of H in  $H_2$  to  $CH_4$   
 446 was suggested (Kawagucci et al., 2014). Since the detailed mechanism of the effect of  $\delta D$ - $H_2$  is  
 447 beyond the scope of this study, the above model includes the effect as the change of the  $^2\alpha$  value  
 448 during the course of the culture. Changing  $R_H$  would produce results identical to changing  $^2\alpha$ .  
 449 Equations (18) to (21) were integrated numerically with three fitting parameters ( $^{13}\alpha$ ,  $^2\alpha$ , and  $\gamma$ ),  
 450 and the results are shown in Figure 2. We used  $\delta^{13}C$  and  $\delta D$  data to fit  $^{13}\alpha$  and  $^2\alpha$ , and the  $\Delta^{13}CH_3D$   
 451 data was used to fit  $\gamma$  value for the derived  $^{13}\alpha$  and  $^2\alpha$  values.

452 The best fit to the experimental data was obtained when  $^{13}\alpha$  was 0.97 and  $^2\alpha$  changed  
 453 linearly from 0.69 to 0.57 from 0 to 25% reaction and remained a constant value of 0.57 afterwards  
 454 (Figure 2 A and B) ( $\delta D_{H_2O}$  of -49.6‰ SMOW, and  $\delta^{13}C_{CO_2}$  of 10.9‰). As  $\delta D_{H_2O}$  does not change  
 455 significantly during the course of the experiment, the change in  $\delta D_{CH_4}$  cannot be explained by the  
 456 closed system effect. The increasing fractionation factor at a later stage indicates  $\delta D_{CH_4}$  is moving  
 457 away from the value expected for equilibrium with water. Previous studies also observed similar  
 458 changes in apparent D/H fractionation factors during early exponential growth phases (Valentine  
 459 et al., 2004; Kawagucci et al., 2014; Okumura et al., 2016). Values of  $\delta D$  of methane produced  
 460 during the early growth phase can be a function of  $\delta D_{H_2}$  as well as  $\delta D_{H_2O}$  (Kawagucci et al., 2014).

461 For those  $^{13}\alpha$  and  $^2\alpha$  values, the best fit for  $\gamma$  was 1.0020 and 1.0032 for the early and late growth  
 462 phases, respectively (Figure 2). The value of  $\gamma$  of higher than unity indicates that the rate of  $^{13}\text{CH}_3\text{D}$   
 463 production is faster than the rate expected from the product of the two fractionation factors ( $^{13}\alpha$   
 464  $^2\alpha$ ).

465 These derived  $\gamma$  values of 1.0020 and 1.0032 translate to closed-system corrected  $\Delta^{13}\text{CH}_3\text{D}$   
 466 values of 2.0 and 3.2‰ respectively (corresponding to 243 and 135°C, for apparent equilibrium  
 467 temperatures). These values are similar to the uncorrected values of 2.1 to 2.5‰, demonstrating  
 468 that the effect of a closed system can only partially explain the non-equilibrium  $\Delta^{13}\text{CH}_3\text{D}$  signals  
 469 of microbial methane. The numerical model above shows that the  $\Delta^{13}\text{CH}_3\text{D}$  value of accumulated  
 470 methane decreases by up to 1.2‰ over the course of reaction given constant  $\gamma$  values, due to the  
 471 effect of mixing of methane formed during early and late exponential growth phases (Figure 2C).  
 472

#### 473 4.2 Origin of H in methane from three methanogenesis pathways

474 The results of the D-spiked series experiments can be used to estimate the origin of C-H  
 475 bonds in methane during three pathways of methanogenesis and associated deuterium isotope  
 476 effects ( $^2\alpha_p$  and  $^2\alpha_s$ ). Assuming methane is formed via mixing of hydrogen atoms both from  $\text{H}_2\text{O}$   
 477 and the methyl group of acetate or methanol, the  $\delta\text{D}$  value of product methane can be written as:

$$478 \delta\text{D}_{\text{CH}_4} + 1 = ^2\alpha_p f (\delta\text{D}_{\text{H}_2\text{O}} + 1) + ^2\alpha_s (1-f) (\delta\text{D}_{\text{CH}_3} + 1) \quad (22)$$

480 where,  $^2\alpha_p$  is the kinetic fractionation factor from  $\text{H}_2\text{O}$  to  $\text{CH}_4$  (primary D-isotope effect), and  $^2\alpha_s$   
 481 is the kinetic isotope fractionation factor from  $\text{CH}_3$  (methyl-H) to  $\text{CH}_4$  (secondary D-isotope effect)  
 482 (Sessions and Hayes, 2005). The value of  $f$  is the fraction of H from  $\text{H}_2\text{O}$ . The canonical value of  
 483  $f$  is 0.25, but this can be higher when scrambling of C-H bonds occurs between C-H and  
 484 intercellular  $\text{H}_2\text{O}$ .  
 485

486 Our experimental results yield the following relationships for  $\text{CO}_2 + \text{H}_2$ , acetate and  
 487 methanol cultures, respectively (Figure 5):

$$488 \delta\text{D}_{\text{CH}_4} + 1 = (0.571 \pm 0.011) (\delta\text{D}_{\text{H}_2\text{O}} + 1) + (0.006 \pm 0.012), \quad (23)$$

$$490 \delta\text{D}_{\text{CH}_4} + 1 = (0.212 \pm 0.004) (\delta\text{D}_{\text{H}_2\text{O}} + 1) + (0.457 \pm 0.005), \text{ and} \quad (24)$$

$$491 \delta\text{D}_{\text{CH}_4} + 1 = (0.269 \pm 0.005) (\delta\text{D}_{\text{H}_2\text{O}} + 1) + (0.361 \pm 0.006). \quad (25)$$

492 For hydrogenotrophic methanogenesis, a small (can be zero within standard error) intercept  
 493 suggests all four hydrogen atoms are derived from water (i.e.,  $f=1$ ), and an  $\alpha_p$  value of 0.571 is  
 494 obtained, which is within the range of previous experiments 0.55 and 0.86 (Valentine et al., 2004;  
 495 Yoshioka et al., 2008; Kawagucci et al., 2014; Okumura et al., 2016).

496 The acetate culture (equation 24) yields:  
497  ${}^2\alpha_p f = 0.211$ , and (26)

498  ${}^2\alpha_s (1-f)(\delta D_{CH_3} + 1) = 0.457$ . (27)

499 The methyl-H of acetate is measured as -123‰ (i.e.,  $\delta D_{CH_3+1} = 0.877$ , Table 3). The two  
500 fractionation factors,  $\alpha_p$  and  $\alpha_s$ , are calculated for a given value of  $f$  in Figure 6. The equations (26  
501 and 27) cannot provide a unique answer, as there are two equations with three unknowns ( $\alpha_p$ ,  $\alpha_s$   
502 and  $f$ ). The likely range of values can be constrained, however, because isotope fractionation  
503 factors are expected to be normal (i.e.,  $\alpha_p < 1$  and  $\alpha_s < 1$ ), and the value of  $f$  is between  $\frac{1}{4}$  and 1.  
504 This yields a range of possible values:  $0.449 < {}^2\alpha_p < 0.844$  and  $0.695 < {}^2\alpha_s < 1.0$ , and  $0.25 < f < 0.48$   
505 (Figure 6). The secondary isotope effect for the formation of methane from methyl-coenzyme M  
506 (the last step of methanogenesis, utilized by all known methanogens Figure 7) is reported to be  
507 0.84 (Scheller et al., 2013). For example, this value for  ${}^2\alpha_s$ , would yield  $f=0.39$  and  ${}^2\alpha_p=0.545$ .  
508 The analysis indicates that the value of  $f$  significantly deviates from the canonical value of 0.25,  
509 which is expected from the reaction stoichiometry. Therefore, this analysis suggests that among  
510 the four hydrogen atoms in methane, up to 1.9 hydrogen atoms ( $=0.48 \times 4$ ) are derived from water,  
511 whereas only one is required from reaction stoichiometry. Since  $\delta D_{CH_3}$  of methanol was not  
512 measured, we cannot carry out the same analysis for methanol experiments. The slightly higher  
513 slope for methanol cultures (0.269) compared to acetate cultures (0.212), however, suggests a  
514 similar or greater contribution of hydrogen atoms from water if we assume the same  ${}^2\alpha_p$  value for  
515 the acetate culture.

516 Previous incubation studies with D-spiked water showed slopes between  $\delta D_{CH_4}$  and  $\delta D_{H_2O}$   
517 of 0.4 for sewage sludge (Schoell, 1980) and 0.48 to 0.61 for paddy soil (Sugimoto and Wada,  
518 1995). These are higher than the value obtained by our experiments, indicating a higher  
519 contribution of water-H to methane in these incubation studies. Since previous experiments were  
520 enriched cultures (not pure cultures), results from previous experiments reflect a mixed  
521 contribution from hydrogenotrophic and acetoclastic methanogenesis. A previous incubation study  
522 of lake and estuary sediments using triply deuterated acetate ( $CD_3COO^-$ ) showed rapid exchange  
523 of the methyl-H of acetate by methanogenic acetate metabolism (de Graaf et al., 1996). Such an  
524 exchange would contribute to the greater slope for the incubation study using natural populations,  
525 compared to pure culture experiments for this study.

526

### 527 4.3 Pathway of D-isotope exchange during acetoclastic methanogenesis

528 Figure 7 illustrates the three pathways of methanogenesis examined in this study. The solid  
529 arrows represent the predominant direction of reaction, and dashed arrows represent the backward  
530 reaction, which is thought to be minor. The two solid arrows are used where the reactions are



531 thought to be reversible. This reversibility was inferred from our results as well as based on  
532 previous studies (e.g., Thauer et al., 1998; Ferry 2010). Hydrogenotrophic methanogenesis  
533 proceeds with a series of two electron reactions, which each adding one H, while the C<sub>1</sub> group is  
534 carried by cofactors: methanofuran (MF), tetrahydrosarcinaopterin (H<sub>4</sub>SPT), and coenzyme M  
535 (CoM). Acetoclastic methanogenesis is a disproportionation reaction where acetate is activated to  
536 acetyl-coA (via acetyl phosphate) and split between a methyl and a carbonyl (CO) moiety, where  
537 the latter is oxidized to CO<sub>2</sub>. The methyl group is transferred to CH<sub>3</sub>-H<sub>4</sub>SPT (methyl-  
538 tetrahydrosarcinaopterin), and then to CH<sub>3</sub>-CoM (methyl-coenzyme M), and finally reduced to  
539 methane. Methylotrophic methanogenesis is also a disproportionation reaction where the oxidation  
540 of one methyl group is coupled with the reduction of three methyl groups. The last step (the  
541 reduction of methyl-CoM) is common to all three pathways, and most (if not all) reactions are  
542 thought to be reversible (Thauer, 2008).

543 Our analysis of the D-spiked water experiments, together with the D-labeled acetate  
544 experiment, suggests that up to 1.9 out of four hydrogen atoms in methane are derived from water-  
545 H. In addition to one hydrogen atom that is added at the last step of methanogenesis (the reduction  
546 of methyl-CoM, Figure 7 and 8), nearly another one of three hydrogen atoms in the methyl group  
547 of acetate could be exchanged with water. Since hydrogen isotope exchange between H<sub>2</sub>O and  
548 CH<sub>4</sub> is sluggish, the exchange probably occurs at the CH<sub>3</sub>-H<sub>4</sub>SPT (methyl-  
549 tetrahydrosarcinaopterin) intermediate via H-abstraction to form CH<sub>2</sub>=H<sub>4</sub>SPT (the exchange can  
550 also occur at the methyl-CoM moiety, Scheller et al., 2013). This part of the pathway is not  
551 required for the acetoclastic metabolism but can occur as a side reaction. A similar D/H exchange  
552 mechanism was suggested to explain the observed scrambling of CD<sub>3</sub>COO<sup>-</sup> during the incubation  
553 of methanogenic sediments (de Graaf et al., 1996). Because up to 0.9 out of three methyl-H atoms  
554 are exchanged, the flux of this side reaction is estimated to be at most 0.3 H (=0.9/3) per uptake of  
555 one acetate (Figure 8-A).

