1	Inverse Kinetic Isotope Fractionation During Bacterial Nitrite Oxidation
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13	FINAL DRAFT ACCEPTED BY GCA
14	DECEMBER 29, 2008
15	

ABSTRACT

16 Natural abundance stable isotopes in nitrate (NO_3^{-}), nitrite (NO_2^{-}), and nitrous oxide (N_2O) 17 have been used to better understand the cycling of nitrogen in marine and terrestrial 18 environments. However, in order to extract the greatest information from the distributions of 19 these isotopic species, the kinetic isotope effects for each of the relevant microbial reactions are 20 needed. To date, kinetic isotope effects for nitrite oxidation and anaerobic ammonium oxidation 21 (anammox) have not been reported. In this study, the nitrogen isotope effect was measured for 22 microbial nitrite oxidation to nitrate. Nitrite oxidation is the second step in the nitrification 23 process, and it plays a key role in the regeneration of nitrate in the ocean. Surprisingly, nitrite 24 oxidation occurred with an inverse kinetic isotope effect, such that the residual nitrite became progressively depleted in ¹⁵N as the reaction proceeded. Three potential explanations for this 25 26 apparent inverse kinetic isotope effect were explored: 1) isotope exchange equilibrium between 27 nitrite and nitrous acid prior to reaction, 2) reaction reversibility at the enzyme level, and 3) true 28 inverse kinetic fractionation. Comparison of experimental data to *ab initio* calculations and 29 theoretical predictions leads to the conclusion that the fractionation is most likely inverse at the 30 enzyme level. Inverse kinetic isotope effects are rare, but the experimental observations reported 31 here agree with kinetic isotope theory for this simple N-O bond-forming reaction. Nitrite 32 oxidation is therefore fundamentally different from all other microbial processes in which N 33 isotope fractionation has been studied. The unique kinetic isotope effect for nitrite oxidation 34 should help to better identify its role in the cycling of nitrite in ocean suboxic zones, and other 35 environments in which nitrite accumulates.

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36 1. INTRODUCTION

37 1.1 Application of Isotopes to Nitrogen Biogeochemistry

38 Stable isotope ratio measurements of nitrogen (N)-containing molecules have long provided 39 insights into the relative rates of biogeochemical processes comprising the global nitrogen cycle. The kinetic fractionation factors ($\alpha_k = {}^{14}k/{}^{15}k$) and isotope effects ($\varepsilon_k = (\alpha_k - 1)*1000$) for key 40 41 microbial processes offer constraints on the mechanisms of enzymatic reactions and provide the 42 basis for interpretation of natural abundance N isotopic distributions in nature. For example, the 43 kinetic isotope effects for both nitrogen fixation and denitrification, which add and remove 44 bioavailable N from the ocean have been used to estimate the relative rates of these fluxes and 45 thus the residence time of fixed N in the ocean (BRANDES and DEVOL, 2002; DEUTSCH et al., 46 2004; ALTABET, 2007).

47 The kinetic isotope effect for nitrogen fixation is small (-2 to +2%; Table 1), which leads to addition of bioavailable N to the ocean with a δ^{15} N value near its source value of 0‰. Other 48 49 sources of N to the ocean from atmospheric deposition and continental runoff are also believed 50 to be near 0‰, although this value is highly uncertain (BRANDES and DEVOL, 2002). The 51 processes that remove bioavailable N from the ocean include denitrification and anammox in 52 suboxic regions within the water column and marine sediments (recently reviewed by 53 DALSGAARD et al., 2005; FRANCIS et al., 2007). If the ocean N budget is in steady state, N would be removed by these processes with a flux-weighted average δ^{15} N value near 0% (ALTABET, 54 2007: BRANDES and DEVOL. 2002: DEUTSCH et al., 2004: DEVOL et al., 2006). While the isotope 55 56 effects for anammox are not known at this time, the isotope effects for water column (20-30‰) and sedimentary (~ 0 %) denitrification have been used to partition N₂ losses between the water 57 column and sediments. The large isotope effect for water column denitrification is a strong 58

constraint on this partitioning and implies that water column denitrification represents a
relatively small proportion of the total N₂ flux (BRANDES and DEVOL, 2002; DEUTSCH et al.,
2004; SIGMAN et al., 2003).

62 The large fractionation factor for nitrate reduction (BARFORD et al., 1999; DELWICHE and STEYN, 1970; GRANGER et al., 2006) also leads to ¹⁵N enrichment of nitrate in ocean suboxic 63 64 zones (BRANDES et al., 1998; CLINE and KAPLAN, 1975; SIGMAN et al., 2005; SIGMAN et al., 2003; Voss et al., 2001). Recently, Casciotti and McIlvin (CASCIOTTI and MCILVIN, 2007) 65 reported $\delta^{15}N_{NO3}$ values up to +21.7‰ (and $\delta^{18}O_{NO3}$ values up to +19.0‰) in the suboxic zone of 66 67 the eastern tropical North Pacific (ETNP). With a large fractionation factor for nitrate reduction to nitrite, it would be expected for NO₂⁻ to be depleted in ¹⁵N relative to NO₃⁻. However, $\delta^{15}N_{NO2}$ 68 69 values within the suboxic zone were as low as -18.5% (CASCIOTTI and MCILVIN, 2007). The difference in $\delta^{15}N$ between nitrate and nitrite ($\Delta\delta^{15}N = \delta^{15}N_{NO3} - \delta^{15}N_{NO2}$) ranged from +28.4 to 70 71 +35.0% within the suboxic zone, with maximal values near the top of the secondary nitrite maximum. These differences in δ^{15} N are larger than expected from the fractionation factor for 72 73 nitrate reduction during denitrification (20-30%; Table 1), particularly if there is concomitant nitrite consumption that acts to increase $\delta^{15}N_{NO2}$. These data suggest that either the fractionation 74 75 factor for nitrate reduction in the water column is larger than previous estimates, with 76 confounding effects on isotope-based oceanic nitrogen budgets (above), or that co-occurring processes act to increase $\Delta \delta^{15}$ N. Such processes would need to decrease δ^{15} N_{NO2} values relative 77 to $\delta^{15}N_{NO3}$, or increase $\delta^{15}N_{NO3}$ values while causing a minimal increase in $\delta^{15}N_{NO2}$. 78 79 In the suboxic zone, nitrite can be produced by nitrate reduction (CODISPOTI and 80 CHRISTENSEN, 1985; CODISPOTI et al., 1986; CODISPOTI and RICHARDS, 1976) or possibly 81 ammonia-oxidation (LIPSCHULTZ et al., 1990; WARD et al., 1989). The isotope effects for these

82 processes range from +13-38‰ (Table 1), but would be expressed to differing degrees depending 83 on the extent of reaction completion. Nitrite consumption may be due to numerous processes, 84 such as nitrite reduction to nitric oxide (NO) by denitrifying bacteria, nitrite reduction to N_2 by 85 anammox bacteria (DALSGAARD et al., 2003; HAMERSLEY et al., 2007; KUYPERS et al., 2005; 86 KUYPERS et al., 2003), or nitrite oxidation to nitrate by nitrite oxidizing bacteria (ANDERSON et 87 al., 1982; CASCIOTTI and MCILVIN, 2007; LIPSCHULTZ et al., 1990; SIGMAN et al., 2005). Nitrite reduction to NO has been shown to occur with a normal kinetic isotope effect ($^{15}\varepsilon_{k \text{ NIR}} \approx +15\%$; 88 (BRYAN et al., 1983)), which would increase $\delta^{15}N_{NO2}$ relative to $\delta^{15}N_{NO3}$ and lead to low $\Delta\delta^{15}N$ 89 90 values. In order to better interpret N (and O) isotopic distributions of nitrate and nitrite in 91 suboxic zones, the isotope effects for anammox and nitrite oxidation are needed. In this study, 92 the kinetic isotope effect for nitrite oxidation was determined for pure cultures of the marine 93 nitrite-oxidizing bacterium Nitrococcus mobilis.

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95 **1.2. Isotopic Fractionation**

96 The differential distribution of isotopes between chemical species or phases, known as isotopic fractionation, has been recognized for many decades and is the basis for many 97 98 applications in geochemistry (BIGELEISEN, 1952; BIGELEISEN, 1965; BIGELEISEN and MAYER, 99 1947; UREY, 1947). Both equilibrium and kinetic isotope effects can lead to isotopic fractionation between phases or chemical species. The equilibrium fractionation factor, $\alpha_{eq} =$ 100 101 $(A_2/A_1)/(B_2/B_1)$, defines the isotope ratios that two species, 'A' and 'B', would have at 102 equilibrium, where '1' and '2' refer to the light and heavy isotopes, respectively. Equilibrium 103 fractionation arises primarily from differences in the zero point energies of the molecules 104 undergoing isotopic exchange, leading to enrichment of the heavy isotopes in the more strongly

bonded form (BIGELEISEN and MAYER, 1947; SCHAUBLE et al., 2001). This can be shown using
the molecular partition functions, which represent the distribution of the relevant molecules
between their available rotational, translational, and vibrational energy states (UREY, 1947;

108 BIGELEISEN and MAYER, 1947; RICHET et al., 1977).

109 Kinetic fractionation factors (α_k) arise from small differences in the rates at which

110 isotopically substituted molecules react, and are often the focus of stable isotope applications in

111 biochemistry. Kinetic fractionation factors are defined by the instantaneous change in the isotope

ratio of the reaction product (R_{Pi}) at a given substrate isotope ratio (R_S): $\alpha_k = R_S/R_{Pi}$ (MARIOTTI

113 et al., 1981; SCOTT et al., 2004). The kinetic fractionation factor is also often represented by the

114 ratio of rate constants, assuming first order reaction kinetics apply: $\alpha_k = k_1/k_2$, where '1' and '2'

115 refer to the light and heavy isotopes, respectively (BIGELEISEN and WOLFSBERG, 1958).

116 For a reversible reaction, the ratio of kinetic isotope effects for forward and reverse reactions

117 is also related to the equilibrium isotope effect: $\alpha_{eq} = \alpha_{kf}/\alpha_{kr}$, where α_{kf} is the kinetic isotope

118 effect for the forward reaction and α_{kr} is the kinetic isotope effect for the reverse reaction

119 (O'LEARY, 1981). Given that the inherent fractionation factors α_{eq} , α_{kf} , and α_{kr} are constant, this 120 relation should hold at all times. However, the expression of this equality in the isotope ratios of 121 substrates and products only holds at equilibrium.