556 When deuterium-labeled acetate (CD<sub>3</sub>COO<sup>-</sup>) was used as a substrate, the major product  
557 (68%) was the isotopologue CD<sub>3</sub>H. However, the isotopologues CD<sub>2</sub>H<sub>2</sub>, CDH<sub>3</sub> and CH<sub>4</sub> were also  
558 formed (Figure 4). Among the four isotopologues, the fraction of CH<sub>4</sub> isotopologues was  
559 disproportionally high. This is because CH<sub>4</sub> and non-deuterated acetate are carried over from the  
560 inoculum. If all CH<sub>4</sub> isotopologues are from the inoculum, the proportion of CD<sub>3</sub>H isotopologues  
561 is 79% (=68/(5+13+68)) over CD<sub>3</sub>H, CH<sub>2</sub>D<sub>2</sub> and CH<sub>3</sub>D isotopologues. This is the maximum  
562 fraction since a small quantity (most likely less than 5%) could be produced from CD<sub>3</sub>COO<sup>-</sup>.  
563 Following the model of D/H exchange at the CH<sub>3</sub>-H<sub>4</sub>SPT step as described above, this means that  
564 less than 100% yield of CD<sub>3</sub>H is explained by the loss of CD<sub>3</sub>-H<sub>4</sub>SPT to CD<sub>2</sub>=H<sub>4</sub>SPT of 0.2 per  
565 one CD<sub>3</sub>COO<sup>-</sup>. The reaction from CD<sub>3</sub>-H<sub>4</sub>SPT to CD<sub>2</sub>=H<sub>4</sub>SPT and from CD<sub>3</sub>-CoM to CD<sub>3</sub>H would  
566 accompany deuterium isotope effects of <sup>2</sup>α<sub>p</sub>, <sup>2</sup>α<sub>s</sub><sup>2</sup> and <sup>2</sup>α<sub>s</sub><sup>3</sup>, respectively. Thus, the corresponding

567 flux for CH<sub>3</sub>COO<sup>-</sup> can be higher by the ratio of the two isotope effects ( $^2\alpha_s/2\alpha_p$ ), which could be as  
 568 high as ~2 based on *in vitro* study of a similar reaction ( $\approx 0.84/0.41$ ) (Scheller et al., 2013). Note  
 569 that the medium water (and thus intercellular water) contain very little deuterium such that there  
 570 is practically no back flux from CD<sub>2</sub>=H<sub>4</sub>SPT to CD<sub>3</sub>-H<sub>4</sub>SPT (Figure 8-B). Thus, H/D scrambling  
 571 at the CH<sub>3</sub>-moiety can be explained by the reverse flux of 0.2 to 0.3 based on the estimate from  
 572 the D-spiked medium (0 to 0.3) and CD<sub>3</sub>COO<sup>-</sup> experiments (0.2 to ~0.4) (Figure 8).

573

#### 574 4.4 Origin of <sup>13</sup>C-H signals in methane from acetoclastic methanogenesis

575 Measured  $\Delta^{13}\text{CH}_3\text{D}$  values are -1.7 to -3.1, -1.1 to -3.8, -4.2 to -5.2‰ for *Methanosarcina*  
 576 cultures grown on acetate, H<sub>2</sub>+CO<sub>2</sub>, and methanol, respectively, and are weakly correlated with  
 577  $\delta\text{D}_{\text{H}_2\text{O}}$  of the medium (Figure 3B). The linear fit and standard error (95% confidence interval) was  
 578 calculated following York et al. (2004). The slope ( $\Delta^{13}\text{CH}_3\text{D}/\delta\text{D}$ ) for acetate is  $0.0016\pm 0.0012$ ,  
 579 hydrogen is  $0.0020\pm 0.0010$  and methanol is  $0.0019\pm 0.0014$ . While this is a small sample size,  
 580 these are statistically not zero. These anti-clumped  $\Delta^{13}\text{CH}_3\text{D}$  values can originate from 1) mixing  
 581 of two or more pools of methane or its precursor with different  $\delta^{13}\text{C}$  and  $\delta\text{D}$  values, 2) transfer of  
 582 the methyl-group of acetate and methanol with pre-existing anti-clumped signals (for acetate and  
 583 methanol cultures), and/or 3) intrinsic kinetic isotope effects associated with enzymatic reactions  
 584 common to three pathways, such as the reduction of methyl-CoM (Figure 7).

585 It has been well known that mixing is non-linear in the clumped isotope system such that  
 586 mixing of two pools of methane yields a  $\Delta^{13}\text{CH}_3\text{D}$  value that is not between the two  $\Delta^{13}\text{CH}_3\text{D}$   
 587 values of original two reservoirs (Eiler and Schauble, 2004; Affek and Eiler, 2006; Affek et al.,  
 588 2007; Defliese and Lohmann, 2015). When two pools of methane (A and B) are mixed, the  
 589  $\Delta^{13}\text{CH}_3\text{D}$  value of the mixture ( $\Delta_{\text{mix}}$ ) can be approximated as:

$$590 \quad \Delta_{\text{mix}} \approx (1-f) \Delta_A + f\Delta_B + f(1-f) (\delta^{13}\text{C}_A - \delta^{13}\text{C}_B)(\delta\text{D}_A - \delta\text{D}_B) \quad (28)$$

591 where  $f$  is the mixing ratio of pool B (Wang et al., 2015). The first and second terms show linear  
 592 mixing between two  $\Delta^{13}\text{CH}_3\text{D}$  values ( $\Delta_A$  and  $\Delta_B$ ), whereas the third term produces a curvature  
 593 following a quadratic function to  $f$ . This bias becomes the largest when the two pools of methane  
 594 are mixed at a 1:1 ratio (i.e.,  $f=0.5$ ), and proportional to the product of the difference of  $\delta$  values  
 595 between the two pools of methane. Mixing of a <sup>13</sup>C- and D-enriched pool with a <sup>13</sup>C- and D-  
 596 depleted pool would produce a positive (low temperature) bias, whereas diagonal mixing (e.g.,  
 597 mixing between a <sup>13</sup>C-enriched and a D-depleted pool with a <sup>13</sup>C-depleted and a D-enriched pool)  
 598 would produce a negative (high temperature) bias in  $\Delta^{13}\text{CH}_3\text{D}$ . By extension, the mixing effect for  
 599 doubly-deuterated clumped methane (<sup>12</sup>CH<sub>2</sub>D<sub>2</sub>) will always produce a positive (low temperature)  
 600 bias (Young et al., 2016). Equation (28) can be used to model reaction branching (i.e., producing

601 two products) or reversible reactions (where *mix* is the source and A and B are forward and  
602 backward reactions).

603 A fully quantitative model for the isotope systematics for the acetoclastic pathway in Figure  
604 8 is the scope of future study, but equation (28) can be used to test the magnitude of any  $\Delta^{13}\text{CH}_3\text{D}$   
605 bias due to mixing. For example, mixing of two reservoirs (or fluxes) in equal portions ( $f=0.5$ )  
606 with  $\delta^{13}\text{C}$  and  $\delta\text{D}$  differing by 40‰ and -300‰, respectively, can produce a non-linear bias of -  
607 3‰ ( $0.5(1-0.5)\times 0.04\times(-0.3)$ ). These magnitudes of kinetic  $^{13}\text{C}$ - or D-isotope effects are possible.  
608 Therefore, an anti-clumped  $\Delta^{13}\text{CH}_3\text{D}$  value of -3‰ can be produced entirely by mixing. For the  
609 acetoclastic cultures in these experiments, however, the value of  $\Delta^{13}\text{CH}_3\text{D}$  was not sensitive to the  
610  $\delta\text{D}$  value of the medium, which changed from -35 to 360‰ (Figure 3B), indicating that the mixing  
611 (between C-H bonds with water-derived H and methyl-derived H) is unlikely to be the major  
612 source of anti-clumped signals observed.

613 Calculations from the D-spiked cultures grown on acetate reported in this study showed  
614 that approximately up to two out of four hydrogen atoms in  $\text{CH}_4$  are derived from water, while  
615 only one is required from stoichiometry. On average, the methyl group of acetate contributes 2 to  
616 3 hydrogen atoms to one methane molecule, presumably also carrying its original  $^{13}\text{C}$ -D signal.  
617 Although the degree of  $^{13}\text{C}$ -D clumping of the methyl group of acetic acid (or methanol) cannot  
618 be measured by our current instrumentation, we expect its  $^{13}\text{C}$ -D clumping signal is not much  
619 different from that of  $\text{CH}_4$ . This is because most industrial acetate (the likely source of sodium  
620 acetate used in this study) is produced from the high temperature (150 to 200°C) catalytic reaction  
621 of methanol and carbon monoxide (e.g., Eby and Singleton, 1983). Industrial methanol, in turn, is  
622 produced from carbon monoxide, carbon dioxide, and hydrogen at high temperature (typically 200  
623 to 300°C) (Cheng and Kung, 1994). Based on theoretical calculations from modeled vibrational  
624 frequencies, Wang et al. (2015) reported that the *equilibrium*  $^{13}\text{C}$ -D clumping of simple carbon  
625 compounds have a relatively narrow range of clumped isotope effects from +5.9 to +6.2‰ at 25°C  
626 for the molecules studied (methane, methanol, formaldehyde, formic acid, methanethiol, acetic  
627 acid). This indicates that the  $^{13}\text{C}$ -D clumped isotope effect is not sensitive to detailed bonding  
628 environments. This is reasonable considering that the  $^{13}\text{C}$ -D clumped isotope effect largely  
629 originates from a zero point energy shift ( $\Delta\text{ZPE}$ ) associated with the C-H stretching vibration  
630 frequency at around  $3000\text{ cm}^{-1}$ ; bending vibration is much lower in energy ( $1350\text{ cm}^{-1}$ ) and the  
631  $\Delta\text{ZPE}$  is relatively small (Whitehill et al., 2017). Therefore, we estimate that the acetate or  
632 methanol may thus carry a  $\Delta^{13}\text{CH}_3\text{D}$  signal of 1.6 to 3.1‰ assuming the near-equilibrium reaction  
633 between 150 and 300°C. If we assume that acetate was produced at the lower end of this  
634 temperature range (150°C) and methanol was produced at the higher end (300°C), this would  
635 correspond to equilibrium values of 3.1‰ for acetate vs. 1.6‰ for methanol. It is possible that this  
636 discrepancy could explain part, but not all of the difference between the  $\Delta^{13}\text{CH}_3\text{D}$  values of the

637 acetoclastic and methylotrophic cultures. Nonetheless, experiments with various sources of acetic  
638 acid, natural and/or synthetic, will be needed to constrain the degree of non-equilibrium signals in  
639 <sup>13</sup>C-D bonds of acetic acids.

640 Another possibility is that the observed anti-clumped signal originates during the addition  
641 of the last hydrogen atom of methane. By applying equation (28), the  $\Delta^{13}\text{CH}_3\text{D}$  value of methane  
642 ( $\Delta_{\text{CH}_4}$ ) is expected to carry  $\frac{3}{4}$  signal from methyl precursor ( $\text{CH}_3\text{-CoM}$ ,  $\Delta_{\text{CH}_3}$ ), and  $\frac{1}{4}$  from the last  
643 H added:

$$644 \quad \Delta_{\text{CH}_4} = \frac{3}{4}(\ln \gamma_S + \Delta_{\text{CH}_3}) + \frac{1}{4} \ln \gamma_P \quad (29)$$

645 where  $\gamma_S$  and  $\gamma_P$  are kinetic clumped isotope effects for the secondary and primary D-addition  
646 (equations 13 and 14). Here, the non-linearity bias in equation (28) does not apply, because <sup>13</sup>C  
647 isotope effects for secondary and primary processes will be nearly identical (i.e.,  $\delta^{13}\text{C}_A - \delta^{13}\text{C}_B \approx$   
648 0).

649 What are the values of  $\gamma_S$  and  $\gamma_P$ ? Applying transition state theory, Whitehill et al. (2017)  
650 presented a detailed analysis of the kinetic clumped effect for the gas phase oxidation of methane  
651 by the OH radical. We use their framework to make an approximate inference for this study.  
652 According to Whitehill et al. (2017), clumped isotope effects can be explained by the difference  
653 of zero-point energy shifts ( $\Delta\text{ZPE}$ ) between <sup>13</sup>CH/<sup>12</sup>CH and <sup>13</sup>CD/<sup>12</sup>CD. For methane, the  $\Delta\text{ZPE}$   
654 for <sup>13</sup>CH<sub>4</sub> vs. <sup>12</sup>CH<sub>4</sub> is 29.8 cm<sup>-1</sup> and that for <sup>13</sup>CH<sub>3</sub>D vs. <sup>12</sup>CH<sub>3</sub>D is 31.9 cm<sup>-1</sup>. The difference  
655 between the two,  $\Delta\Delta\text{ZPE}$  of 2.1 cm<sup>-1</sup>, is the origin of clumped effect (Whitehill et al., 2017).  
656 Transition state for the last step of methanogenesis involves a methyl radical intermediate with  
657 trigonal planar geometry (Scheller et al., 2013; Wongnate et al., 2017). We estimated the  $\Delta\Delta\text{ZPE}$   
658 of the methyl radical of 2.1cm<sup>-1</sup> using unrestricted MP2 with basis set aug-cc-pVQZ basis set.  
659  $\Delta\Delta\text{ZPE}$  for reactant ( $\text{CH}_3\text{-S-CoM}$ , approximated by methylthiol, Wang et al., 2015) and transition  
660 state (approximated by the methyl radical) suggests that there is little kinetic clumped effect for  
661 the secondary reaction such that  $\gamma_S$  is expected to be close to unity. Whitehill et al. (2017) also  
662 showed that the imaginary frequency and tunneling terms do not produce clumped effects (within  
663 transition state theory), although the Wigner tunneling correction used in the study is highly  
664 approximated.

665 For the primary reaction, the C-D bond becomes loose at a transition state, contributing a  
666 smaller  $\Delta\Delta\text{ZPE}$  between 0 and 2.1 cm<sup>-1</sup> (it is 0.5cm<sup>-1</sup> for  $\text{CH}_3\text{-D-OH}$  transition state, Whitehill et  
667 al., 2017). For the bond forming reaction, <sup>13</sup>C-D is slightly preferred but not as much as equilibrium  
668 ( $\gamma_P < 1.006$ ). However, the anti-clumped effect ( $\gamma_P < 1$ ) is unlikely because it requires the  $\Delta\Delta\text{ZPE}$  to  
669 have the opposite sign (i.e., smaller  $\Delta\text{ZPE}$  for <sup>13</sup>CD/<sup>12</sup>CD than <sup>13</sup>CH/<sup>12</sup>CH) at the transition state.  
670 Thus, based on this transition state model,  $\gamma_P$  could take any value between 1.000 and 1.006.

671 For example, if  $\ln(\gamma_S)$  and  $\Delta_{CH_3}$  are 0.0 and 1.6 ‰, respectively, the value of  $\ln(\gamma_P)$  of -13  
672 ‰ is required to explain the observed  $\Delta^{13}CH_3D$  value in the acetate culture of  $\sim -3\%$  (equation  
673 (29). The magnitude of this kinetic anti-clumped effect has been implicated by Wang et al. (2015)  
674 as well as Stolper et al. (2015) to explain the observed values for natural samples. Low  $\Delta^{13}CH_3D$   
675 values for methanol cultures (-4.2 to -4.9 ‰) suggest that the methyl group of methanol may carry  
676 lower  $\Delta^{13}CH_3D$  values.

677 Another possibility is that the anti-clumping effect for methanol cultures is related to the  
678 faster growth rate of methylotrophic compared to acetoclastic and hydrogenotrophic cultures.  
679 Cultures were incubated until they had produced enough methane for our analyses (>5mL, STP)  
680 which was reached over the course of different incubation times as noted in Table 2. Although  
681 exact experiment durations and methane concentrations were not controlled for in all experimental  
682 setups, incubation times and volume of methane produced are reported for Set 2 cultures (Table  
683 2). Despite some measurement uncertainty, it appears methanol cultures generated methane more  
684 quickly. Quantifying the effect of growth rate warrants investigation in future work.

685 A series of studies suggest the potential importance of quantum mechanical tunneling in  
686 some H transfer reactions, in particular, during the oxidation of alcohol by dehydrogenase (Cha et  
687 al., 1989; Klinman and Kohen, 2013). Experimental evidence that supports tunneling includes the  
688 observed mass-independent fractionation among H/D/T, departure from the rule of geometric  
689 mean for multiply-deuterated (clumped) substrates, rate enhancement at low temperatures  
690 (Klinman and Kohen, 2013; Srinivasan et al., 1985; Amin et al., 1988; Huskey, 2007), and the anti-  
691 clumped  $CH_2D_2$  abundance (Young et al., 2017). The large anti-clumped effect may be due to the  
692 tunneling effect, which is only approximated in the above transition state model. The test for this  
693 hypothesis would include *in vitro* enzyme assay experiments similar to Scheller et al. (2013), or  
694 high-level quantum mechanical modeling with accurate geometry and potential energy surface at  
695 the key transition state (e.g., Chen et al., 2012; Klinman and Kohen, 2013; Wongnate et al., 2016).  
696

#### 697 **4.5 Non-equilibrium vs. equilibrium $\Delta^{13}CH_3D$ signals of methane in the** 698 **environment**

699 This study corroborates previous studies (Stolper et al., 2015; Wang et al., 2015; Douglas et  
700 al., 2016; Wang et al., 2016; Young et al., 2016; Young et al., 2017) and demonstrates that non-  
701 equilibrium (i.e., kinetic) clumped methane isotopologue signals are common for methane  
702 produced by microbial methanogenesis in laboratory cultures (Figure 1, Figure 9).

703 Our results are consistent with kinetic signals in methane sampled from freshwater  
704 environments (e.g. swamps and lakes), where acetoclastic methanogenesis, as opposed to  
705 hydrogenotrophic methanogenesis, is thought to be the dominant source of methane (e.g., Conrad

706 2005; Ferry 2010). Nonetheless, our results demonstrate that low  $\Delta^{13}\text{CH}_3\text{D}$  values are consistently  
707 out of equilibrium in batch culture and in environmental samples, but that this is not necessarily  
708 dependent on the metabolic pathway (hydrogenotrophic, methyotrophic, or acetoclastic). Thus,  
709  $^{13}\text{CH}_3\text{D}$  isotopologue compositions cannot be used alone to resolve which methanogenic pathway  
710 is dominant in the environment.

711 Although dominantly microbial in origin, methane in marine environments (e.g. pore water  
712 and hydrate) tends to carry equilibrium or near-equilibrium  $^{13}\text{CH}_3\text{D}$  abundances (Stolper et al.,  
713 2015; Wang et al., 2015). This is also corroborated by  $\delta\text{D}_{\text{CH}_4}$  values that are relatively constant at  
714  $-180\pm 10\text{‰}$  (e.g., Whiticar, 1999; Okumura et al., 2016), which is close to the expected value for  
715 methane in equilibrium with seawater water (with  $\delta\text{D}\approx 0\text{‰}$ ) (Horibe and Craig, 1995). It was  
716 suggested that “slow” methanogenesis under small thermodynamic drive (low environmental  $\text{H}_2$   
717 concentration) would produce near-equilibrium clumped isotopologue signals (Wang et al., 2015;  
718 Stolper et al., 2015). Oxidative cycles of methane also modify clumped isotopologue signals of  
719 environmental methane. The aerobic oxidation of methane by *Methylococcus*, for example, is  
720 characterized by kinetic clumped isotope effects ( $\gamma\approx 1$ ), and residual methane can exhibit strong  
721 non-equilibrium signals (Wang et al., 2016). In contrast, anaerobic oxidation of methane (AOM),  
722 thought to be the reverse process of methanogenesis (Hallam et al., 2004; Scheller et al., 2013;  
723 Wang et al., 2014), has been shown to operate with significant back-flux (Holler et al., 2011) and  
724 produce near-equilibrium carbon isotope fractionation (e.g., Yoshinaga et al., 2014). Methane  
725 cycling involving AOM and methanogenesis likely promotes near-equilibrium isotopologue  
726 distributions, in particular, in marine environments, and may contribute to near-equilibrium  $\delta\text{D}$  as  
727 well as  $\Delta^{13}\text{CH}_3\text{D}$  signals. Conversely, bulk  $\delta\text{D}$  and  $\Delta^{13}\text{CH}_3\text{D}$  values of methane may primarily  
728 reflect kinetic versus equilibrium signals and the rate of methanogenesis more so than metabolic  
729 pathways.

## 730 5. Conclusions

731 This study reports the  $\Delta^{13}\text{CH}_3\text{D}$  systematics of microbial methane produced by pure  
732 cultures of methanogens. Our results show that the  $\Delta^{13}\text{CH}_3\text{D}$  signals are not directly pathway  
733 dependent, as cultures of *M. barkeri* and *M. mazei* grown on acetate, methanol, and  $\text{H}_2+\text{CO}_2$  all  
734 yield methane that is depleted in  $^{13}\text{CH}_3\text{D}$ , which seems to be characteristic of microbial  
735 methanogenesis in near-surface environments (lakes, swamps and ruminants). For mesophilic  
736 methanogens, the lowest  $\Delta^{13}\text{CH}_3\text{D}$  values were produced for methane from cultures grown on  
737 methanol. Methanol cultures grew faster (incubation time of  $\sim 3$  days) than those grown on acetate  
738 or  $\text{H}_2+\text{CO}_2$  (incubation time of  $\sim 30$  days). Thus, the  $\Delta^{13}\text{CH}_3\text{D}$  values of methane may be related  
739 to the rate at which methane is produced rather than to the substrate used. Mesophilic methanogens

740 (*M. barkeri*) produced anti-clumped  $\Delta^{13}\text{CH}_3\text{D}$  values ( $<0$ ), while thermophilic and  
741 hyperthermophilic (*Methanothermococcus* and *Methanocaldococcus*) methanogens produced  
742 less kinetic signals.

743         Experiments with deuterated water or acetate aid in determining the source of hydrogen  
744 atoms in methane. The deuterated water experiments confirm that the four hydrogen atoms that  
745 form methane in hydrogenotrophic methanogenesis are derived from water. For the acetoclastic  
746 culture, 1.6 to 1.9 H atoms are derived from water, whereas only one is required by stoichiometry,  
747 suggesting some reversibility and isotope exchange at the methyl precursor. The deuterium spiked  
748 experiments also demonstrate that the observed non-equilibrium signals cannot be explained by  
749 the mixing of two pools of C-H bonds (e.g., from methyl group of acetate and one C-H bond  
750 formed during acetoclastic methanogenesis). The production of low  $\Delta^{13}\text{CH}_3\text{D}$  values independent  
751 of the methanogenic pathway, suggests, although not exclusively, that the most of the kinetic signal  
752 is produced during the enzymatic reactions common in the three methanogenic pathways, such as  
753 the reduction of methyl-coenzyme M.  
754

755 **Acknowledgements:**

756 We thank H. Taubner, J. Wendt, X. Prieto, S. Burns, A. Colman, and G. Olack for technical  
757 assistance and  $\delta^{13}\text{C}$  and  $\delta\text{D}$  measurements. We also thank Alex Sessions, Shinsuke Kawagucci,  
758 one anonymous reviewer, and the associate editor, Ed Hornibrook, for constructive feedback.  
759 Grants from the National Science Foundation (EAR-1250394 to S.O.), N. Braunsdorf and D. Smit  
760 of Shell PTI/EG (to S.O.), the Deep Carbon Observatory (to S.O., M.K., K.-U.H., D.S.G.), the  
761 Gottfried Wilhelm Leibniz Program of the Deutsche Forschungsgemeinschaft (HI 616-14-1 to K.-  
762 U.H.), and the Heisenberg Program (KO3651-3-1 to M.K.) of the Deutsche  
763 Forschungsgemeinschaft supported this study. D.S.G. was also supported by a National Science  
764 Foundation Graduate Research Fellowship, the Neil and Anna Rasmussen Foundation Fund, the  
765 Grayce B. Kerr Fellowship, and a Shell-MIT Energy Initiative Graduate Fellowship. D.T.W. was  
766 supported by a National Defense Science and Engineering Graduate Fellowship. L.C.S. was  
767 supported by a NASA Earth and Space Science Fellowship (grant NNX11AP78H).