Kinetic fractionation factors can also be calculated from first principles using transition state theory, where the substrate is assumed to be in thermal equilibrium with the transition state of the reaction (BIGELEISEN, 1949; BIGELEISEN, 1952; BUDDENBAUM and SHINER, 1977). If the isotopically substituted species is more stably bonded in the substrate than in the transition state, then molecules containing the heavy isotope will react more slowly. This behavior is termed "normal" kinetic isotope fractionation, and typically occurs in bond-breaking reactions. On the

128 other hand, if the isotopically substituted species is more stably bonded in the transition state, 129 which may be the case in some bond-forming reactions, then molecules containing the heavy 130 isotope may react more quickly. This phenomenon is referred to as "inverse" kinetic isotope 131 fractionation, and may be expected to occur in some bond-forming reactions (FRY, 1970). In this study, measurements of nitrate and nitrite δ^{15} N in pure culture experiments show that 132 133 microbial nitrite oxidation occurs with an inverse kinetic isotope effect. Three potential 134 mechanisms that could lead to this unusual fractionation are discussed: 1) a reversible reaction 135 prior to the nitrite oxidation step ('pre-equilibrium'), 2) reversibility of the nitrite oxidation 136 reaction itself ('enzyme reversibility'), and 3) a true enzyme-level inverse kinetic isotope effect. 137 Examination of each of these mechanisms by comparison of experimental data to equilibrium 138 and kinetic isotope fractionation theory suggests that this relatively simple bond-forming 139 reaction proceeds with a true enzyme-level inverse kinetic isotope effect. Although inverse 140 kinetic isotope effects are unusual in biochemistry, here I argue that its occurrence in nitrite 141 oxidation should be expected and may have important implications for understanding marine 142 nitrogen isotope biogeochemistry and nitrite cycling in suboxic zones.

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144 **2. MATERIALS AND METHODS**

145 **2.1 Growth and Maintenance of Bacterial Cultures**

Nitrococcus mobilis, a marine nitrite-oxidizing bacterium (NOB) isolated from the Equatorial
Pacific Ocean near the Galapagos Islands (WATSON and WATERBURY, 1971), was generously
provided by Frederica Valois (Woods Hole Oceanographic Institution). *N. mobilis* is a member
of the gamma-*Proteobacteria* that grows optimally at 25-30 °C and at pH 7.5-8.0 in 70-100%
seawater (WATSON and WATERBURY, 1971). It is an obligate chemolithoautotroph, which fixes

151 carbon dioxide into biomass using NO_2^- as an electron source (WATSON and WATERBURY, 1971). 152 In addition to oxidizing NO_2^- to NO_3^- , most nitrite oxidizing bacteria can assimilate NO_3^- , NO_2^- , 153 and NH_4^+ during growth (STARKENBURG et al., 2006). The physiology of *N. mobilis* has not been 154 studied exhaustively, but its draft genome contains genes required to transport nitrite, assimilate 155 nitrite (*nirB*), and reduce nitrite to nitric oxide (*nirK*). No genes have been found in this organism 156 for ammonium oxidation to nitrite.

157 In this study, N. mobilis was maintained in marine NOB medium containing 75% seawater 158 and 25% distilled water amended with 400 µM MgSO4, 30 µM CaCl₂, 5 µM K₂HPO4, 2.3 µM 159 Fe(III)-EDTA ("Geigy iron"), 0.1 µM Na₂MoO₄, 0.25 µM MnCl₂, 0.002 µM CoCl₂, and 0.08 160 µM ZnSO₄ (WATSON and WATERBURY, 1971). Filter-sterilized NaNO₂ was added to a final 161 concentration of 10 mM after autoclaving the base medium. One-liter cultures were inoculated 162 by transfer of 100 mL cultures (that had exhausted their NO₂) into 900 mL NOB medium. The 163 one-liter maintenance cultures were then grown for approximately one month before transfer or 164 experimentation.

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166 **2.2 Nitrite Oxidation Experiments**

To collect *N. mobilis* cells for isotopic fractionation experiments, 300 mL of maintenance culture (cell density ~ 10^8 cells/mL) were harvested by centrifugation at 7600 g for 20 minutes in six 50 mL sterile centrifuge tubes. The cell pellets were each washed with 2 mL sterile marine NOB medium containing 100 μ M NaNO₂. Washed cells were pooled, then re-aliquotted into sterile 2 mL centrifuge tubes and collected by centrifugation at 10,000 g. The supernatant was pipetted away and the cells were resuspended and pooled in 8 mL sterile marine NOB medium with 100 μ M NaNO₂.

174 Results from three independent experiments are reported here. The first experiment (hereafter 175 referred to as 'Experiment 1') was initiated by adding 0.5 mL of washed cells to flask "A" and 1.0 mL to flask "B", each containing 100 mL marine NOB medium with 100 µM NaNO₂ at pH 176 8.2. A third flask received no cells and was incubated with $100 \,\mu M \, \text{NO}_2^-$ as an uninoculated 177 178 control. In the second experiment (hereafter referred to as 'Experiment 2'), nine 250 mL flasks 179 were prepared, each with 100 mL of NOB medium at pH 7.8. Experiment 2 was initiated by 180 adding 0.5 mL of washed cells to triplicate "A" flasks and 1.0 mL to triplicate "B" flasks. The 181 three final flasks received no cells (uninoculated controls). In the third experiment (hereafter 182 referred to as "Experiment 3'), three 250 mL flasks were prepared, each with 100 mL of medium 183 with 100 µM NaNO₂ at pH 8.8. Experiment 3 was initiated by adding 0.5 mL of washed cell 184 suspension to flask "A" and 1.0 mL to flask "B". The third flask received no cells and was 185 incubated as an uninoculated control. In each experiment, final cell densities were estimated to be on the order of $10^6 - 10^7$ cells/mL, with half as many cells in "A" flasks compared with "B" 186 187 flasks.

Flasks were incubated at room temperature with shaking (Experiments 1 and 2) or without shaking (Experiment 3). Time course samples were collected immediately after inoculation and periodically thereafter until approximately all of the NO_2^- had been converted to NO_3^- (2-6 days, depending on the experiment). At each time point, 10-15 mL were withdrawn for nitrite and nitrate concentration analyses, as well as for nitrite and nitrate isotopic analyses; these samples were passed through a 0.2 μ M pore size membrane filter and stored frozen.

195 2.3 Chemical Analyses

196 NO_2^{-1} concentrations were analyzed in duplicate using the Greiss-Islovay colorimetric 197 reaction (STRICKLAND and PARSONS, 1972). Combined NO₂⁺+NO₃⁻ concentrations were 198 determined from the m/z = 44 peak area obtained by isotopic analyses using the denitrifier method (SIGMAN et al., 2001; CASCIOTTI et al., 2002). A two-point calibration for NO₂⁺+NO₃⁻ 199 200 concentration was derived from blank (0 µM added) and 20 µM standards, analyzed after every 201 nine samples. Most samples were analyzed in duplicate, and uncertainties in $[NO_2 + NO_3]$ were 202 derived from replicate analyses where possible. NO₃⁻ concentrations were calculated by 203 difference between $[NO_2^++NO_3^-]$ and $[NO_2^-]$, and standard deviations for $[NO_3^-]$ measurements 204 were calculated by error propagation from standard deviations of $[NO_2^-]$ and $[NO_2^-+NO_3^-]$ 205 measurements.

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207 2.4 Isotopic Analyses and Corrections

Nitrite $\delta^{15}N$ values ($\delta^{15}N_{NO2} = ([^{15}N/^{14}N]_{NO2}/[^{15}N/^{14}N]_{AIR} - 1) \times 1000$) from the time course 208 209 samples were analyzed by reaction with acetic acid-buffered sodium azide to form nitrous oxide 210 (N₂O) via the "azide method" (MCILVIN and ALTABET, 2005). Isotope ratios of the produced 211 N₂O were measured by continuous flow isotope ratio mass spectrometry (CF-IRMS) using an 212 on-line purge and trap system for extraction and purification of the N₂O analyte (see CASCIOTTI 213 et al., 2007). Analyses were calibrated to the AIR-N₂ reference scale using sodium nitrite salts 214 with calibrated $\delta^{15}N_{NO2}$ values: N-23 (+3.7‰), N-7373 (-79.6‰), and N-10219 (+2.5‰) (CASCIOTTI et al., 2007). Solutions of these standard nitrite materials were reacted and analyzed 215 216 in parallel to nitrite samples, with an aliquot of each standard analyzed initially and after every 217 nine experimental samples. Sample and standard quantities were matched to achieve 10 nmoles

of either sample or standard NO₂⁻ converted to 10 nmoles of N₂O by reaction with sodium azide. In this technique, one nitrogen atom in the N₂O arises from NO₂⁻ and one from azide (MCILVIN and ALTABET, 2005). Analyses were conducted only on samples with NO₂⁻ concentrations greater than 1 μ M. Each sample was analyzed in duplicate and the δ^{15} N values are reported as the average \pm one standard deviation.

Nitrate plus nitrite δ^{15} N values (δ^{15} N_{NO2+NO3} = ($[^{15}$ N/ 14 N]_{NO2+NO3}/ $[^{15}$ N/ 14 N]_{AIR} - 1)×1000) 223 were analyzed by bacterial conversion of nitrogen oxides (NO2⁻ and NO3⁻) to N2O via the 224 225 "denitrifier method" (CASCIOTTI et al., 2002; SIGMAN et al., 2001). The isotopes of N₂O were 226 analyzed via CF-IRMS, as described above. Analyses were calibrated to the AIR-N₂ reference scale using potassium nitrate reference materials USGS32 ($\delta^{15}N = +180.0\%$), USGS34 ($\delta^{15}N = -$ 227 1.8‰), and USGS35 ($\delta^{15}N = +2.7\%$) (BOHLKE et al., 2003). Standards were reacted and 228 229 analyzed in parallel to samples, with an aliquot of each standard analyzed initially and after 230 every nine experimental samples. Sample and standard quantities were matched to achieve 20 nmoles of sample $(NO_2^- + NO_3^-)$ or standard (NO_3^-) converted to 10 nmoles N₂O by the 231 232 denitrifier method. Most samples were analyzed in duplicate and are reported as average values 233 ± one standard deviation.

The majority of the time course samples contained a mixture of NO₂⁻ and NO₃⁻. Therefore, in order to estimate nitrate $\delta^{15}N(\delta^{15}N_{NO3})$ values, the procedure previously described (CASCIOTTI and MCILVIN, 2007) was followed for mass and isotope balance of the mixture components: $\delta^{15}N_{NO3} = (\delta^{15}N_{NO2+NO3} \times [NO_2^- + NO_3^-] - \delta^{15}N_{NO2} \times [NO_2^-]) \div [NO_3^-]$. The approximation required by this equation that ¹⁵N is a minor constituent does not introduce significant error in $\delta^{15}N_{NO3}$ estimation for the samples described here. Standard deviations for $\delta^{15}N_{NO3}$, $\delta^{15}N_{NO2+NO3}$, and

241 concentration measurements. For samples with a single $\delta^{15}N_{NO2+NO3}$ analysis, the standard 242 deviation was assumed to be 0.3‰ (SIGMAN et al., 2001).