768

769

770

771



772 **References**

- 773 Affek H. P. and Eiler J. M. (2006) Abundance of mass  $^{47}\text{CO}_2$  in urban air, car exhaust, and human  
774 breath. *Geochim. Cosmochim. Acta* **70**, 1–12.
- 775 Affek H. P., Xu X. and Eiler J. M. (2007) Seasonal and diurnal variations of  $^{13}\text{C}^{18}\text{O}^{16}\text{O}$  in air:  
776 Initial observations from Pasadena, CA. *Geochim. Cosmochim. Acta* **71**, 5033–5043.
- 777 Alperin M. J. and Blair N. E. (1992) Factors that control the stable carbon isotopic composition of  
778 methane produced in an anoxic marine sediment. *Global Biogeochem. Cycles* **6**, 271–291.
- 779 Amin, M., Price, R.C. and Saunders, W.H. (1988) Isotope effects on isotope effects. Failure of the  
780 rule of the geometric mean as evidence for tunneling. *Journal of the American Chemical*  
781 *Society* **110**, 4085–4086.
- 782 Balch W. E., Fox G. E., Magrum L. J., Woese C. R. and Wolfe R. S. (1979) Methanogens:  
783 Reevaluation of a unique biological group. *Microbiol. Rev.* **43**, 260–296.
- 784 Bapteste E., Brochier C. and Boucher Y. (2005) Higher-level classification of the Archaea:  
785 evolution of methanogenesis and methanogens. *Archaea* **1**, 353–363.
- 786 Bigeleisen J. and Mayer (1947) Calculation of equilibrium constants for isotopic exchange  
787 reactions. *J. Chem. Phys.* **15**, 261–267.
- 788 Blair N. E. and Carter W. D. (1992) The carbon isotope biogeochemistry of acetate from a  
789 methanogenic marine sediment. *Geochim. Cosmochim. Acta* **56**, 1247–1258.
- 790 Blaser M. and Conrad R. (2016) Stable carbon isotope fractionation as tracer of carbon cycling in  
791 anoxic soil ecosystems. *Curr. Opin. Biotechnol.* **41**, 122–129.
- 792 Burke R. A. (1993) Possible influence of hydrogen concentration on microbial methane stable  
793 hydrogen isotopic composition. *Chemosphere* **26**, 55–67.
- 794 Cha Y., Murray, C. and Klinman, J. (1989) Hydrogen tunneling in enzyme reactions. *Science* **243**,  
795 1325–1330.
- 796 Chen S. L., Blomberg M. R. A. and Siegbahn P. E. M. (2012) How is methane formed and oxidized  
797 reversibly when catalyzed by Ni-containing methyl-coenzyme M reductase? *Chem. - A*  
798 *Eur. J.* **18**, 6309–6315.
- 799 Cheng, W.-H. and Kung, H.H. (1994) Methanol production and use. Marcel Dekker, New York.
- 800 Conrad R. (2005) Quantification of methanogenic pathways using stable carbon isotopic  
801 signatures: a review and a proposal. *Org. Geochem.* **36**, 739–752.
- 802 Conrad R., Claus P. and Casper P. (2009) Characterization of stable isotope fractionation during  
803 methane production in the sediment of a eutrophic lake, Lake Dagow, Germany. **54**, 457–  
804 471.
- 805 Coplen T. B. (2011) Guidelines and recommended terms for expression of stable-isotope-ratio and  
806 gas-ratio measurement results. *Rapid Commun. Mass Spectrom.* **25**, 2538–2560.
- 807 Daniels L., Fulton, G., Spencer, R. and Orme-Johnson, W. (1980) Origin of hydrogen in methane  
808 produced by *Methanobacterium thermoautotrophicum*. *J. Bacteriol* **141**, 694–698.
- 809 Defliese W. F. and Lohmann K. C. (2015) Non-linear mixing effects on mass-47  $\text{CO}_2$  clumped  
810 isotope thermometry: Patterns and implications. *Rapid Commun. Mass Spectrom.* **29**, 901–  
811 9.
- 812 Douglas P. M. J., Stolper D. A., Smith D. A., Walter Anthony K. M., Paull C. K., Dallimore S.,  
813 Wik M., Crill P. M., Winterdahl M., Eiler J. M. and Sessions A. L. (2016) Diverse origins  
814 of Arctic and Subarctic methane point source emissions identified with multiply-  
815 substituted isotopologues. *Geochim. Cosmochim. Acta* **188**, 163–188.
- 816 Douglas P. M. J., Stolper D. A., Eiler J. M., Sessions A. L., Lawson M., Shuai Y., Bishop A.,  
817 Podlaha O. G., Ferreira A. A., Santos Neto E. V., Niemann M., Steen A. S., Huang L.,

818 Chimia L., Valentine D. L., Fiebig J., Luhmann A. J., Seyfried W. E., Etiope G., Schoell  
819 M., Inskip W. P., Moran J. J. and Kitchen N. (2017) Methane clumped isotopes: Progress  
820 and potential for a new isotopic tracer. *Org. Geochem.* **113**, 262–282.

821 Eby R. T. and Singleton T. C. (1983) Chapter 10: Methanol Carbonylation to Acetic Acid. In  
822 *Applied Industrial Catalysis* (ed. B. E. Leach). Academic Press. p. 275.

823 Eiler J. M. (2007) “Clumped-isotope” geochemistry—The study of naturally-occurring, multiply-  
824 substituted isotopologues. *Earth Planet. Sci. Lett.* **262**, 309–327.

825 Eiler J. M. (2013) The isotopic anatomies of molecules and minerals. *Annu. Rev. Earth Planet.*  
826 *Sci.* **41**, 411–441.

827 Eiler J. M. and Schauble E. (2004)  $^{18}\text{O}^{13}\text{C}^{16}\text{O}$  in Earth’s atmosphere. *Geochim. Cosmochim. Acta*  
828 **68**, 4767–4777.

829 Ermler U., Grabarse W., Shima S., Goubeaud M. and Thauer R. K. (1997) Crystal structure of  
830 methyl coenzyme M reductase: The key enzyme of biological methane formation. *Science*  
831 **278**, 1457–1462.

832 Farquhar J., Johnston D. T. and Wing B. A. (2007) Implications of conservation of mass effects  
833 on mass-dependent isotope fractionations: Influence of network structure on sulfur isotope  
834 phase space of dissimilatory sulfate reduction. *Geochim. Cosmochim. Acta* **71**, 5862–5875.

835 Ferry J. G. (2010) How to make a living by exhaling methane. *Annu. Rev. Microbiol.* **64**, 453–473.

836 Grabarse W., Mahlert F., Shima S., Thauer R. K. and Ermler U. (2000) Comparison of three  
837 methyl-coenzyme M reductases from phylogenetically distant organisms: Unusual amino  
838 acid modification, conservation and adaptation. *J. Mol. Biol.* **303**, 329–344.

839 Hallam S. J., Putnam N., Preston C. M., Detter J. C., Rokhsar D., Richardson P. M. and DeLong  
840 E. F. (2004) Reverse methanogenesis: Testing the hypothesis with environmental  
841 genomics. *Science* **305**, 1457–1462.

842 Hartmann G. C., Santamaria E., Fernandez V. M., and Thauer R. K. (1996) Studies on the catalytic  
843 mechanism of  $\text{H}_2$ -forming methylenetetrahydromethanopterin dehydrogenase: Paraortho  
844  $\text{H}_2$  conversion rates in  $\text{H}_2\text{O}$  and  $\text{D}_2\text{O}$ . *J. Biol. Inorg. Chem.* **1**(5), 446–450.

845 Heuer V., Elvert M., Tille S., Krummen M., Mollar X. P., Hmelo L. R. and Hinrichs K.-U. (2006)  
846 Online  $\delta^{13}\text{C}$  analysis of volatile fatty acids in sediment/porewater systems by liquid  
847 chromatography-isotope ratio mass spectrometry. *Limnol. Oceanogr. Methods*, 346–357.

848 Holler T., Wegener G., Niemann H., Ferdelman T. G., Boetius A., Kristiansen T. Z., Molina H.,  
849 Pandey A., Werner J. K., Juluri K. R., Xu Y., Glenn D., Parang K. and Snyder S. H. (2012)  
850 Correction for Holler et al., Carbon and sulfur back flux during anaerobic microbial  
851 oxidation of methane and coupled sulfate reduction. *Proc. Natl. Acad. Sci.* **109**, 21170–  
852 21170.

853 Huskey, W.P. (2007) Multiple-Isotope Probes of Hydrogen Tunneling, Hydrogen-Transfer  
854 Reactions. Wiley-VCH Verlag GmbH & Co. KGaA, pp. 1285-1309.

855 Huber H., Thomm M., König H., Thies G., and Stetter K.O. (1982) *Methanococcus*  
856 *thermolithotrophicus*, a novel thermophilic lithotrophic methanogen. *Arch. Microbiol.* **132**,  
857 47-50.

858 Hut G. (1987) Consultants’ group meeting on stable isotope reference samples for geochemical  
859 and hydrological investigations. *Rep. to Dir. Gen.*, 16–18.

860 Inagaki F., Hinrichs, K.-U., Kubo Y., Bowles M. W., Heuer V. B., Ijiri A., Imachi H., Ito M.,  
861 Kaneko M., Lever M. A., Morita S., Morono Y., Tanikawa W., Bihan M., Bowden S. A.,  
862 Elvert M., Glombitza C., Gross D., Harrington G. J., Hori T., Li K., Limmer D., Murayama  
863 M., Ohkouchi N., Ono S., Purkey M., Sanada Y., Sauvage J., Snyder G., Takano Y.,

864 Tasumi E., Terada T., Tomaru H., Wang D. T. and Yamada Y. (2015) Exploring deep  
865 microbial life in coal-bearing sediment down to ~2.5 km below the ocean floor. *Science*  
866 **349**, 420–424.

867 Jones W., Leigh J., Mayer F., Woese C., Wolfe R. (1983) *Methanococcus jannaschii* sp. nov., an  
868 extremely thermophilic methanogen from a submarine hydrothermal vent. *Arch.*  
869 *Microbiol.* **136**, 254–261.

870 Kawagucci S., Kobayashi M., Hattori S., Yamada K., Ueno Y., Takai K. and Yoshida N. (2014)  
871 Hydrogen isotope systematics among H<sub>2</sub>–H<sub>2</sub>O–CH<sub>4</sub> during the growth of the  
872 hydrogenotrophic methanogen *Methanothermobacter thermautotrophicus* strain ΔH.  
873 *Geochim. Cosmochim. Acta* **142**, 601–614.

874 Klapp S. a., Bohrmann G., Kuhs W. F., Mangir Murshed M., Pape T., Klein H., Techmer K. S.,  
875 Heeschen K. U. and Abegg F. (2010) Microstructures of structure I and II gas hydrates  
876 from the Gulf of Mexico. *Mar. Pet. Geol.* **27**, 116–125.

877 Klein A. R. and Thauer R. K. (1995) *Re*-face specificity at C14a of  
878 methylenetetrahydromethanopterin and *Si*-face specificity at C5 of coenzyme F<sub>420</sub> for  
879 coenzyme F<sub>420</sub>-dependent methylenetetrahydromethanopterin dehydrogenase from  
880 methanogenic archaea. *Eur. J. Biochem.* **227** (1–2), 169–174. K

881 Klein A. R., Fernandez V. M., and Thauer R. K. (1995a) H<sub>2</sub>-forming N<sup>5</sup>,N<sup>10</sup>-  
882 methylenetetrahydromethanopterin dehydrogenase: Mechanism of H<sub>2</sub> formation analyzed  
883 using hydrogen isotopes. *FEBS Lett.* **368** (2), 203–206.

884 Klein A. R., Hartmann G. C., and Thauer R. K. (1995b) Hydrogen isotope effects in the reactions  
885 catalyzed by H<sub>2</sub>-forming N<sup>5</sup>, N<sup>10</sup>-methylenetetrahydromethanopterin dehydrogenase from  
886 methanogenic archaea. *Eur. J. Biochem.* **233** (1), 372–376.

887 Klinman J. P. and Kohen A. (2013) Hydrogen tunneling links protein dynamics to enzyme  
888 catalysis. *Annu. Rev. Biochem.* **82**, 471–496.

889 Kung, Harold H & Cheng, Wu-Hsun, 1954- (1994). *Methanol production and use*. M. Dekker,  
890 New York

891 Kvenvolden K. A. (1993) Gas hydrates-geological perspective and global change. *Rev. Geophys.*  
892 **31**, 173–187.

893 Liu Q. and Liu Y. (2016) Clumped-isotope signatures at equilibrium of CH<sub>4</sub>, NH<sub>3</sub>, H<sub>2</sub>O, H<sub>2</sub>S and  
894 SO<sub>2</sub>. *Geochim. Cosmochim. Acta* **175**, 252–270.

895 Lopes J. C., de Matos L. F., Harper M. T., Giallongo F., Oh J., Gruen D., Ono S., Kindermann M.,  
896 Duval S. and Hristov A. N. (2016) Effect of 3-nitrooxypropanol on methane and hydrogen  
897 emissions, methane isotopic signature, and ruminal fermentation in dairy cows. *J. Dairy*  
898 *Sci.* **99**, 5335–5344.

899 Ma Q., Wu S. and Tang Y. (2008) Formation and abundance of doubly-substituted methane  
900 isotopologues (<sup>13</sup>CH<sub>3</sub>D) in natural gas systems. *Geochim. Cosmochim. Acta* **72**, 5446–  
901 5456.

902 McCalley C. K., Woodcroft B. J., Hodgkins S. B., Wehr R. a., Kim E.-H., Mondav R., Crill P. M.,  
903 Chanton J. P., Rich V. I., Tyson G. W. and Saleska S. R. (2014) Methane dynamics  
904 regulated by microbial community response to permafrost thaw. *Nature* **514**, 478–481.

905 Okumura T., Kawagucci S., Saito Y., Matsui Y., Takai K. and Imachi H. (2016) Hydrogen and  
906 carbon isotope systematics in hydrogenotrophic methanogenesis under H<sub>2</sub>-limited and H<sub>2</sub>-  
907 enriched conditions: implications for the origin of methane and its isotopic diagnosis. *Prog.*  
908 *Earth Planet. Sci.* **3**, 14.

909 Ono S., Wang D. T., Gruen D. S., Sherwood Lollar B., Zahniser M. S., McManus B. J. and Nelson  
910 D. D. (2014) Measurement of a doubly substituted methane isotopologue,  $^{13}\text{CH}_3\text{D}$ , by  
911 Tunable Infrared Laser Direct Absorption Spectroscopy. *Anal. Chem.* **86**, 6487–6494.

912 Penger J., Conrad R. and Blaser M. (2012) Stable carbon isotope fractionation by methylotrophic  
913 methanogenic archaea. *Appl. Environ. Microbiol.* **78**, 7596–602.

914 Piasecki A., Sessions A., Lawson M., Ferreira A. A., Neto E. V. S. and Eiler J. M. (2016) Analysis  
915 of the site-specific carbon isotope composition of propane by gas source isotope ratio mass  
916 spectrometer. *Geochim. Cosmochim. Acta* **188**, 58–72.

917 Piasecki A., Sessions A., Peterson B. and Eiler J. (2016) Prediction of equilibrium distributions of  
918 isotopologues for methane, ethane and propane using density functional theory. *Geochim.*  
919 *Cosmochim. Acta* **190**, 1–12.

920 Pine M. J. and Barker H. A. (1956) Studies on the methane fermentation. XII. The pathway of  
921 hydrogen in the acetate fermentation. *J. Bacteriol.* **71**, 644–8.

922 Rees C. E. (1973) A steady-state model for sulphur isotope fractionation in bacterial reduction  
923 processes. *Geochim. Cosmochim. Acta* **37**, 1141–1162.

924 Reeves E. P., Seewald J. S. and Sylva S. P. (2012) Hydrogen isotope exchange between n-alkanes  
925 and water under hydrothermal conditions. *Geochim. Cosmochim. Acta* **77**, 582–599.

926 Riedel, M., Collett, T.S., Malone, M.J., and Expedition 311 Scientists (2006) *Proceedings of the*  
927 *IODP Integrated Ocean Drilling Program* **311**, Integrated Ocean Drill. Program, College  
928 Station, Tex.

929 Röckmann T., Popa M. E., Krol M. C. and Hofmann M. E. G. (2016) Statistical clumped isotope  
930 signatures. *Sci. Rep.* **6**, 31947.

931 Scheller S., Goenrich M., Thauer R. K. and Jaun B. (2013) Methyl-coenzyme M reductase from  
932 methanogenic archaea: Isotope effects on the formation and anaerobic oxidation of  
933 methane. *J. Am. Chem. Soc.* **135**, 14975–14984.

934 Schleucher J., Griesinger C., Schworer B., and Thauer R. K. (1994)  $\text{H}_2$ -forming  $\text{N}^5, \text{N}^{10}$ -  
935 methylenetetrahydromethanopterin dehydrogenase from *Methanobacterium*  
936 *thermoautotrophicum* catalyzes a stereoselective hydride transfer as determined by 2-  
937 dimensional NMR- spectroscopy. *Biochemistry* **33** (13), 3986–3993.

938 Schoell M. (1980) The hydrogen and carbon isotopic composition of methane from natural gases  
939 of various origins. *Geochim. Cosmochim. Acta* **44**, 649–661.

940 Schworer B., Fernandez V. M., Zirngibl C., and Thauer R. K. (1993)  $\text{H}_2$ -forming  $\text{N}^5, \text{N}^{10}$ -  
941 methylenetetrahydromethanopterin dehydrogenase from *Methanobacterium*  
942 *thermoautotrophicum*: Studies of the catalytic mechanism of  $\text{H}_2$  formation using hydrogen  
943 isotopes. *Eur. J. Biochem.* **212** (1), 255–261.

944 Sessions A. L. and Hayes J. M. (2005) Calculation of hydrogen isotopic fractionations in  
945 biogeochemical systems. *Geochim. Cosmochim. Acta* **69**, 593–597.

946 Sim M. S., Ono S., Donovan K., Templer S. P. and Bosak T. (2011) Effect of electron donors on  
947 the fractionation of sulfur isotopes by a marine *Desulfovibrio* sp. *Geochim. Cosmochim.*  
948 *Acta* **75**, 4244–4259.

949 Srinivasan, R. and Fisher, H.F. (1985) Deuterium isotope effects for the nonenzymic and glutamate  
950 dehydrogenase catalyzed reduction of an  $\alpha$ -imino acid by NADH. *J. Am. Chem. Soc.* **107**,  
951 4301–4305.

952 Stewart L. C., Jung J., Kim Y., Kwon S., Park C., Holden J. F. and Holden J. F. (2015)  
953 *Methanocaldococcus bathoardescens* sp. nov., a hyperthermophilic methanogen isolated

- 954 from a volcanically active deep-sea hydrothermal vent. *Int. J. Syst. Evol. Microbiol.* **65**,  
955 1280-1283.
- 956 Stolper D. A., Lawson M., Davis C. L., Ferreira A. A., Neto E. V. S., Ellis G. S., Lewan M. D.,  
957 Martini A. M., Tang Y., Schoell M., Sessions A. L. and Eiler J. M. (2014) Formation  
958 temperatures of thermogenic and biogenic methane. *Science* **344**, 1500–1503.
- 959 Stolper D. A., Martini A. M., Clog M., Douglas P. M., Shusta S. S., Valentine D. L., Sessions A.  
960 L. and Eiler J. M. (2015) Distinguishing and understanding thermogenic and biogenic  
961 sources of methane using multiply substituted isotopologues. *Geochim. Cosmochim. Acta*  
962 **161**, 219–247.
- 963 Stolper D. A., Sessions A. L., Ferreira A. A., Santos Neto E. V., Schimmelmann A., Shusta S. S.,  
964 Valentine D. L. and Eiler J. M. (2013) Combined <sup>13</sup>C–D and D–D clumping in methane:  
965 Methods and preliminary results. *Geochim. Cosmochim. Acta* **126**, 169–191.
- 966 Sugimoto A. and Wada E. (1995) Hydrogen isotopic composition of bacterial methane: CO<sub>2</sub>/H<sub>2</sub>  
967 reduction and acetate fermentation. **59**, 1329–1337.
- 968 Thauer R. K. (1998) Biochemistry of methanogenesis: a tribute to Marjory Stephenson.  
969 *Microbiology* **144**, 2377–406.
- 970 Thauer R. K., Kaster A.-K., Seedorf H., Buckel W. and Hedderich R. (2008) Methanogenic  
971 archaea: ecologically relevant differences in energy conservation. *Nat. Rev. Microbiol.* **6**,  
972 579–91.
- 973 Urey H. C. (1947) The thermodynamic properties of isotopic substances. *J. Chem. Soc.*, 562.
- 974 Vanwonterghem I., Evans P. N., Parks D. H., Jensen P. D., Woodcroft B. J., Hugenholtz P. and  
975 Tyson G. W. (2016) Methylophilic methanogenesis discovered in the archaeal phylum  
976 Verstraetearchaeota. *Nat. Microbiol.* **1**, 16170.
- 977 Valentine D. L., Chidthaisong A., Rice A., Reeburgh W. S. and Tyler S. C. (2004) Carbon and  
978 hydrogen isotope fractionation by moderately thermophilic methanogens. *Geochim.*  
979 *Cosmochim. Acta* **68**, 1571–1590.
- 980 Wagner T., Kahnt J., Ermler U. and Shima S. (2016) Didehydroaspartate Modification in Methyl-  
981 Coenzyme M Reductase Catalyzing Methane Formation. *Angew. Chemie - Int. Ed.* **55**,  
982 10630–10633.
- 983 Waldron, S., Lansdown, J., Scott, E., Fallick, A. and Hall, A. (1999) The global influence of the  
984 hydrogen isotope composition of water on that of bacteriogenic methane from shallow  
985 freshwater environments. *Geochim. Cosmochim. Acta* **63**, 2237-2245.
- 986 Wang D. T., Gruen D. S., Lollar B. S., Hinrichs K., Stewart L. C., Holden J. F., Hristov A. N.,  
987 Pohlman J. W., Morrill P. L., Könneke M., Delwiche K. B. and Reeves E. P. (2015)  
988 Nonequilibrium clumped isotope signals in microbial methane. *Science* **348**, 428–431.
- 989 Wang D. T., Welander P. V. and Ono S. (2016) Fractionation of the methane isotopologues <sup>13</sup>CH<sub>4</sub>,  
990 <sup>12</sup>CH<sub>3</sub>D, and <sup>13</sup>CH<sub>3</sub>D during aerobic oxidation of methane by *Methylococcus capsulatus*  
991 (Bath). *Geochim. Cosmochim. Acta* **192**, 186–202.
- 992 Wang F.-P., Zhang Y., Chen Y., He Y., Qi J., Hinrichs K.-U., Zhang X.-X., Xiao X. and Boon N.  
993 (2014) Methanotrophic archaea possessing diverging methane-oxidizing and electron-  
994 transporting pathways. *ISME J.* **8**, 1069–1078.
- 995 Webster C. R., Mahaffy P. R., Atreya S. K., Flesch G. J., Mischna M. A., Meslin P.-Y., Farley K.  
996 A., Conrad P. G., Christensen L. E., Pavlov A. A., Martin-Torres J., Zorzano M.-P.,  
997 McConnochie T. H., Owen T., Eigenbrode J. L., Glavin D. P., Steele A., Malespin C. A.,  
998 Archer P. D., Sutter B., Coll P., Freissinet C., McKay C. P., Moores J. E., Schwenzer S.

999 P., Bridges J. C., Navarro-Gonzalez R., Gellert R. and Lemmon M. T. (2015) Mars  
1000 methane detection and variability at Gale crater. *Science* **347**, 415–417.

1001 Wecht K. J., Jacob D. J., Frankenberg C., Jiang Z. and Blake D. R. (2014) Mapping of North  
1002 American methane emissions with high spatial resolution by inversion of SCIAMACHY  
1003 satellite data. *J. Geophys. Res. Atmos.* **119**, 7741–7756.

1004 Whitehill A. R., Joelsson L. M. T., Schmidt J. A., Wang D. T., Johnson M. S. and Ono S. (2017)  
1005 Clumped isotope effects during OH and Cl oxidation of methane. *Geochim. Cosmochim.*  
1006 *Acta* **196**, 307–325.

1007 Whiticar M. (1990) A geochemical perspective of natural gas and atmospheric methane. *Org.*  
1008 *Geochem.* **16**, 531–547.

1009 Whiticar M. J. (1999) Carbon and hydrogen isotope systematics of bacterial formation and  
1010 oxidation of methane. *Chem. Geol.* **161**, 291–314.

1011 Yeung L. Y. (2016) Combinatorial effects on clumped isotopes and their significance in  
1012 biogeochemistry. *Geochim. Cosmochim. Acta* **172**, 22–38.

1013 Yoshinaga M. Y., Holler T., Goldhammer T., Wegener G., Pohlman J. W., Brunner B., Kuypers  
1014 M. M. M., Hinrichs K.-U. and Elvert M. (2014) Carbon isotope equilibration during  
1015 sulphate-limited anaerobic oxidation of methane. *Nat. Geosci.* **7**, 190–194.

1016 Yoshioka H., Sakata S. and Kamagata Y. (2008) Hydrogen isotope fractionation by  
1017 *Methanothermobacter thermoautotrophicus* in coculture and pure culture conditions.  
1018 *Geochim. Cosmochim. Acta* **72**, 2687–2694.