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244 **2.5 Fractionation Factor Estimates**

The kinetic fractionation factor for nitrite oxidation $({}^{15}\alpha_{kNXR})$ was calculated from the 245 $\delta^{15}N_{NO2}$ data using a closed-system Rayleigh distillation model (SCOTT et al., 2004): ln($^{15}R_{NO2}$) 246 = $(1^{15}\alpha_{k,NXR} - 1) \cdot \ln(C) + \ln({}^{15}R_{NO2.0}/C_0^{(1/\alpha,kNXR-1)})$, where ${}^{15}R_{NO2}$ is the ${}^{15}N/{}^{14}N$ ratio of NO₂⁻ 247 (unitless) and ${}^{15}R_{NO2.0}$ is its initial value. C is the NO₂⁻ concentration (μ M) and C₀ is its initial 248 value. ${}^{15}\alpha_{k NXR}$ (unitless) is the ratio of the first order rate constants for ${}^{14}NO_2^{-1}$ and ${}^{15}NO_2^{-1}$ 249 oxidation, respectively $({}^{14}k_{NXR}/{}^{15}k_{NXR})$. Fractionation factors calculated in this way are the 250 251 reciprocal of those discussed by Mariotti et al (MARIOTTI et al., 1981). Dummy variables were used to obtain the average and standard deviation of ${}^{15}\alpha_{k,NXR}$ for all 252 experiments (SCOTT et al., 2004). This procedure utilizes the equation $y = \beta x + \gamma_1 + \gamma_2 D_2 + \gamma_3$ 253 D₃, where $y = \ln({}^{15}R_{NO2})$, $x = \ln(C)$, and $\beta = (1/{}^{15}\alpha_{k NXR} - 1)$. First, experimental data were 254 255 combined from different flasks to obtain the y-intercept for experiment 1 (γ_1) and the y-intercept 256 offset from γ_1 for experiments 2 (γ_2) and 3 (γ_3). Cell density did not appreciably change the observed kinetic isotope effect, so "A" and "B" flasks from each experiment were treated as 257 replicates for this purpose. Next, the β value was determined from each x-y pair, and the ${}^{15}\alpha_{k,NXR}$ 258 value was calculated from each β value. Finally, the average and standard deviation for ${}^{15}\alpha_{k,NXR}$ 259 260 were calculated from the entire dataset.

262	2.6 Ab Ir	nitio Calcu	lations of N	Molecular V	Vibration F	requencies

Molecular vibration frequencies for two isotopic forms (containing either ¹⁵N or ¹⁴N) of NO₂ 263 , HNO₂, and NO₃⁻ were calculated using the Gaussian interface through the WebMO (Version 264 265 8.0.011e) guest server at buchner.chem.hope.edu. The geometry of each molecule was first 266 optimized in either 'water' or 'no solvent', depending on the desired phase of the vibration 267 calculations. The optimized molecules were then used to calculate the molecular vibration 268 frequencies using B3LYP density function theory with the 'routine' basis set (6-31G(d)) (ANBAR 269 et al., 2005) in both 'no solvent' and in aqueous solution. The geometry optimization and 270 vibration calculations were performed three times each, and the frequencies obtained varied by 271 less than 0.01%. 272

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3.1 Time Course Experiments

3. RESULTS

275 In each independent experiment, NO_2^- was oxidized to NO_3^- by pure cultures of N. mobilis 276 over a course of 2-6 days (Figure 1A, C, E). NO₂⁻ removal was nearly complete in each 277 inoculated flask. Each experiment was conducted using a two-fold range of cell densities (approximately 2-4 x 10^7 cells mL⁻¹), depending on the volume of concentrated cell suspension 278 279 added. The net rate of nitrite oxidation apparently increased with cell density (Figure 1A, C, E). 280 For the most part, mass balance was maintained between NO₂⁻ and NO₃⁻ pools as NO₂⁻ was 281 consumed. Concentrations of NO₂⁺+NO₃⁻ in experimental flasks showed up to 10-20% variation at some time points, but as discussed below, there was no change in $\delta^{15}N_{NO2+NO3}$ corresponding 282 283 to these apparent changes in $[NO_2^++NO_3^-]$. Control flasks also showed similar variability in 284 $[NO_2^++NO_3^-]$, suggesting that the variability was most likely not biological in nature. At this

time, the most likely explanation is error in concentration measurement at particular time pointsin the time course.

In every case, the nitrite became progressively depleted in ${}^{15}N$, decreasing from initial $\delta^{15}N$ 287 288 values of +0.7‰ to less than -50‰ as it was oxidized to NO₃⁻ (Figure 1B, D, F). In contrast, uninoculated control flasks showed no change $\delta^{15}N_{NO2}$ values over the time course. Control 289 flasks also showed no change in the δ^{18} O value of NO₂⁻+NO₃⁻ (data not shown), a measurement 290 291 that is very sensitive to changes in the fractions of nitrate and nitrite in the mixture (CASCIOTTI et al., 2007). Finally, within the standard error of the measurement (~0.3‰) $\delta^{15}N_{NO2+NO3}$ did not 292 293 change over the time course in any flask, whether nitrite oxidation occurred (experimental flasks) or not (control flasks) (Figure 1 B, D, F). The consistency of $\delta^{15}N_{NO2+NO3}$ over the time course is 294 a strong constraint on mass balance between NO₂⁻ and NO₃⁻ because any potential loss of N from 295 the NO₂⁻ and NO₃⁻ pools would have needed to possess a constant $\delta^{15}N$ value of +1‰, even as 296 $\delta^{15}N_{NO2}$ decreased from +1% to -52%. Isotopically, it would be difficult to account for los of 297 NO_2 to other pools of N that were not measured (cell biomass, NO, NH_4^+). 298 Mass balance allows the calculation of the δ^{15} N of accumulated NO₃⁻ over the time course 299 300 (CASCIOTTI and MCILVIN, 2007). As expected from a closed-system isotope mass balance, $\delta^{15}N_{NO3}$ was highest initially, then decreased over the time course to a value within error of the 301 initial $\delta^{15}N_{NO2}$ (not shown). The maintenance of mass (Figure 1A, C, E) and isotope (Figure 1B, 302 303 D, F) balance between nitrate and nitrite in experimental flasks and the lack of nitrite oxidation 304 to nitrate in control flasks points towards biological nitrite oxidation as the process responsible 305 for NO_2^- consumption in these experiments.

307 3.2 Rayleigh Fractionation

Analysis of the observed trends in $\delta^{15}N_{NO2}$ using a closed system Rayleigh model yields estimates for ${}^{15}\alpha_{k,NXR}$ of 0.9873 ± 0.0010, 0.9867 ± 0.0018, and 0.9882 ± 0.0005 for the three experiments, respectively (Figure 2). Averaging over all experiments yields an estimate of 0.9872 ± 0.0015 for ${}^{15}\alpha_{k,NXR}$. In all three experiments, the kinetic fractionation factor appears to be inverse. In addition, the three experiments were all very similar to each other, with individual estimates for ${}^{15}\alpha_{k,NXR}$ within two standard deviations of the overall ${}^{15}\alpha_{k,NXR}$.

314 As observed in Figure 2, the fractionation factor appeared to increase slightly (become less 315 inverse) after approximately 80% of the nitrite has been oxidized, as indicated by the deviation from the ${}^{15}\alpha_{k,NXR} = 0.9872$ line at ln([NO₂⁻]) values below 2 (corresponding to [NO₂⁻] less than 316 317 \sim 7 μ M) (Figure 2). Both experiments 1 and 2 showed this change of slope. Inclusion of these points in the analysis may be one reason for the slight variation between ${}^{15}\alpha_{k NXR}$ in experiments 318 1, 2, and 3. The apparent change in ${}^{15}\alpha_{k,NXR}$ at later time points in the experiment may result 319 320 from departures of the system from the basic assumption of a unidirectional closed system 321 inherent in the Rayleigh distillation equation. Many systems show a tendency towards smaller 322 isotope effects at high extents of substrate consumption as a result of diffusive limitation or 323 active transport of substrate (FARQUHAR et al., 1982; GRANGER et al., 2008; GRANGER et al., 324 2006; NEEDOBA et al., 2004; O'LEARY, 1981). Given that the N. mobilis experiments were 325 conducted with whole cells, these factors may play a role in the expressed isotope effect, which I 326 will return to below. However, both diffusion of NO_2^- and uptake by the cell are likely to have 327 small, normal kinetic isotope effects (GRANGER et al., 2008; WASER et al., 1998b), and are likely 328 to be of secondary importance in the current analysis.

4. DISCUSSION

331 Given that true inverse kinetic isotope effects are rare, two alternative explanations were considered for the observed results. First, one must consider whether nitrite with low δ^{15} N could 332 333 have been produced in the experiments over time, thus mimicking an inverse kinetic isotope 334 effect. The only mechanism possible for nitrite production in *N. mobilis* is through reduction of 335 nitrate by reversible reaction of the nitrite oxidoreductase enzyme. The membrane-bound nitrite 336 oxidoreductase enzymes isolated from Nitrobacter agilis and Nitrobacter hamburgensis have 337 been shown to be reversible under low pH (TANAKA et al., 1983) and reducing conditions 338 (MEINCKE et al., 1992; SUNDERMEYER-KLINGER et al., 1984). Although oxygen concentrations 339 are not expected to have been limiting in the current N. mobilis experiments, enzyme-level 340 reversibility is one mechanism that could produce an apparent inverse kinetic isotope effect, if 341 the reverse reaction (nitrate reduction) is more highly fractionating than the forward reaction 342 (nitrite oxidation). Therefore, the possibility for enzymatic back-reaction of NO_3^- to NO_2^- was 343 carefully examined in section 4.1.2.

344 The second scenario under which a normal enzyme-level kinetic fractionation factor may 345 appear to be inverse involves equilibrium between the true substrate for the reaction and a larger 346 pool of non-reactive substrate. This is termed a "pre-equilibrium" scenario because the 347 equilibrium reaction would occur prior to the enzymatic reaction, rather than at the enzyme level. 348 For example, a pH-dependent pre-equilibrium between H₂S and HS⁻ was used to explain the 349 small inverse kinetic isotope effect observed in anaerobic sulfide oxidation (FRY et al., 1984). 350 The pH-dependent isotope exchange between NO_2^- and HNO_2 is examined here in section 4.1.1. 351 Rejecting these two scenarios, in which the apparent inverse fractionation factor would be driven 352 by equilibrium or reversibility, it is shown that transition state theory supports the observation of

an inverse kinetic fractionation factor for nitrite oxidoreductase at the enzyme level (Section4.1.3).

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4.1 Potential Explanations for the Inverse Isotope Effect

357 *4.1.1 Pre-equilibrium*

358 Apparent inverse fractionation factors have been observed in systems where a fractionating 359 equilibrium reaction precedes a unidirectional enzymatic reaction. An example of this is in anaerobic oxidation of sulfide to elemental sulfur (S^0) by the purple photosynthetic bacterium 360 361 Chromatium vinosum (FRY et al., 1984). An inverse isotope effect was observed, whereby the sulfide undergoing oxidation became depleted ³⁴S over time. The S⁰ produced by this reaction 362 was initially enriched in ³⁴S, but then also became depleted over time, approaching the δ^{34} S of 363 364 the initial substrate. However, in the sulfide oxidation reaction, the apparent inverse kinetic 365 isotope effect was explained by co-occurring equilibrium and kinetic isotope effects, rather than 366 a true inverse kinetic isotope effect (FRY et al., 1984). The equilibrium isotope effect between hydrogen sulfide (H_2S) and bisulfide (HS^-) caused the substrate for the reaction, H_2S , to be 367 enriched in ³⁴S by approximately 6‰ relative to the HS⁻ (FRY et al., 1984). Therefore, when H₂S 368 reacted with a small *normal* kinetic isotope effect the overall process produced S⁰ that was 369 enriched in ³⁴S relative to the total reduced sulfur pool, yielding an apparently inverse kinetic 370 371 isotope effect.