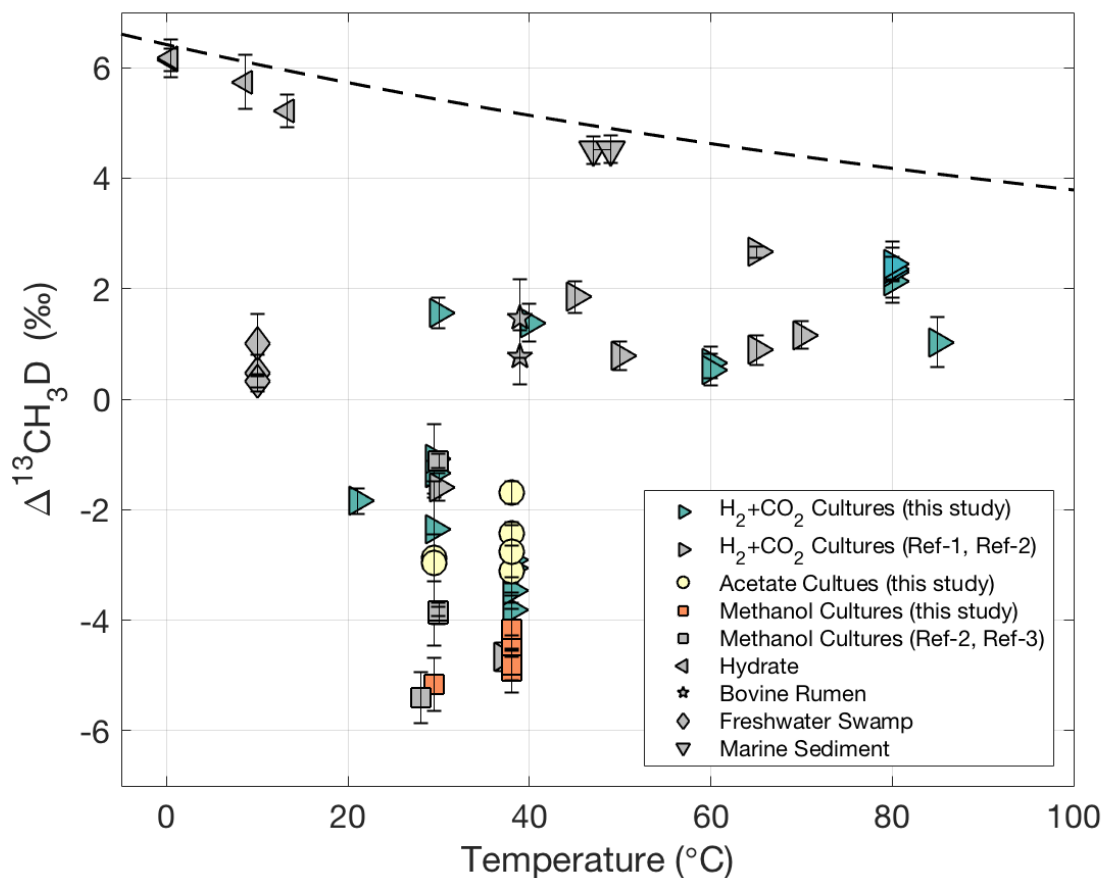
1019 York D., Evensen, N.M., Martinez, M.L. and Delgado, J.D.B. (2004) Unified equations for the  
1020 slope, intercept, and standard errors of the best straight line. *Am. J. Phys.* **72**, 367-375.

1021 Young E. D., Rumble D., Freedman P. and Mills M. (2016) A large-radius high-mass-resolution  
1022 multiple-collector isotope ratio mass spectrometer for analysis of rare isotopologues of O<sub>2</sub>,  
1023 N<sub>2</sub>, CH<sub>4</sub> and other gases. *Int. J. Mass Spectrom.* **401**, 1–10.

1024 Young E. D., Kohl I. E., Lollar B. S., Etiope G., Rumble D., Li S., Haghnegahdar M. a., Schauble  
1025 E. a., McCain K. A., Foustoukos D. I., Sutcliffe C., Warr O., Ballentine C. J., Onstott T. C.,  
1026 Hosgormez H., Neubeck A., Marques J. M., Pérez-Rodríguez I., Rowe A. R., LaRowe D.  
1027 E., Magnabosco C., Yeung L. Y., Ash J. L. and Bryndzia L. T. (2017) The relative  
1028 abundances of resolved <sup>12</sup>CH<sub>2</sub>D<sub>2</sub> and <sup>13</sup>CH<sub>3</sub>D and mechanisms controlling isotopic bond  
1029 ordering in abiotic and biotic methane gases. *Geochim. Cosmochim. Acta* **203**, 235–264.

1030

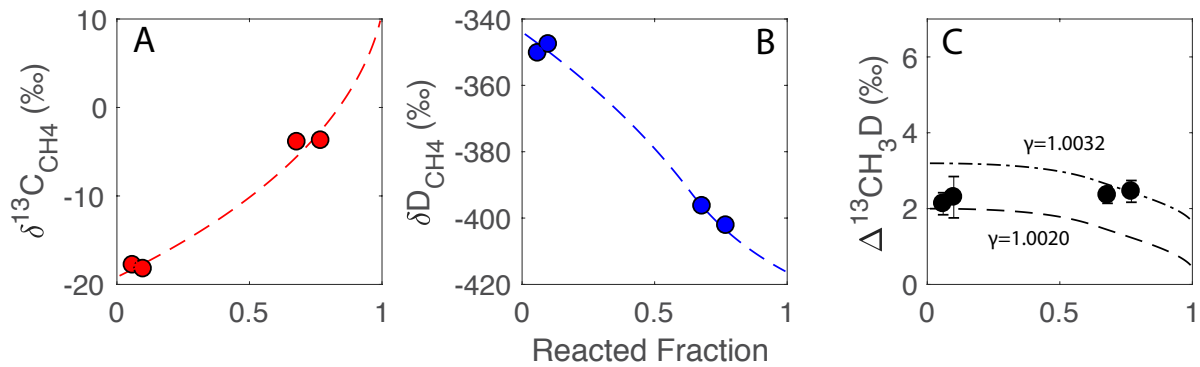
1031



1033  
 1034 **Figure 1. Clumped methane,  $\Delta^{13}\text{CH}_3\text{D}$ , plotted against environmental temperatures.** The  
 1035 dashed line represents the equilibrium  $\Delta^{13}\text{CH}_3\text{D}$  values calibrated experimentally using heated  
 1036 methane calibrations between 400 and 150°C (Ono et al., 2014; Wang et al., 2015) and  
 1037 extrapolated for lower temperatures. Colored triangles, circles, and squares represent laboratory  
 1038 cultures from this study. Right-facing triangles refer to  $\text{H}_2+\text{CO}_2$  cultures, circles to acetate cultures  
 1039 and squares to methanol cultures. A subset of samples was previously published in Wang et. (2015)  
 1040 as noted in Table 2. For comparison to this work, previously reported  $\text{H}_2+\text{CO}_2$  and methanol  
 1041 cultures from Stolper et al., 2014 (Ref-1), Young et al., 2017 (Ref-2) and Douglas et al., 2016  
 1042 (Ref-3) are plotted with grey symbols which correspond to the substrate used. Also plotted are  
 1043 environmental methane samples reported in Wang et al. (2015) (Grey symbols). Bovine rumen  
 1044 samples are published in Lopes et al. (2016). In situ temperatures for hydrate samples are  
 1045 calculated using depths and geothermal gradients listed in IODP reports (Riedel et al., 2006). All  
 1046 previously published culture data are reported according to their original measurement notation  
 1047 ( $\Delta^{13}\text{CH}_3\text{D}$  or  $\Delta 18$ ).

1048  
 1049  
 1050  
 1051

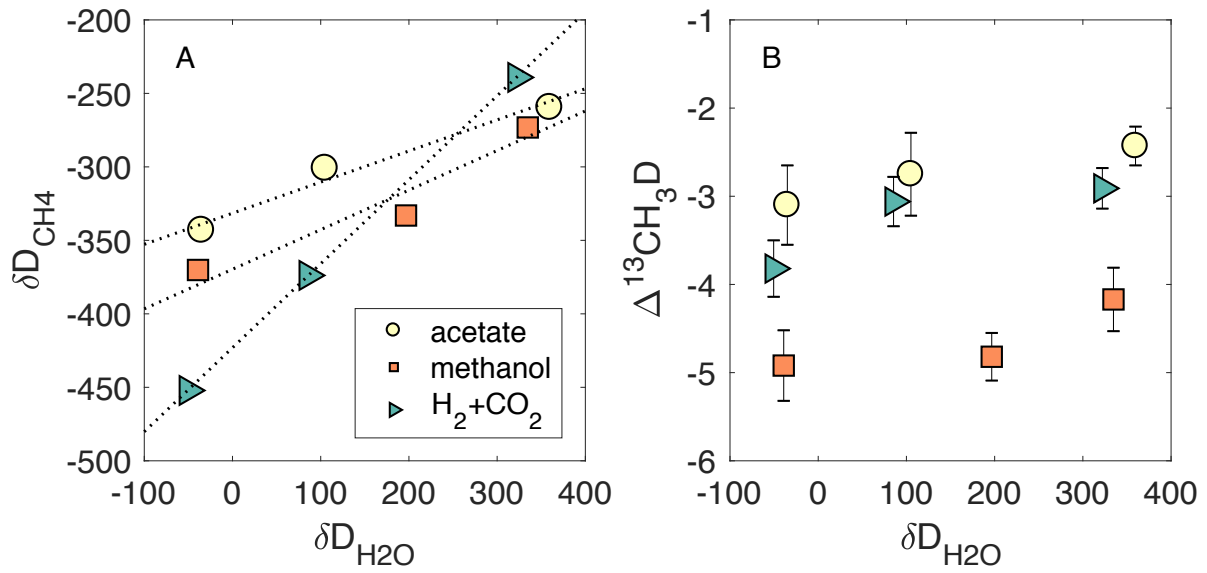
1052



1053  
1054  
1055  
1056  
1057  
1058  
1059  
1060  
1061  
1062

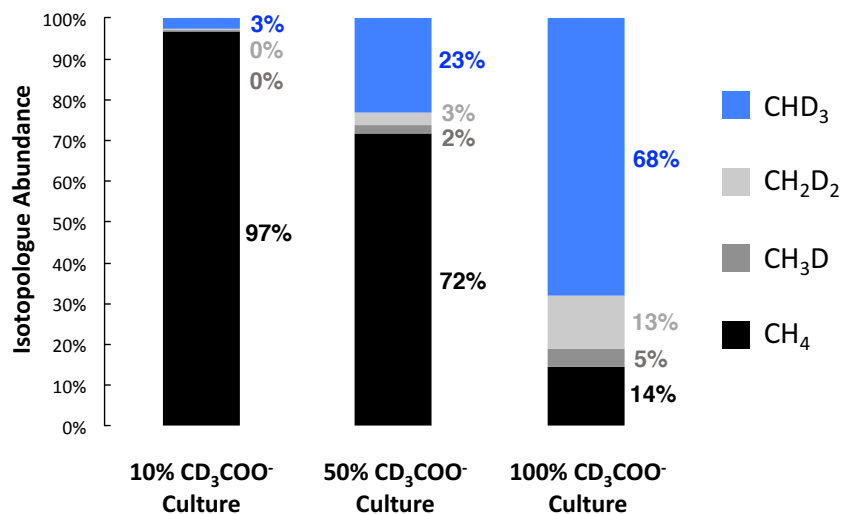
**Figure 2. Isotope systematics of  $\delta^{13}\text{C}$ ,  $\delta\text{D}$ ,  $\Delta^{13}\text{CH}_3\text{D}$  over the course of a batch culture experiment.** Hydrogenotrophic methanogens (*Methanocaldococcus bathoardescens*) grown at 80°C. Reacted fraction refers to the fraction of carbon dioxide converted to methane. The filled circles represent methane measured from culture experiments. The dashed lines show the results of a closed system model discussed in section 4.1.  $\delta^{13}\text{C}$ ,  $\delta\text{D}$  and  $\Delta^{13}\text{CH}_3\text{D}$  of methane are shown in A, B, and C, respectively. Results for two  $\gamma$  values are shown in C. See text for other fractionation factors. The  $\delta^{13}\text{C}$  value of initial  $\text{CO}_2$  was fitted to 10.9‰.