In the case of nitrite oxidation to nitrate, it is often assumed that the substrate for the reaction is NO₂⁻, although few experiments have been specifically designed to distinguish between HNO₂ and NO₂⁻ as substrates (BOON and LAUDELOUT, 1962). Given this uncertainty, the possibility that HNO₂ could have a significantly different δ^{15} N from the total N(III) pool (NO₂⁻ + HNO₂) must

376 be considered. Nitrous acid (HNO₂) has a pKa of 2.8 (RIORDAN et al., 2005), and at pH 8 more 377 than 99.999% of the total N(III) occurs in the form of NO_2^- , and less than 0.001% of the N(III) 378 occurs in the form of HNO₂. In order for the pH-dependent pre-equilibrium to explain the 379 observed inverse kinetic isotope effect, 1) the minor pool (in this case, HNO_2) must be the substrate for the reaction, 2) HNO₂ must be enriched in 15 N relative to NO₂, and 3) the 380 381 magnitude of the equilibrium isotope effect must be large enough to overcome a normal enzyme-382 level kinetic isotope effect for the oxidation to nitrate. The scenarios that would allow the 383 observed inverse isotope effect to be consistent with a normal enzyme-level kinetic isotope 384 effect, were evaluated through *ab initio* calculation of the equilibrium isotope effect for N 385 isotope exchange between HNO₂ and NO₂, in combination with six different Rayleigh model 386 simulations (see below for additional details). .

387 The nitrogen isotope exchange between HNO_2 and NO_2^- is represented by equation (1): 388

$$389 \qquad H^{14}NO_2 + {}^{15}NO_2^{-} \longleftrightarrow^{15}\alpha_{eq,HNO_2/NO_2} \to H^{15}NO_2 + {}^{14}NO_2^{-}$$
(1)

390

where ${}^{15}\alpha_{eq,HNO2/NO2}$ is the fractionation factor for this isotope exchange equilibrium. As discussed earlier, ${}^{15}\alpha_{eq,HNO2/NO2}$ defines the ratio of isotope ratios of HNO₂ and NO₂⁻ at equilibrium (${}^{15}R_{HNO2}/{}^{15}R_{NO2})_{eq}$. This equilibrium fractionation factor can be calculated (with some approximations that are suitable for isotopes other than hydrogen) using the reduced partition functions (Q) if the vibration frequencies of the ${}^{14}N$ and ${}^{15}N$ -containing molecules of both species are known (UREY, 1947). Equation 2 illustrates how ${}^{15}\alpha_{eq,HNO2/NO2}$ is related to the reduced partition function ratios for the two species.

$$399 \qquad {}^{15}\alpha_{eq,HNO2/NO2} = \frac{{}^{15}Q_{HNO2}}{{}^{15}Q_{NO2}}{}^{14}Q_{HNO2}} \tag{2}$$

401 Equation 3 shows how the reduced partition function ratio for either HNO_2 or NO_2^- is calculated 402 from the normal mode molecular vibration frequencies of ¹⁵N and ¹⁴N-containing molecules. 403

$$404 \qquad \frac{Q_{15}}{Q_{14}} = \frac{\sigma_{14}}{\sigma_{15}} \prod_{i} \frac{u_{15i}}{u_{14i}} \cdot \frac{e^{-u_{15i}/2}}{e^{-u_{14i}/2}} \cdot \frac{1 - e^{-u_{14i}}}{1 - e^{-u_{15i}}}$$
(3)

405

406 In equation 3, Q_{14} and Q_{15} are the reduced partition function ratios relative to the separated atoms, for ¹⁴N and ¹⁵N-containing molecules, respectively. σ_{14} and σ_{15} are the symmetry 407 numbers for the two isotopically substituted species. In addition, $u_i = hc\omega_i/kT$, where ω_i 408 corresponds to each of the ith normal frequencies (bent polyatomic molecules with N atoms have 409 410 3N-6 normal frequencies), h is Plank's constant, k is Boltzman's constant, c is the speed of light, 411 and T is the absolute temperature in K (here, T = 298 K). Therefore, in order to obtain a firstorder estimate for ${}^{15}\alpha_{eq,HNO2/NO2}$, we need only the molecular vibration frequencies for H¹⁴NO₂, 412 H¹⁵NO₂, ¹⁴NO₂, and ¹⁵NO₂. For additional discussion of the assumptions and simplifications 413 414 involved in using equations 2 and 3, please see Bigeleisen and Mayer (1947) and Urey (1947). 415 A collection of observed (BEGUN and FLETCHER, 1960; DEELEY and MILLS, 1985) and 416 calculated (*this study*) vibration frequencies for NO₂⁻ and HNO₂ are given in Table 2. Aqueous 417 vibration frequencies were available for NO₂⁻ and NO₃⁻ (BEGUN and FLETCHER, 1960), but not for HNO₂. For HNO₂ the only available vibration frequencies were in gas phase (DEELEY and 418 419 MILLS, 1985). Given that the molecular vibrations could be quite different in aqueous vs. gas 420 phase (RICHET et al., 1977), vibrational frequencies were needed for HNO₂ in the aqueous phase. 421 To check the accuracy of the calculation, the aqueous phase vibration frequencies for NO_2^- and 422 gas phase vibration frequencies for HNO_2 were also calculated to compare with observed 423 frequencies (Table 2). Finally, to test the sensitivity to solvation the vibration frequencies for 424 NO_2^- were also calculated for the 'no solvent' case.

The calculated frequencies for NO₂⁻ and HNO₂ were similar to observed frequencies, with 425 differences ranging from -6.8 to +1.5% for ¹⁴N-containing species and -6.7 to +1.7% for ¹⁵N-426 427 containing species (Table 2). The frequency shifts upon isotopic substitution were also similar. On average, a 1.13% vibrational frequency shift upon substitution with ¹⁵N was obtained in the 428 429 calculated frequencies, compared with a 1.16% vibrational shift in the observed frequencies. 430 Finally, the normal modes that appeared insensitive to isotopic substitution in the modeled 431 frequencies also showed little isotopic shift in the observed frequencies. The comparison 432 between modeled and observed frequencies gives some confidence in the calculated results, 433 although both observed and calculated frequencies could be inaccurate for different reasons. The 434 differences between calculated aqueous and gas phase frequencies were less than 2.5% in all 435 cases.

436 The modeled frequencies were used to calculate the equilibrium fractionation factor for 437 nitrogen exchange between NO₂⁻ and HNO₂ using equations 2 and 3. Little difference was obtained for ${}^{15}\alpha_{eq,HNO2/NO2}$ in aqueous (0.9975) versus gas (0.9972) phase (Table 3). In addition, 438 use of the observed frequencies (from mixed phases) yields an ${}^{15}\alpha_{eq,HNO2/NO2}$ value of 0.9970, 439 which is very similar to the ${}^{15}\alpha_{eq HNO2/NO2}$ values obtained from the *ab initio* frequencies. In all 440 cases, the calculated ${}^{15}\varepsilon_{eq,HNO2/NO2}$ values were small at room temperature (-3 to -2.5%), and 441 HNO_2 was predicted to be depleted in ¹⁵N relative to NO_2^- . Given these results for HNO_2/NO_2^- 442 equilibration, it is anticipated that HNO₂ would have been depleted in ¹⁵N relative to NO₂⁻ by 2.5 443

444 to 3.0‰ in the *N. mobilis* experiments (room temperature, pH 7.8 - 8.8). In contrast to the

445 H_2S/HS^2 equilibrium where protonation and deprotonation directly affect the bonding

446 environment of the sulfur atom (a primary isotope effect), protonation and deprotonation of NO₂⁻

447 does not occur at the central N atom. Therefore, it is not surprising that protonation and

448 deprotonation of the NO₂⁻/HNO₂ couple should exhibit a small isotope effect.

449 Given the small equilibrium isotope effect between HNO_2 and NO_2^- , it can be shown that

450 reaction of either NO_2^- or HNO_2 with a normal fractionation factor cannot explain an apparent

451 inverse kinetic fractionation factor of the magnitude that we observe. This is illustrated in detail

452 through Rayleigh simulations using various estimates for the kinetic (normal vs. inverse) and

453 equilibrium isotope effects, and using either HNO_2 or NO_2^- as substrate for the reaction

454 (*Electronic Annex* 1). The pre-equilibrium between HNO_2 and NO_2^- can therefore be rejected as 455 a mechanism contributing to the apparent inverse kinetic fractionation factor.

456

457 4.1.2. Enzyme-Level Reversibility

Enzyme reversibility could lead to an apparent inverse isotope effect, even if the forward and reverse reactions are normal, if 1) the reverse reaction has a larger isotope effect than the forward reaction, and 2) there is sufficient mass flux in the reverse reaction to lead to isotopic depletion of the primary substrate pool. The flux required to obtain an apparent inverse isotope effect will vary inversely with the magnitude of the isotope effect for the reverse reaction, in this case nitrate reduction to nitrite.

464 Nitrite oxidoreductase is known to be reversible (SUNDERMEYER-KLINGER et al., 1984;

465 TANAKA et al., 1983). For example, some studies have shown growth of *Nitrobacter* under

466 anaerobic conditions using nitrate as a terminal electron acceptor (AHLERS et al., 1990; BOCK et

467 al., 1988; FREITAG and BOCK, 1990; FREITAG et al., 1987), and other studies have shown that 468 Nitrobacter can assimilate nitrate-N (KUMAR and NICHOLAS, 1982). Both processes require a 469 mechanism for reducing nitrate to nitrite; however, nitrate reductase has not been found in any of 470 the nitrite oxidizer genomes studied thus far, including Nitrococcus mobilis (STARKENBURG et 471 al., 2008). Instead, it is believed that nitrite oxidoreductase acts as the nitrate reductase under 472 certain growth conditions. This theory is based on the biochemical similarity (YAMANAKA, 1996; 473 YAMANAKA and FUKUMORI, 1988) and evolutionarily relatedness of the genes for nitrite 474 oxidoreductase to the *narGH* family of dissimilatory nitrate reductases (STARKENBURG et al., 475 2008), as well as observed reversibility of partially purified nitrite oxidoreductase from 476 Nitrobacter species (TANAKA et al., 1983; YAMANAKA and FUKUMORI, 1988). The biochemistry 477 of nitrite oxidoreductase has not been studied in Nitrococcus species directly, but the nitrite 478 oxidizing systems appear to be similar between the two genera (BARTOSCH et al., 1999; SPIECK 479 and BOCK, 2001), and some reversibility of this enzyme might be expected under some 480 circumstances.