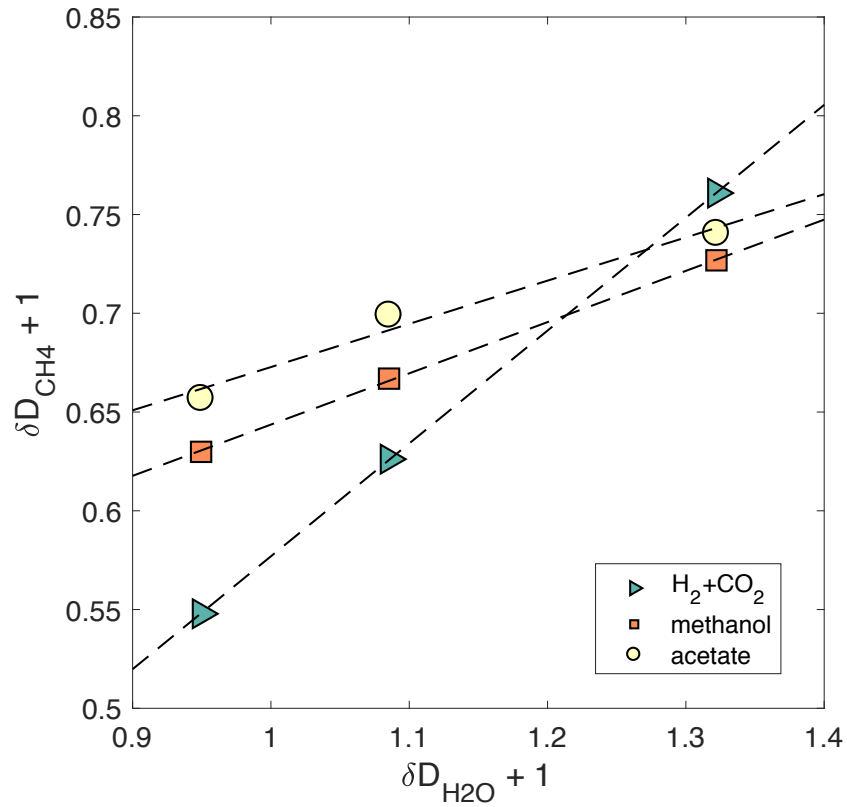
1063  
 1064  
 1065  
 1066  
 1067  
 1068

**Figure 3. The  $\delta D$  and  $\Delta^{13}CH_3D$  values of methane produced by *M. barkeri* in deuterium-spiked medium.** Cultures were grown at 38°C, and isotopic compositions of methane are compared against  $\delta D$  values of media water in D-spiked experiments. Includes Set 2 data.



1069  
 1070  
 1071  
 1072  
 1073  
 1074  
 1075

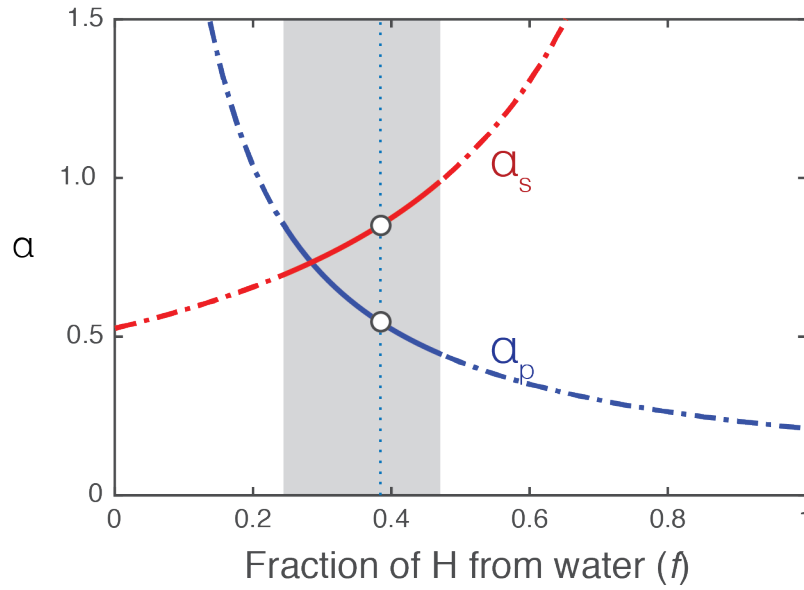
**Figure 4. Relative isotopologue abundances of methane produced by *M. barkeri* in batch cultures spiked with CD<sub>3</sub>COO<sup>-</sup> (10, 50, 100%).** Isotopologue composition was determined by FTIR, and relative abundances of methane-d<sub>n</sub> isotopologues were calibrated against high-purity synthetic standards. CD<sub>4</sub> was not detected in any of the experiments.



1076  
 1077  
 1078  
 1079  
 1080  
 1081  
 1082  
 1083

**Figure 5. The hydrogen isotope composition of microbial methane vs. the deuterium composition of media water in spiked culture experiments. *M. barkeri* grown on three different substrates. In comparison to Figure 3A, the axes in this figure are plotted as  $\delta D + 1$  to take into account the non-linearity in  $\delta D$ .**

1084



1085

1086

1087 **Figure 6. The fractionation factor ( $\alpha$ ) as a function of the fraction of hydrogen derived from**

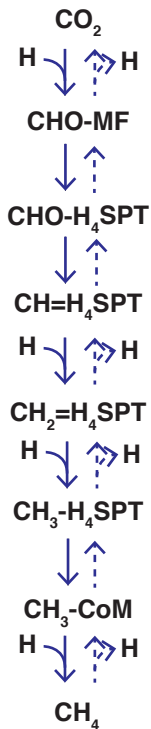
1088 **water ( $f$ ) in acetoclastic methanogenesis in  $D_2O$  spiked water experiments. Grey shading**

1089 **indicates the range of  $f$  values that satisfy  $\alpha_p$  and  $\alpha_s < 1$  and  $0.25 < f < 1$ . The dashed vertical line and**

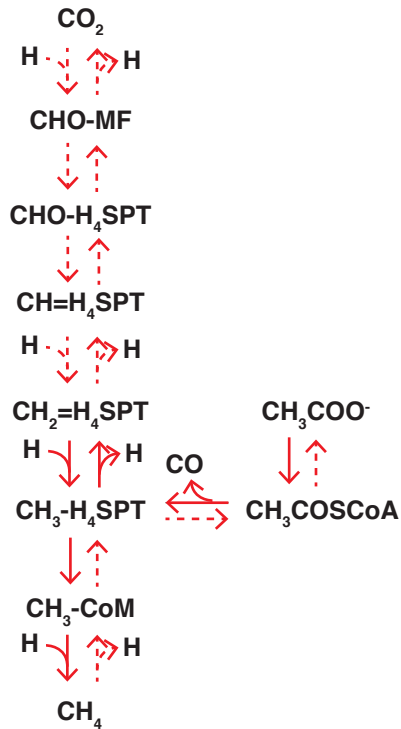
1090 **open circles correspond to the solution when  $\alpha_s = 0.85$ .**

1090

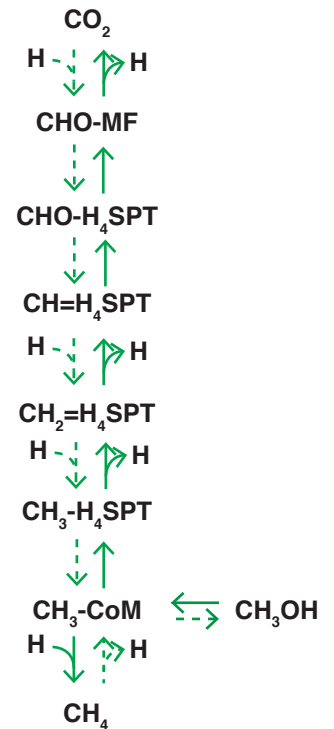
## 1) Hydrogenotrophic Pathway



## 2) Acetoclastic Pathway



## 3) Methylotrophic Pathway



1092

1093

1094

**Figure 7: Three methanogenic pathways for *Methanosarcina*.** The solid arrows represent the predominant direction of the reaction, and the dashed arrows represent the backward reaction, which is thought to be minor. Two solid arrows are used to indicate that the reactions are thought to be reversible, according to the literature, not specifically evaluated in this study. In the

hydrogenotrophic pathway (blue),  $\text{CO}_2$  is reduced by a series of four two-electron processes, each adding one H atom.  $\text{C}_1$  compounds are carried with cofactors (MF, methanofuran,  $\text{H}_4\text{SPT}$ , tetrahydrosarcinaopterin, CoM, coenzyme M). For the acetoclastic pathway, acetate is first

activated by acetyl co-A (via acetyl phosphate), acetyl co-A is split to methyl ( $\text{CH}_3-$ ) and a carbonyl moiety (CO), and the latter is oxidized to  $\text{CO}_2$ . The methylotrophic pathway is overall a

disproportionation reaction in which one methyl group is oxidized to  $\text{CO}_2$  via a reversed methanogenic pathway, and three additional methyl groups are reduced to methane.

1098

1099

1100

1101

1102

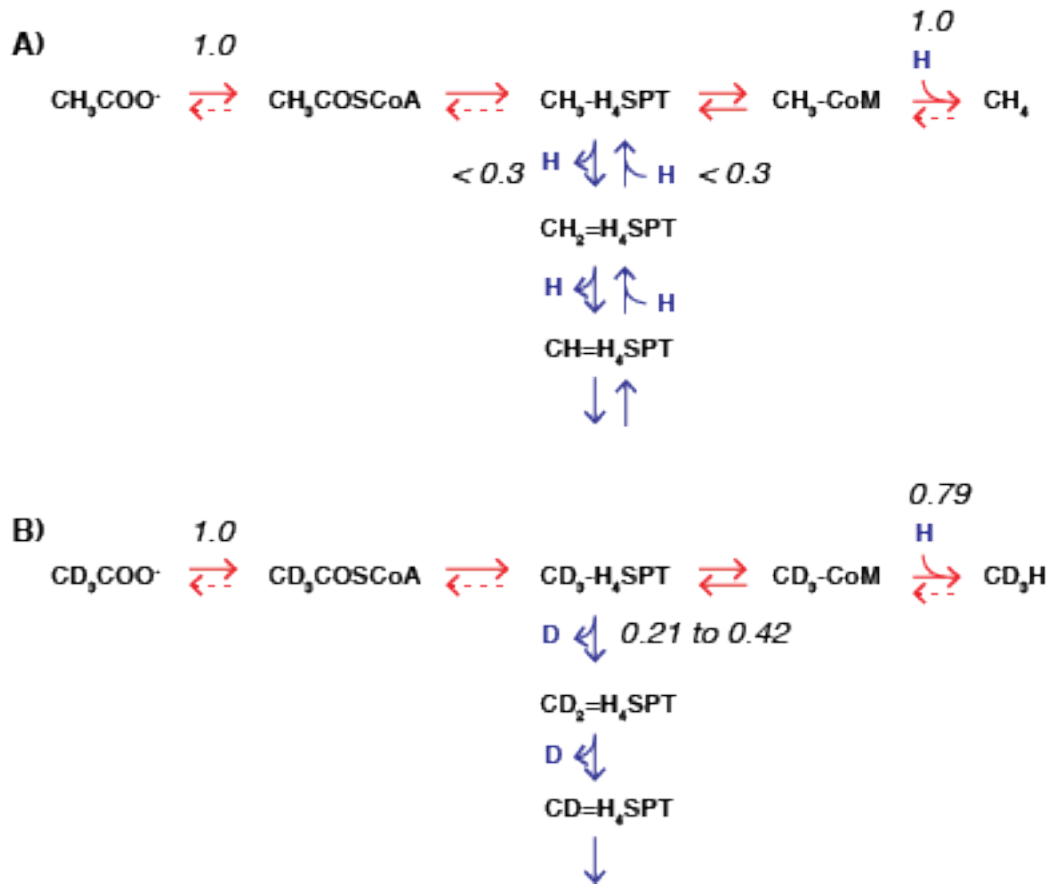
1103

1104

1105

1106

1107



1108

1109

1110 **Figure 8: Proposed mechanism of D-isotope exchange by acetate metabolism by *M. barkeri*.**

1111 The numbers in italics are relative fluxes estimated from A) D-spiked medium, and B)  $\text{CD}_3\text{COO}^-$

1112 experiments, respectively. Reactions with two solid lines indicate that the reaction is thought to be

1113 reversible, whereas reactions with solid and dashed lines indicate that the reaction proceeds

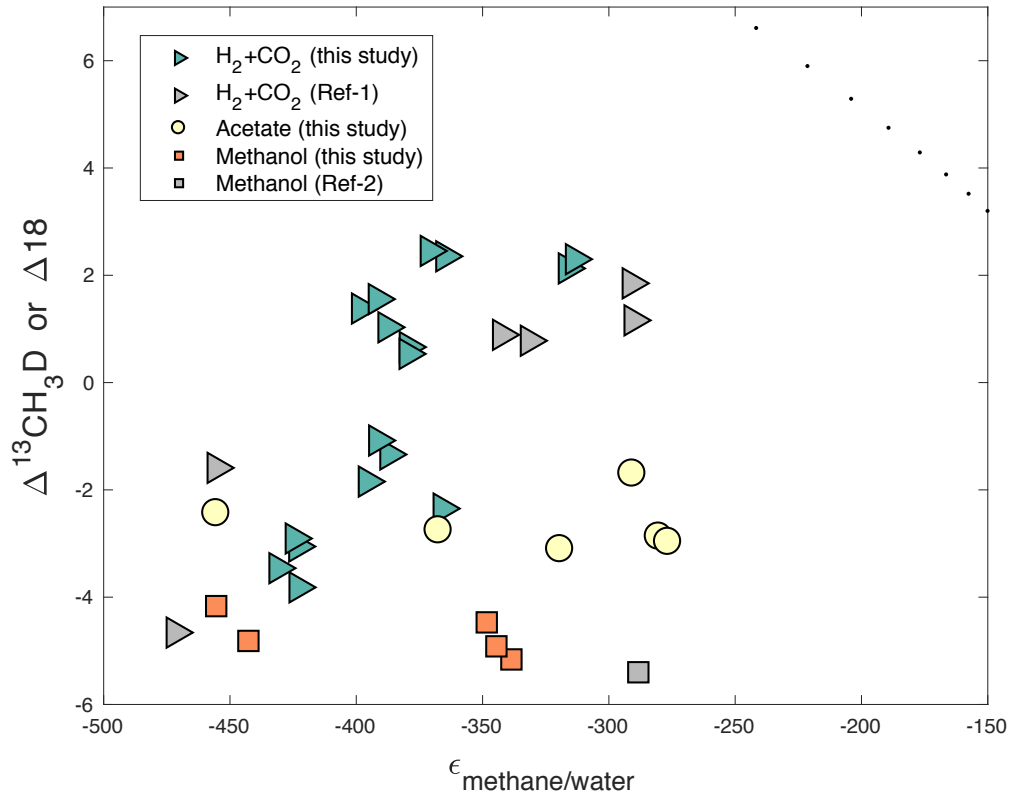
1114 predominantly the direction of the solid arrow. As in Figure 7, the solid arrows represent the

1115 predominant direction of the reaction, and the dashed arrows represent the backward reaction. Red

1116 lines indicate the reactions required for the acetoclastic pathway. Blue lines indicate side reactions.

1117

1118



1119  
 1120  
 1121  
 1122  
 1123  
 1124  
 1125

**Figure 9. The values of D13CH3D (or D18) of methane plotted against D/H fractionation factor between methane and water ( $\epsilon_{\text{methane/water}}$ ) for methanogen cultures grown on different substrate. All culture data previously published with corresponding  $\delta\text{D}_{\text{H}_2\text{O}}$  values is included for reference (grey symbols). Ref-1 refers to Stolper et al. (2015) and Ref-2 refers to Douglas et al. (2016). Black dots represent isotopic equilibrium (Wang et al., 2015).**

1126 **Table 1. Summary of experiments.**

Experiment (Method)	Organism(s)	Purpose	Variables	Substrates Used	Presented in
Temperature Series* (2.1.1)	<i>Methanocaldococcus jannaschii</i> , <i>Methanocaldococcus bathoardescens</i> , <i>Methanothermococcus thermolithotrophicus</i>	Effect of growth temperature	Growth Temperature	H <sub>2</sub> +CO <sub>2</sub>	Figures 1, 9 Tables 2, 3
Time Series (2.1.2)	<i>Methanocaldococcus bathoardescens</i>	Closed system isotope effects	Incubation Time	H <sub>2</sub> +CO <sub>2</sub> ,	Figures 1, 2, 9 Tables 2, 3
Substrate Series (2.1.3)	<i>Methanosarcina barkeri</i> , <i>Methanosarcina mazei</i>	Substrate & Pathway effects	Substrate	H <sub>2</sub> +CO <sub>2</sub> , acetate, methanol	Figures 1, 3, 4, 5, 9 Tables 2, 3
Deuterated Water (2.1.4)	<i>Methanosarcina barkeri</i>	Hydrogen source	δD Water	H <sub>2</sub> +CO <sub>2</sub> , acetate, methanol	Figures 1, 3, 5, 6, 8, 9 Tables 2, 3
Deuterated Acetate (2.1.5)	<i>Methanosarcina barkeri</i>	Hydrogen source	CD <sub>3</sub> COOD Spike	acetate	Figures 1, 4, 8, 9 Tables 2, 3

1127 \*A part of this data was reported in Wang et al. (2015) as noted in Table 2.

1128



1129

**Table 2. Results for methanogen culture experiments.**

Methanogen	Substrate* <sup>1</sup>	T (°C)	δD <sub>H2O</sub>	δ <sup>13</sup> C <sub>CH4</sub>	+/- ‰	δD <sub>CH4</sub>	+/- ‰	Δ <sup>13</sup> CH3D	+/- ‰	Incubation (hrs)	Methane (mL) <sup>7</sup>
<b>Temperature Series Experiments (TS)</b>											
<i>M. jannaschii</i> <sup>*f</sup>	H <sub>2</sub> +CO <sub>2</sub>	80	-49.0	-18.79	0.03	-415.46	0.05	2.29	0.23	5.25 <sup>*5</sup>	NA
<i>M. bathoardescens</i> <sup>*f</sup>	H <sub>2</sub> +CO <sub>2</sub>	85	-49.0	-12.58	0.07	-417.80	0.07	1.03	0.45	8 <sup>*5</sup>	NA
<i>M. thermolithotrophicus</i> <sup>*f</sup>	H <sub>2</sub> +CO <sub>2</sub>	40	-49.0	-16.47	0.04	-427.76	0.04	1.38	0.34	28 <sup>*5</sup>	NA
<i>M. thermolithotrophicus</i>	H <sub>2</sub> +CO <sub>2</sub>	30	-49.0	-17.05	0.08	-421.44	0.12	1.56	0.28	64 <sup>*5</sup>	NA
<i>M. thermolithotrophicus</i>	H <sub>2</sub> +CO <sub>2</sub>	60	-49.0	-17.15	0.06	-409.77	0.05	0.66	0.28	6 <sup>*5</sup>	NA
<i>M. thermolithotrophicus</i> <sup>*f</sup>	H <sub>2</sub> +CO <sub>2</sub>	60	-49.0	-17.05	0.05	-409.84	0.05	0.54	0.28	6 <sup>*5</sup>	NA
<b>Time Series Experiments (CS)</b>											
<i>M. bathoardescens</i>	H <sub>2</sub> +CO <sub>2</sub> (6%)	80	-49.6	-17.82	0.07	-350.30	0.21	2.13	0.29	2.75	NA
<i>M. bathoardescens</i>	H <sub>2</sub> +CO <sub>2</sub> (10%)	80	-49.6	-18.25	0.10	-347.63	0.10	2.30	0.55	3.25	NA
<i>M. bathoardescens</i>	H <sub>2</sub> +CO <sub>2</sub> (68%)	80	-49.6	-3.91	0.04	-396.41	0.04	2.35	0.21	3.5	NA
<i>M. bathoardescens</i>	H <sub>2</sub> +CO <sub>2</sub> (77%)	80	-49.6	-3.74	0.03	-402.25	0.03	2.45	0.29	4	NA
<b>Methanosarcina Substrate, Spike, and Temperature Experiments - set 1<sup>3</sup></b>											
<i>M. barkeri</i> <sup>*f</sup>	H <sub>2</sub> +CO <sub>2</sub>	21-38	-51.2	-59.90	0.05	-418.40	0.05	-1.34	0.22	336 <sup>*6</sup>	NA
<i>M. barkeri</i> <sup>*f</sup>	H <sub>2</sub> +CO <sub>2</sub>	21-38	-51.2	-59.30	0.07	-422.67	0.07	-1.08	0.63	336 <sup>*6</sup>	NA
<i>M. barkeri</i>	H <sub>2</sub> +CO <sub>2</sub>	21-38	100 <sup>*3</sup>	-59.15	0.06	-340.47	0.05	-1.32	0.23	336 <sup>*6</sup>	NA
<i>M. barkeri</i>	H <sub>2</sub> +CO <sub>2</sub>	21-38	260 <sup>*3</sup>	-60.93	0.11	-201.10	0.11	-2.35	0.56	336 <sup>*6</sup>	NA
<i>M. barkeri</i>	methanol	21-38	-51.2	-116.30	0.11	-372.46	0.11	-5.16	0.48	336 <sup>*6</sup>	NA
<i>M. barkeri</i>	acetate	21-38	-51.2	-66.83	0.08	-317.08	0.09	-2.87	0.42	336 <sup>*6</sup>	NA
<i>M. barkeri</i>	acetate	21-38	-51.2	-66.78	0.34	-313.61	0.55	-2.97	1.49	336 <sup>*6</sup>	NA
<b>Methanosarcina Substrate, Spike, and Temperature Experiments - set 2</b>											
<i>M. barkeri</i>	H <sub>2</sub> +CO <sub>2</sub>	38	85.4	-62.02	0.05	-373.86	0.05	-3.06	0.28	730	22
<i>M. barkeri</i>	H <sub>2</sub> +CO <sub>2</sub>	38	322.1	-58.08	0.05	-239.13	0.05	-2.91	0.23	730	39
<i>M. barkeri</i>	H <sub>2</sub> +CO <sub>2</sub>	38	-50.5	-57.40	0.05	-452.14	0.05	-3.82	0.32	730	30
<i>M. barkeri</i> <sup>*2</sup>	H <sub>2</sub> +CO <sub>2</sub>	21	-49.4	-70.52	0.05	-425.43	0.05	-1.85	0.23	730	15
<i>M. barkeri</i>	methanol	38	-39.1	-117.01	0.05	-370.16	0.05	-4.92	0.40	72	22
<i>M. barkeri</i>	methanol	38	196.7	-116.32	0.06	-333.09	0.06	-4.82	0.27	72	33
<i>M. barkeri</i>	methanol	38	334.7	-118.10	0.05	-273.19	0.06	-4.17	0.36	72	96
<i>M. barkeri</i>	acetate	38	105.1	-73.74	0.07	-300.93	0.07	-2.75	0.47	730	11
<i>M. barkeri</i>	acetate	38	359.7	-67.68	0.05	-259.54	0.05	-2.43	0.22	730	9
<i>M. barkeri</i>	acetate (+YE)	38	-35	-72.86	0.10	-343.12	0.10	-3.10	0.45	730	12
<i>M. mazei</i>	H <sub>2</sub> +CO <sub>2</sub>	38	-51.1	-56.07	0.05	-460.03	0.05	-3.46	0.22	730	23
<i>M. mazei</i>	acetate	38	-34.1	-60.40	0.05	-314.90	0.05	-1.70	0.20	730	11
<i>M. mazei</i>	methanol	38	-23	-120.82	0.04	-363.34	0.04	-4.47	0.20	72	86

1130

\*<sup>1</sup> Substrate (%) refers to the percent of substrate consumed or reaction completion.

1131

\*<sup>2</sup> The culture was grown at 21°C and excluded from Figure 3, 4 and analysis.

1132

\*<sup>3</sup> δD<sub>H2O</sub> values represent estimates. These were measured for Set 2.

1133

\*<sup>4</sup> From Wang et al., 2015

1134

\*<sup>5</sup> Values represent an estimate based on previous culture data using this this culture strain grown at similar temperatures.

1135

\*<sup>6</sup> Values represent estimated incubation time.

1137

\*<sup>7</sup> Volume of methane (STP) in the culture headspace at the end of the incubation time. NA is recorded for all cultures that were not injected with NaOH at the end of incubation and prior to isotope measurements.

1138

1139

1140

1141

1142

1143

1144 **Table 3. Isotopic compositions of substrates and medium water.**

<b>Material</b>	<b><math>\delta^{13}\text{C}</math> ‰</b>	<b><math>\delta\text{D}_{\text{H}_2\text{O}}</math> ‰</b>	<b><math>\delta\text{D}_{\text{CH}_3}</math> ‰</b>
H <sub>2</sub> - CO <sub>2</sub> gas mix* <sup>1</sup>	-34.4		
N <sub>2</sub> - CO <sub>2</sub> gas mix* <sup>1</sup>	-35.8		
CH <sub>3</sub> COONa <sup>1</sup>	-40.2		-123
CH <sub>3</sub> OH* <sup>1</sup>	-49.5		
Bremen DI Water* <sup>2</sup>		-51.2	
UMass DI Water* <sup>3</sup>		-49.6	

1145

1146 \*1: These materials were used for Set 2 of substrate and D-spike experiments.

1147 \*2: Bremen DI water (pre-inoculation) was used in Set 2 of substrate and D-spike experiments.

1148 \*3: UMass DI water (pre-inoculation) was used in the temperature and time-series experiments.

1149

1150

1151

1152