Here the role of enzyme reversibility in the observed inverse kinetic isotope effect for nitrite oxidation was evaluated. Two aspects of this question are discussed: first, whether equilibrium calculations for the nitrite/nitrate redox system support a normal kinetic isotope effect for the forward reaction (nitrite oxidation to nitrate) and second, whether significant back reaction (nitrate reduction to nitrite) can be detected during experiments 1-3. To address the first component, the equilibrium fractionation factor for isotope exchange between nitrate and nitrite (equation 4) was first computed.

488

$$489 \qquad {}^{14}NO_2^- + {}^{15}NO_3^- \xleftarrow{}^{15}\alpha_{eq,NO2/NO3}} \longrightarrow {}^{15}NO_2^- + {}^{14}NO_3^- \tag{4}$$

The equilibrium fractionation factor for the isotope exchange reaction $({}^{15}\alpha_{eq,NO2/NO3})$ can be 491 492 calculated using the same principles as applied above to calculate the isotope effect for nitrogen 493 isotope exchange between NO_2^- and HNO_2 (equations 2 and 3). Table 2 contains the relevant molecular vibration frequencies for NO₂⁻ and NO₃⁻ for use in equation 3. Using the calculated 494 molecular vibrations for NO₂⁻ and NO₃⁻ in aqueous solution, it is estimated that ${}^{15}\alpha_{eq,NO2/NO3}$ is 495 496 approximately 0.945 in aqueous solution (Table 3). Furthermore, using experimentally-derived molecular vibrations for aqueous NO₂⁻ and NO₃⁻ ions yields ${}^{15}\alpha_{eq,NO2/NO3}$ values of 0.939 (BEGUN 497 and FLETCHER, 1960) or 0.922 (SPINDEL, 1954), which compare favorably to the results 498 499 presented in Table 3.

The equilibrium fractionation factor for this isotope exchange reaction can be represented by the ratio of kinetic fractionation factors for forward and reverse reactions (O'LEARY, 1981). In this case, the forward and reverse reactions are nitrite oxidation and nitrate reduction, respectively. The kinetic fractionation factor for nitrite oxidation (${}^{15}\alpha_{k,NXR}$) can therefore be estimated from the computed equilibrium fractionation factor (${}^{15}\alpha_{eq,NO2/NO3}$) and the kinetic fractionation factor for nitrate reduction (${}^{15}\alpha_{k,NAR}$):

506

507
$${}^{15}\alpha_{eq,NO2/NO3} = \left[\frac{{}^{15}NO_2^{-}/{}^{14}NO_2^{-}}{{}^{15}NO_3^{-}/{}^{14}NO_3^{-}}\right]_{eq} = \frac{{}^{15}\alpha_{k,NXR}}{{}^{15}\alpha_{k,NAR}}$$
 (5)

508

Each estimate for ${}^{15}\alpha_{eq,NO2/NO3}$ was less than unity (Table 3), suggesting that ${}^{15}\alpha_{k,NXR}$ is lower than ${}^{15}\alpha_{k,NAR}$ and that ${}^{15}N$ accumulates preferentially in NO₃⁻ at equilibrium. Assuming that the fractionation factor for nitrate reduction by the enzyme nitrite oxidoreductase (operating in 512 reverse) is between 1.020 and 1.030, as has been observed for nitrate reductases (Table 1; 513 (BARFORD et al., 1999; DELWICHE and STEYN, 1970; GRANGER et al., 2006), the value of $^{15}\alpha_{k \text{ NXR}}$ estimated from equation 5 would vary between 0.964 and 0.973. Because of 514 515 simplifications made in calculating the equilibrium isotope effects, they should be seen as 516 approximations. However, rather than indicating a role for the reverse reaction in driving an 517 artificial inverse kinetic isotope effect, these calculations suggest that the isotope effect for nitrite oxidation may actually be inverse. In order for the calculated ${}^{15}\alpha_{eq,NO2/NO3}$ values to be consistent 518 with a normal ${}^{15}\alpha_{k,NXR}$, ${}^{15}\alpha_{k,NAR}$ would have to be greater than 1.058, which is unlikely though 519 520 not impossible (TCHERKEZ and FARQUHAR, 2006).

It is interesting that the observed ${}^{15}\alpha_{k NXR}$ value of 0.9872 is actually closer to 1 ('more 521 522 normal') than expected from the results of these calculations. This may reflect partial rate 523 determination by steps prior to nitrite oxidation to nitrate, such as transport into and out of the 524 cell, diffusion to the enzyme active site, or fractionation associated with enzyme/substrate 525 association and dissociation. The expression of enzyme-level isotope effects in the external 526 medium does indeed require exchange of substrates and products between the interior and 527 exterior of the cell (FARQUHAR et al., 1982; GRANGER et al., 2008; NEEDOBA et al., 2004; 528 O'LEARY, 1981). If diffusion or uptake of nitrite is partially rate-limiting, these processes could 529 invalidate the use of a simple unidirectional Rayleigh model to fit the N. mobilis data. However, 530 in this case, the observed kinetic isotope effect should provide a minimum estimate of the inverse 531 kinetic isotope effect because the processes of diffusion and transport would be expected to 532 occur with normal isotope effects. There are a series of assumptions that should be tested by 533 experiment, such as the isotope effect for nitrate reduction by nitrite oxidoreductase and the 534 equilibrium fractionation between nitrite and nitrate. However, current calculations suggest that

in addition to an inverse kinetic fractionation in the oxidation reaction, any reversibility wouldtend to further lower the observed kinetic isotope effect.

537 The data from experiments 1-3 were also compared with a simple reaction reversibility 538 model (Electronic Annex 2) to determine whether enzyme reversibility was detectable in the culture conditions employed here. Figure 3 shows the simulated $\delta^{15}N_{NO2}$ and $\delta^{15}N_{NO3}$ values for 539 540 four scenarios with different rates of nitrate reduction, relative to nitrite oxidation. This ratio of 541 forward (nitrite oxidation) and reverse (nitrate reduction) reactions is represented by the parameter $x = {}^{14}k_{NXR}/{}^{14}k_{NAR}$, which is set to values of 0, 0.2, 0.5, and 1.0 in the different 542 543 simulations. This formulation assumes both nitrite oxidation and nitrate reduction are first order 544 with respect to their substrates.

545 As the amount of reaction reversibility (x) increases from 0 to 1, significant differences are seen in the relationship of $\delta^{15}N_{NO2}$ and $\delta^{15}N_{NO3}$ to the apparent extent of reaction (as determined 546 547 from $[NO_2^-]$ (Figure 3). With any significant amount of nitrate reduction, the forward reaction 548 fails to go to completion and there is always some NO_2^- that remains in the simulation (Figure EA2.1). This is also indicated by f values failing to reach 0 (Figure 3). Nitrate reduction also 549 leads to enhanced separation of ¹⁵N between NO₂⁻ and NO₃⁻ (Figure 3), and at equilibrium, the 550 remaining NO₂⁻ would be expected to have a δ^{15} N value that is approximately 50% lower than 551 $\delta^{15}N_{NO3}.$ Furthermore, at values of x greater than 0.2, the $\delta^{15}N_{NO2}$ and $\delta^{15}N_{NO3}$ values no longer 552 appear to follow a closed-system "Rayleigh" model. Although $\delta^{15}N_{NO2}$ decreases in all 553 simulations, the behavior of $\delta^{15}N_{NO3}$ changes dramatically. Rather than decreasing towards the 554 initial $\delta^{15}N_{NO2}$ value, $\delta^{15}N_{NO3}$ gradually increases to its equilibrium $\delta^{15}N$ value. This value is 555 556 based on the concentrations of NO_2^- and NO_3^- that remain in the system at equilibrium and the assumed value of ${}^{15}\alpha_{eq,NO2/NO3}$. 557

558	Comparison of these results to the <i>N. mobilis</i> data from experiments 1-3 suggests that little
559	reversibility occurred (Figure 3). The $\delta^{15}N_{NO2}$ and $\delta^{15}N_{NO3}$ observations fall between the x = 0
560	and $x = 0.2$, consistent with little or no reversibility in the experiments described here.
561	Furthermore, nitrite concentrations decreased to less than 1% of their initial value in all
562	experiments, suggesting that there was at most 1% reaction reversibility ($x = 0.01$). At this time,
563	both the current observations and calculations are consistent with little reversibility of nitrite
564	oxidoreductase under near-neutral pH conditions (HOLLOCHER, 1984; STRAAT and NASON, 1965;
565	TANAKA et al., 1983; YAMANAKA and FUKUMORI, 1988).
566	

. . .

567 4.1.3. Enzyme-Level Inverse Isotope Effect

568 Evaluations of both the pre-equilibrium and the enzyme reversibility scenarios imply that the 569 kinetic fractionation factor for nitrite oxidation is likely to be inverse at the enzyme level. 570 Enzyme-level inverse kinetic isotope effects have rarely, if ever, been demonstrated 571 unambiguously in biochemical systems, making the results of the current study both surprising 572 and novel. Several studies have reported apparent inverse kinetic isotope effects but these were 573 shown to result from either pH-dependent pre-equilibrium (FRY et al., 1984), reaction branching 574 (FRY et al., 1985; MCCREADY et al., 1975), or reaction reversibility (WEISS et al., 1988). In one 575 case, the mechanism resulting in the observed inverse kinetic isotope effect was attributed to association and dissociation kinetics of the enzyme and substrate (GAWLITA et al., 1995), while 576 577 in another case the mechanism was never identified (MCCREADY et al., 1975). However, the 578 current results suggest that an inverse isotope effect is plausible, and even predictable, for the simple N-O bond forming reaction involved in nitrite oxidation. This hypothesis is examined 579

further using transition state theory applied to the current understanding of the mechanism fornitrite oxidation.

The reaction mechanism for nitrite oxidation has been studied using oxygen isotope transfer between H₂O, NO₂⁻, and NO₃⁻ species, and the currently accepted reaction mechanism involves a relatively simple oxygen (O) atom addition to NO₂⁻ (Equation 6). Previous studies have suggested that the reaction proceeds by initial oxygenation of the enzyme's metal cofactor (M) from either NO₃⁻ or H₂O to form an oxo complex in the +2 oxidation state (M=O). This step is followed by NO₂⁻ binding to the enzyme-bound O, reduction of the metal cofactor to the +0 state, and dissociation of the produced NO₃⁻ (DISPIRITO and HOOPER, 1986; FRIEDMAN et al., 1986).

589

590
$$M = O + NO_2^{-} \longleftrightarrow \left[M - O - NO_2 \right] \longleftrightarrow M + NO_3^{-}$$
 (6)

591

592 Using this reaction mechanism as a model, one can predict the isotope effect for nitrite 593 oxidation using transition state theory (TST), which has often been used to describe the kinetics 594 and isotopic fractionation of chemical and biochemical reactions (CLELAND, 1987; NORTHROP, 595 1981; O'LEARY, 1977; THORNTON and THORNTON, 1978). Key assumptions in TST are that the 596 transition state activated complex is in thermal equilibrium with the reactants and that the rate of 597 reaction depends on the transmission coefficient and the concentration of molecules in the 598 activated complex (BIGELEISEN, 1949; BIGELEISEN, 1952; BUDDENBAUM and SHINER, 1977). 599 With these assumptions, the kinetic fractionation factor can be estimated with knowledge of the 600 vibrational frequencies for the substrate and transition state, using equation 7 (BIGELEISEN, 1949; 601 BIGELEISEN and WOLFSBERG, 1958; HUSKEY, 1991):.

602

603
$$\alpha_{k} = k_{1}/k_{2} = (v_{1L}'/v_{2L}) \left[1 + \sum_{i}^{3n-6} G(u_{i})\Delta u_{i} - \sum_{i}^{3n'-6} G(u_{i}')\Delta u_{i}' \right]$$
(7)

In equation 7, u_i is defined as in equation 2, $G(u_i) = [1/2 - (1/u_i) + 1/(e^{u_i} - 1)]$, and primed 605 606 variables refer to properties of the transition state ([M-O-NO₂] in equation 6). The term v'_{1L}/v'_{2L} 607 is the ratio of imaginary frequencies in the transition state along the reaction coordinate for light 608 and heavy molecules, respectively; this term is always greater than 1 (FRY, 1970; SHEPPARD et 609 al., 1954) and can be determined from the reduced masses of the isotopic atoms involved in the 610 bond being formed or broken (SHEPPARD et al., 1954). The term in square brackets, which is 611 based on the zero point energy difference between the substrate and transition state for the light 612 and heavy molecules (BIGELEISEN, 1952), can be greater than or less than 1, depending on the 613 change in the bonding environment at the isotopically substituted atom between substrate and 614 activated complex (SHEPPARD et al., 1954). If N is more strongly bonded in the substrates than in 615 the transition state (a typical bond-breaking reaction), the sum of $G(u_i)\Delta u_i$ for the substrates will 616 be greater than the sum of $G(u_i)\Delta u_i$ in the transition state and the term in square brackets (eq. 7) will be greater than 1, leading to a normal kinetic isotope effect. This is the case for most 617 618 nitrogen-oxide reduction reactions (nitrate reduction, nitrite reduction, nitrous oxide reduction), 619 which involve N-O bond breaking (BARFORD et al., 1999; BRYAN et al., 1983; GRANGER et al., 620 2006). If, however, N is more strongly bonded in the transition state than in the substrates (a 621 bond forming reaction), the sum of $G(u_i)\Delta u_i$ in the reactants will be less than the sum of 622 $G(u_i)\Delta u_i$ in the transition state, and the term in the square brackets will be less than one. This 623 can lead to an inverse kinetic isotope effect if the product of this term and the imaginary 624 frequency term (v'_{11}/v'_{2L}) leads to a ratio of rate constants (k_1/k_2) that is less than one. If,

625 however, the product of the imaginary frequency term and the zero point energy term is greater 626 than 1, a normal kinetic isotope effect would be expected, even for a bond-forming reaction. 627 Here, the kinetic isotope effect for nitrite oxidation was estimated from equation 7 assuming 628 that the transition state contains the newly formed N-O bond (FRIEDMAN et al., 1986) and that 629 the geometry and bonding environment of the N atom in this complex are similar to nitrate 630 (FRIEDMAN et al., 1986; SHEPPARD et al., 1954). Assuming N-O bonding in the transition state is 631 similar to that of the product (nitrate) may overestimate the bonding difference between the 632 substrate and the transition state. This will be considered in the discussion below. The value of 633 the imaginary frequency term was estimated to be 1.0182, from the square root of the ratio of reduced masses of the bond forming atoms, ¹⁵N-¹⁶O and ¹⁴N-¹⁶O (SHEPPARD et al., 1954). The 634 635 zero point energy terms were estimated from the molecular vibration frequencies of both 636 isotopically substituted forms of NO₂⁻ (assumed substrate) and NO₃⁻ (assumed geometry for N in the activated complex). Using both experimental and *ab initio* frequencies for NO₂⁻ and NO₃⁻ 637 (Table 2), the ZPE term was estimated to range from 0.933 to 0.941, depending on whether 638 639 observed or calculated frequencies in aqueous or gas phase are used. Multiplying by the 640 imaginary frequency term, the inherent kinetic isotope effect for nitrite oxidation to nitrate would therefore be predicted from TST (eqn. 7) to be on the order of 0.950-0.958. This ${}^{15}\alpha_{k NXR}$ value 641 642 estimated from TST is lower (more inverse) than was observed experimentally (Figure 2). It is also similar, but a bit lower than the ${}^{15}\alpha_{k,NXR}$ value predicted from equilibrium calculations 643 644 (0.964-0.973; section 4.1.2.). However both of these estimates are consistent with the expectation 645 of an inverse kinetic isotope effect for nitrite oxidation. 646 There are many reasons for the exact values of measured and calculated kinetic isotope

647 effects to differ. Uncertainties involved in estimating ${}^{15}\alpha_{k,NXR}$ from the equilibrium exchange

648	reaction for NO_2^- and NO_3^- were discussed above (section 4.1.2). Application of transition state
649	theory also involves inherent assumptions about the enzyme-bound transition state structure and
650	vibrational frequencies, as well as the potential energy surfaces for isotopically substituted
651	molecules (NORTHROP, 1981; THORNTON and THORNTON, 1978). Assuming that the transition
652	state bonding and vibration frequencies are similar to NO ₃ ⁻ is likely to exaggerate the difference
653	in molecular properties between NO_2^- and the transition state. If the transition state properties are
654	more intermediate between NO_2^- and NO_3^- , the fractionation factor may indeed be closer to 1
655	(less inverse). Because of this uncertainty, it may not be surprising that the observed
656	fractionation factor is less inverse than the theoretical value calculated from TST.
657	Even if the estimates of ${}^{15}\alpha_{k,NXR}$ based on TST provide an accurate estimate for the inherent
658	kinetic isotope effect for nitrite oxidation, it may differ from the isotope effect observed
659	experimentally because of the influence of cellular processes and enzymatic steps prior to the
660	oxidation of nitrite (NORTHROP, 1981; O'LEARY, 1981; TCHERKEZ and FARQUHAR, 2006).
661	Cellular processes that could cause some deviation of the observed isotope effect relative to the
662	predicted (inherent) value cannot currently be ruled out (transport, diffusion, or leakage of NO_2^-
663	from the cell). Indeed, for the enzyme-level kinetic isotope effect to be expressed, some leakage
664	of NO_2^- from the cell must occur. Given that diffusion and transport processes are expected to
665	have normal kinetic isotope effects, these processes could cause some "renormalization" of the
666	observed kinetic isotope effect relative to the theoretical values.
667	Two of the common causes for apparent inverse isotope fractionation have now been
668	explored: pre-equilibrium (Section 4.1.1.) and enzyme reversibility (Section 4.1.2.), neither of
669	which appear to be the underlying mechanism for inverse kinetic isotope fractionation during
670	nitrite oxidation. Indeed, evaluation of the equilibrium isotope effect for ¹⁵ N exchange between

 NO_2^- and NO_3^- suggested that the fractionation factor for nitrite oxidation is likely to be inverse.

672 Evaluation of the fractionation factor using transition state theory provided an independent

673 confirmation of the inverse enzyme-level inverse isotope effect for nitrite oxidoreductase. While

there are uncertainties in each of these calculations, the fact that a new N-O bond is being formed

675 without concurrent bond breakage in NO_2^- is a unique aspect of the nitrite oxidation reaction.

676 Overall, the reaction results in increased bonding to the central N atom, and the observed inverse

677 isotope effect suggests that the enzyme-bound transition state contains the new N-O bond.

678

679 4.2. Implications for Nitrogen Isotope Biogeochemistry

680 An enzyme-level inverse kinetic isotope effect is unusual in biochemistry. The observation 681 here of a large inverse kinetic isotope effect in the oxidation of nitrite by N. mobilis is supported 682 qualitatively by theory, although further work is required to understand whether other cellular 683 processes might mask an even higher inverse kinetic isotope effect. Understanding the cellular 684 controls on the isotope effect for nitrite oxidation will allow us to better predict how the isotope 685 effect for this process is expressed in the environment and how it should be parameterized in 686 ocean biogeochemical models. At this point, it is important to recognize that this process occurs 687 with a fractionation factor that is inverse (less than 1), which challenges the commonly held 688 assumption that all biochemical processes fractionate against the heavier isotope.

An inverse isotope effect for nitrite oxidation may help explain the large $\Delta \delta^{15}N$ (= $\delta^{15}N_{NO3}$ - $\delta^{15}N_{NO2}$) values observed in the eastern tropical north Pacific (ETNP), mentioned above. In a previous study, nitrite oxidation was considered as a mechanism for altering the $\delta^{15}N_{NO3}$ and $\delta^{18}O_{NO3}$ profiles, particularly on the upper edge of the suboxic zone, where calculations suggested that nitrite and oxygen may be consumed concurrently (CASCIOTTI and MCILVIN,

2007). However, adequate explanations for the low $\delta^{15}N_{NO2}$ values and high $\Delta\delta^{15}N$ values were 694 not found, holding to constraints of a 25% isotope effect for nitrate reduction ($^{15}\varepsilon_{k,NAR} =$ 695 $({}^{15}\alpha_{k,NAR}-1)*1000)$ and either a 'normal' or non-fractionating nitrite consumption process. For 696 example, nitrite reduction with a normal isotope effect (${}^{15}\varepsilon_{k,NIR}$) of +15% (BRYAN et al., 1983; 697 CASCIOTTI, 2002; CASCIOTTI et al., 2002) would tend to cause isotopic enrichment in the NO₂⁻ 698 699 pool, counteracting the isotopic depletion of the nitrite pool generated by nitrate reduction. At steady state, that is for nitrite concentrations and $\delta^{15}N_{NO2}$ values not changing (or changing 700 slowly relative to the production and consumption fluxes), the $\Delta \delta^{15}$ N value would tend towards 701 $^{15}\varepsilon_{k,NAR}$ - $^{15}\varepsilon_{k,NIR}$, or 25‰ - 15‰ = 10‰. The observed $\Delta\delta^{15}N$ values of 28-35‰ are therefore 702 703 difficult to explain with nitrite reduction as the sole sink for nitrite. Non steady-state conditions between nitrite production and consumption processes may contribute to $\Delta \delta^{15}N$ values greater 704 than 25‰, but are unlikely to explain the low $\delta^{15}N_{NO2}$ values. 705

If nitrite oxidation by marine NOB routinely occurs with an inverse kinetic isotope effect, as observed here for *N. mobilis*, then nitrite oxidation would tend to lower $\delta^{15}N_{NO2}$ values and enhance the ¹⁵N depletion of NO₂⁻ achieved by nitrate reduction. A simple steady-state box model including the fluxes of nitrite reduction (F_{NIR}), nitrite oxidation (F_{NXR}), and nitrate reduction (F_{NAR} = F_{NIR} + F_{NXR}), was used to estimate the dependence of $\Delta\delta^{15}N$ on F_{NXR}/F_{NAR}.

711

712
$$\Delta \delta^{15} N = {}^{15} \varepsilon_{k,NAR} - {}^{15} \varepsilon_{k,NXR} * F_{NXR} / F_{NAR} - {}^{15} \varepsilon_{k,NIR} * F_{NIR} / F_{NAR}$$
(8)

713

Equation 8 predicts that $\Delta \delta^{15}$ N values should increase as nitrite oxidation becomes a greater fraction of the nitrite consumption flux. Figure 4 shows the steady state δ^{15} N_{NO2}, δ^{15} N_{NO3}, and $\Delta \delta^{15}$ N values obtained from this exercise, as well as the relationship between $\Delta \delta^{15}$ N and 717 F_{NXR}/F_{NAR} (inset). For these calculations, ${}^{15}ε_{k,NAR}$ was assumed to be +25‰, ${}^{15}ε_{k,NXR} = -13‰$,718and ${}^{15}ε_{k,NIR} = +15‰$. Data from the ETNP are overlaid on $\Delta\delta^{15}N$ contours in Figure 4,719illustrating the high $\Delta\delta^{15}N$ values and low $\delta^{15}N_{NO2}$ values observed in the ETNP suboxic zone720(CASCIOTTI and MCILVIN, 2007). The data from the ETNP oxygen minimum zone suggest that721nitrite oxidation may be an important sink for NO2⁻ in this region ($\Delta\delta^{15}N > 28‰$, F_{NXR}/F_{NAR} 722>0.7) (Figure 4). This is consistent with previous interpretations based on $\Delta(15,18)$ (CASCIOTTI

723 and MCILVIN, 2007).

724 More generally, the opposing isotope effects for nitrite oxidation and reduction processes suggest that $\Delta \delta^{15}$ N might be used to distinguish between oxidative and reductive fates for nitrite. 725 Where reductive fates for nitrite dominate over oxidative fates ($F_{NIR} < 0.5 * F_{NAR}$), $\Delta \delta^{15}$ N would 726 be expected to be less than the isotope effect for nitrate reduction to nitrite in a given setting 727 728 (approximately 25‰ in suboxic zones, or 5‰ in the euphotic zone) (Equation 8). Conversely, 729 where the consumption of nitrite through oxidation exceeds the flux of nitrite consumed by reductive processes ($F_{NXR} > 0.5 * F_{NAR}$), $\Delta \delta^{15}$ N would be expected to be greater than the isotope 730 731 effect for nitrate reduction (Equation 8). These isotope effects are most likely to be observed in 732 or near ocean suboxic zones with high concentrations of NO₂. However, the isotope effect for 733 nitrite oxidation may also be expressed in the euphotic zone or in the primary nitrite maximum at 734 the base of the euphotic zone, where competing fates for nitrite may lead to branching between 735 nitrite assimilation into particulate nitrogen and nitrite oxidation to nitrate.

736

737 **4.3. Future Work**

738 It will be important to expand these observations beyond *N. mobilis* to examine the generality 739 of these results among nitrite oxidizing microorganisms and to obtain comparable data on δ^{18} O

740 fractionation by nitrite-oxidizing bacteria. There are 4 known genera of nitrite oxidizing bacteria,

741 which differ from each other in many ways: Nitrococcus, Nitrobacter, Nitrospira, and

742 *Nitrospina*, and it will be important to determine whether the inverse isotope effect is universal

among different species and growth conditions.

Examination of the δ^{18} O fractionation factor during nitrite oxidation holds promise for 744 independently verifying the conclusions from the current study based on δ^{15} N. For example, 745 746 these additional measurements have the power to independently verify the reversibility of the 747 reaction, the structure of the transition state, and the reaction mechanism. At the same time, interpreting the δ^{18} O data is more complex because of the analytical challenges posed by abiotic 748 749 equilibration of oxygen atoms between nitrite and water (CASCIOTTI et al., 2007), as well as the 750 potential for biochemical exchange of nitrite with water catalyzed by the bacteria themselves 751 (DISPIRITO and HOOPER, 1986; FRIEDMAN et al., 1986; HOLLOCHER, 1984). Further work is 752 required to fully resolve the effects of oxygen isotope equilibration from the interpretation of the 753 kinetic isotope effect for oxygen isotopes, since in some situations equilibration will act in the same direction to lower $\delta^{18}O_{NO2}$, and in other cases it will have counteracting effects on $\delta^{18}O_{NO2}$. 754 755 It is also important to examine the kinetic isotope effects with isolated nitrite oxidoreductase enzyme to resolve the contribution of complicating factors, such as NO₂⁻ transport and diffusion, 756 757 to the observed enzyme-level kinetic isotope effect. Nitrite oxidoreductase remains active in cell 758 extracts (TANAKA et al., 1983; YAMANAKA and FUKUMORI, 1988) and can be partially purified 759 for use in these experiments.

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Acknowledgements. The author would like to acknowledge stimulating conversations with John
 Hayes and Carolyn Buchwald during the preparation of this manuscript. Comments on an earlier

- 763 draft were provided by Julie Granger, Tom Trull, and by the author's lab group: Carolyn
- 764 Buchwald, Dan Rogers, Erin Banning, Matthew McIlvin, Edward Leadbetter, and Caitlin Frame.
- 765 Excellent technical assistance was also provided by Matthew McIlvin. Funding from NSF award
- 766 OCE 05-26277 to KLC is also gratefully acknowledged.

768 TABLES

Process	Reaction	$^{15}\alpha_{\rm k}$	$^{15}\epsilon_{k}$ (‰)	References
		$= {}^{14}k/{}^{15}k$	$=(^{15}\alpha_{k}-1)\times 10^{3}$	
Nitrate reduction (Denitrification)	$NO_3^- \rightarrow NO_2^-$	1.013 to 1.030	+13‰ to +30‰	BARFORD et al., 1999; DELWICHE and STEYN, 1970; GRANGER et al., 2006
Nitrite reduction (Denitrification)	$NO_2^- \rightarrow NO$	1.005 to 1.025	+5‰ to +25‰	BRYAN et al., 1983
Nitrous oxide reduction (Denitrification)	$N_2O \rightarrow N_2$	1.004 to 1.013	+4‰ to +13‰	BARFORD et al., 1999; OSTROM et al., 2007
Nitrate reduction (Nitrate assimilation)	$NO_3^- \rightarrow NO_2^-$	1.005 to 1.010	+5‰ to +10‰	GRANGER et al., 2004; NEEDOBA and HARRISON, 2004; WASER et al., 1998a
Nitrogen fixation	$N_2 \rightarrow N_{org}$	0.998 to 1.002	-2‰ to +2‰	DELWICHE and STEYN, 1970; HOERING and FORD, 1960; MEADOR et al., 2007
Ammonium assimilation	$\rm NH_4^+ \rightarrow N_{org}$	1.014-1.027	+14‰ to +27‰	HOCH et al., 1992; WASER et al., 1998a
Ammonia oxidation* (Nitrification)	$\mathrm{NH_4}^+ \rightarrow \mathrm{NO_2}^-$	1.014 to 1.038	+14‰ to +38‰	CASCIOTTI et al., 2003; MARIOTTI et al., 1981; YOSHIDA, 1988; DELWICHE AND STEYN, 1970
Nitrite oxidation (Nitrification)	$NO_2^- \rightarrow NO_3^-$	0.9872	-12.8‰	This study

769 **TABLE 1**: Nitrogen isotope effects for microbial processes in pure culture

770 * ${}^{15}\alpha_k$ for ammonia oxidation is reported relative to NH₄⁺, which is not the substrate for ammonia oxidation. Instead,

771 the observed ${}^{15}\alpha_k$ relative to NH₄⁺ should be corrected for NH₄⁺/NH₃ equilibrium fractionation, dividing the

772 observed ${}^{15}\alpha_k$ by a correction factor of 1.0183 at pH 8 (WEISS et al., 1988), which yields isotope effects of 0.996 to

773 1.0194 for ammonia oxidation relative to NH₃.

	¹ Aqueous phase		² Gas phase		³ Aqueous and ⁴ gas phase		
	ab initio c	alculations	ab initio calculations		observations		
	ω, (cm^{-1})	ω, (c	m^{-1})	ω_{ι}	(cm^{-1})	
	^{14}N	^{15}N	^{14}N	¹⁵ N	^{14}N	¹⁵ N	
NO ₂	795.770	790.733	792.684	787.911	808	805	
	1316.199	1288.788	1338.523	1310.613	1232	1208	
	1386.788	1362.958	1377.186	1353.122	1326	1305	
	This	study	This s	study	³ BEGUN and FL	³ BEGUN and FLETCHER, 1960	
NO ₃ ⁻	708.514	706.678	706.396	704.737	716.8	714.8	
	708.800	706.953	706.518	704.865	716.8	714.8	
	829.003	807.405	837.870	816.030	830.9	809.0	
	1089.308	1089.308	1077.496	1077.496	1049.2	1049.2	
	1438.454	1404.619	1471.120	1436.134	1375.6	1343.7	
	1438.681	1404.820	1471.151	1436.177	1375.6	1343.7	
	This	study	This study		³ BEGUN and FLETCHER, 1960		
HNO ₂	596.345	595.140	600.998	599.845	543.880	542.90	
(trans)	671.375	668.799	630.216	627.954	595.620	593.20	
	906.346	886.079	863.359	844.223	790.118	773.62	
	1308.670	1308.112	1325.358	1324.390	1263.183	1262.45	
	1731.762	1700.779	1794.972	1762.652	1699.800	1669.60	
	3210.523	3210.478	3698.573	3698.576	3590.711	N/D	
	This study		This study		⁴ DEELEY and MILLS, 1985		

774 **TABLE 2**: Molecular vibration frequencies used to calculate isotope exchange equilibria.

¹Calculated in the WebMO interface using Gaussian, B3LYP, 'routine', in aqueous solution.

²Calculated in the WebMO interface using Gaussian, B3LYP, 'routine', in gas phase.

³Observed in aqueous solution (BEGUN and FLETCHER, 1960).

⁴Observed in gas phase as trans-HONO (DEELEY and MILLS, 1985).

TABLE 3: Calculated isotope exchange equilibria

	Medium	$^{15}\alpha_{eq}^{1}$
HNO ₂ /NO ₂	aqueous	0.9975
	gas phase	0.9972
NO ₂ ⁻ /NO ₃ ⁻	aqueous	0.9454
	gas phase	0.9418

782 ¹Calculated from the ratios of Q_2/Q_1 using equations 2 and 3 at T = 298 K and *ab initio* frequencies from Table 2.

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FIGURE 1: Time course results from *N. mobilis* experiments1, 2, and 3.

Figure 1. $[NO_2^-]$ (filled symbols) and $[NO_2^- + NO_3^-]$ (open symbols) are shown for time course experiments 1 (panel A), 2 (panel C), and 3 (Panel E). Please note the use of different x-axes for experiments 1, 2, and 3. $\delta^{15}N_{NO2}$ (filled symbols) and $\delta^{15}N_{NO2+NO3}$ (open symbols) are shown for the same experiments in panels B, D, and F, respectively. In every panel circles represent data from flask A, triangles represent data from flask B, and squares denote data from the control flask. In experiment 2 (panels C and D), there were triplicate A, B, and control flasks, denoted by unique line dashing (solid, dotted, dashed) for replicate flasks. Error bars for $\delta^{15}N_{NO2+NO3}$ and $\delta^{15}N_{NO2+NO3}$ are smaller than the symbols.





Figure 2. Rayleigh plot for nitrite oxidation, using $\ln({}^{15}R_{NO2})$ vs. $\ln([NO_2^-])$, for experiments 1 (filled circles), 2 (open squares), and 3 (filled triangles). The fractionation factor for nitrite oxidation (${}^{15}\alpha_{k,NXR}$) estimated using dummy variables and averaging over all three experiments was 0.9872 ± 0.0015 . For reference, the fractionation line ${}^{15}\alpha_{k,NXR} = 0.9872$ is shown (dotted line). Error bars are smaller than the data points.

FIGURE 3: Modeled effect of enzyme reversibility on expression of KIE



Figure 3. Model simulations for enzyme-level reversibility of nitrite oxidoreductase. Simulated $\delta^{15}N_{NO2}$ (black lines) and $\delta^{15}N_{NO3}$ (grey lines) are plotted versus $f(=[NO_2^{-1}]/[NO_2^{-1}]_0)$ for varying amounts of back-reaction (x = ${}^{14}k_{NAR}/{}^{14}k_{NXR}$; see *Electronic Supplement 2* for model details). Line dashing denotes simulations with a specified amount of back-reaction: x = 0.0 (solid lines), x = 0.2 (dotted lines), x = 0.5 (dashed lines), and x = 1.0 (dot-dashed lines). Overlain in each panel are experimental observations for $\delta^{15}N_{NO2}$ (open black circles) and $\delta^{15}N_{NO3}$ (open grey squares) from experiments 1, 2, and 3 (see text). Error bars for $\delta^{15}N_{NO2}$ measurements are smaller than the symbols. Experimental data are consistent with very low ($0.0 \le x \le 0.2$) amounts of back-reaction.





Figure 4. Contour plot of simulated suboxic zone $\delta^{15}N_{NO2}$ (‰ vs. AIR), $\delta^{15}N_{NO3}$ (‰ vs. AIR), and $\Delta\delta^{15}N$ (= $\delta^{15}N_{NO3}$ - $\delta^{15}N_{NO2}$) for varying F_{NXR}/F_{NAR} . Assumptions for the simulation include: $^{15}\varepsilon_{k,NAR} = +25\%$, $^{15}\varepsilon_{k,NXR} = -12.8\%$, and $^{15}\varepsilon_{k,NIR} = +15\%$, with nitrite concentration and $\delta^{15}N_{NO2}$ held in steady state (see text/supplement for details). Contouring denotes 2‰ increments of $\Delta\delta^{15}N$ (‰) with a range of 12-36‰. Overlaid on the contour plot are observations from the Eastern Tropical North Pacific (ETNP) (black circles; [†]Casciotti and McIlvin, 2007). Inset plot shows the dependence of $\Delta\delta^{15}N$ (‰) on F_{NXR}/F_{NAR} in the simulation. ETNP data show $\Delta\delta^{15}N$ values of 18-36‰, corresponding to simulated F_{NXR}/F_{NAR} values of 0.3-0.9.

CASCIOTTI (W5860) ELECTRONIC ANNEX 1

EA.1. Simulations for Pre-equilibrium Scenarios

In order to illustrate the role of a pre-equilibrium between HNO₂ and NO₂⁻ on the expressed kinetic isotope effect for nitrite oxidation, six different closed-system simulations were conducted with different assumptions for ${}^{15}\alpha_{eq,HNO2/NO2}$, ${}^{15}\alpha_{k,NXR}$, and enzyme substrate (Table EA1.1).

Simulation	Substrate	$^{15}\alpha_{eq,HNO2/NO2}$	$^{15}\alpha_{k,NXR}$
S1	HNO ₂	1.0280	1.0150
S2	NO ₂ ⁻	1.0280	1.0150
S3	HNO ₂	0.9972	1.0150
S4	NO ₂ ⁻	0.9972	1.0150
<u>S</u> 5	HNO ₂	0.9972	0.9872
S6	NO ₂ ⁻	0.9972	0.9872

Table EA1.1. Adjustable parameters for pre-equilibrium simulations.

In each simulation, the reaction was taken to be first order in substrate concentration. It was also assumed that isotope exchange equilibration between HNO_2 and NO_2^- was rapid relative to the rate of oxidation, and that the oxidation reaction itself was irreversible. For simulations 1, 3, and 5 where HNO_2 was taken to be the substrate for oxidation, the following finite-difference equations were used:

$$[^{14}N_{T}] = [^{14}NO_{3}] + [^{14}NO_{2}] + [H^{14}NO_{2}]$$
(EA1.1)

$$[^{14}NO_{3}^{-}]_{t} = [^{14}NO_{3}^{-}]_{t-1} + {}^{14}k_{NXR} \cdot [H^{14}NO_{2}]_{t-1} \cdot dt$$
(EA1.2)

$$[^{14}NO_{2}^{-}]_{t} = ([^{14}N_{T}]_{t} - [^{14}NO_{3}^{-}]_{t})/(1+10^{pKa-pH})$$
(EA1.3)

$$[H^{14}NO_2]_t = ([^{14}N_T]_t - [^{14}NO_3]_t) \cdot (10^{pKa-pH})/(1+10^{pKa-pH})$$
(EA1.4)

$$[{}^{15}N_{T}] = [{}^{15}NO_{3}] + [{}^{15}NO_{2}] + [H^{15}NO_{2}] = constant$$
(EA1.5)

$$[{}^{15}NO_{3}^{-}]_{t} = [{}^{15}NO_{3}^{-}]_{t-1} + ({}^{14}k_{NXR}/{}^{15}a_{k,NXR}) \cdot [H^{15}NO_{2}]_{t-1} \cdot dt$$
(EA1.6)

$$[{}^{15}NO_{2}^{-}]_{t} = ([{}^{15}N_{T}]_{t} - [{}^{15}NO_{3}^{-}]_{t})/(1 + {}^{15}\alpha_{eq,HNO2/NO2} \cdot [H^{14}NO_{2}]_{t}/[{}^{14}NO_{2}^{-}]_{t})$$
(EA1.7)

$$[H^{15}NO_2]_t = {}^{15}\alpha_{eq,HNO2/NO2} \cdot [H^{14}NO_2]_t \cdot [{}^{15}NO_2^{-}]_t / [{}^{14}NO_2^{-}]_t$$
(EA1.8)

For simulations 2, 4, and 6 where NO_2^- was taken to be the substrate for oxidation, the following finite-difference equations were used:

$$[^{14}NO_3^{-}]_t = [^{14}NO_3^{-}]_{t-1} + {}^{14}k_{NXR} \cdot [^{14}NO_2]_{t-1} \cdot dt$$
(EA1.9)

$$[{}^{15}NO_{3}^{-}]_{t} = [{}^{15}NO_{3}^{-}]_{t-1} + ({}^{14}k_{NXR}/{}^{15}a_{k,NXR}) \cdot [{}^{15}NO_{2}]_{t-1} \cdot dt$$
(EA1.10)

Equations for $[{}^{14}N_T]$, $[{}^{15}N_T]$, $[{}^{14}NO_2^-]_t$, $[{}^{15}NO_2^-]_t$, $[H^{14}NO_2]_t$, and $[H^{15}NO_2]_t$ were the same as in simulations 1, 3, and 5 (above).

The simulated δ^{15} N of NO₂⁻, HNO₂, and NO₃⁻ are plotted versus simulated $f(=[NO_2^{-1}]/[NO_2^{-1}]_0)$ in Figure EA1.1. The $\delta^{15}N_{NO2}$ and $\delta^{15}N_{NO3}$ data collected from experiments 1-3 are overlaid on the simulated $\delta^{15}N_{NO2}$ and $\delta^{15}N_{NO3}$ in each panel. Simulations 1, 2, 3, and 4 were all conducted with a normal ${}^{15}\alpha_{k,NXR}$ acting on either HNO₂ (S1, Figure EA1.1A; S3, Figure EA1.1C) or NO₂⁻ (S2, Figure EA1.1B; S4, Figure EA1.1D) as the substrate. Of these first four simulations, only S1 fit the observed data. These results illustrate that in order for the observed $\delta^{15}N_{NO2}$ and $\delta^{15}N_{NO3}$ data to be explained using a normal kinetic isotope effect, HNO₂ (the minor pool) must be the substrate for the reaction, it must be enriched in ${}^{15}N$ relative to the bulk pool, and the equilibrium isotope effect must be about 13‰ higher than the kinetic isotope effect (in this case ${}^{15}\alpha_{eq,HNO2/NO2} = 1.0280$ and ${}^{15}\alpha_{k,NXR} = 1.0150$; Table EA1.1). For example, S1 appears to explain the observations with HNO₂ as the substrate, whereas S2 does not fit the observations with identical fractionation factors, but with the enzyme acting on NO₂⁻ as the substrate rather than HNO₂. However, in order to fit the observations in S1 using a normal kinetic isotope effect, ${}^{15}\alpha_{eq,HNO2/NO2}$ was set to 1.028. As discussed in the main text, however, our best estimate for the equilibrium fractionation factor is estimated to be small (0.9975-0.9978), and HNO₂ is expected to be depleted in 15 N relative to NO₂⁻. Therefore, S1 and S2 can be eliminated.

Constraining ${}^{15}\alpha_{eq,HNO2/NO2}$ to 0.9975 in S3-S6 illustrates that an inverse kinetic isotope effect for either HNO₂ (S5, Figure EA1.1E) or NO₂⁻ (S6, Figure EA1.1F) oxidation is required to explain the observations. A normal kinetic isotope effect for either HNO₂ (S3, Figure EA1.1C) or NO₂⁻ (S4; Figure EA1.1D) cannot explain the observations.



FIGURE EA1.1: Simulated effects of HNO₂/NO₂⁻ pre-equilibrium on expression of ${}^{15}\alpha_{k,NXR}$.

FIGURE EA1.1 LEGEND. Simulated effects of HNO₂/NO₂⁻ pre-equilibrium on expression of ${}^{15}\alpha_{k,NXR}$. The different simulations were carried out using equations EA1.1-EA1.10 and the adjustable parameters (substrates, kinetic and equilibrium isotope effects) given in Table EA1.1. Simulated $\delta^{15}N_{NO2}$ (solid black line), $\delta^{15}N_{HNO2}$ (dashed black line), and $\delta^{15}N_{NO3}$ (solid grey line) are plotted versus simulated f (=[NO₂⁻]/[NO₂⁻]₀), with pH held constant at 8.2. Overlaid in each panel are experimental observations for $\delta^{15}N_{NO2}$ (open black circles) and $\delta^{15}N_{NO3}$ (open grey squares) vs. *f* from experiments 1, 2, and 3 (see main text). Error bars for $\delta^{15}N_{NO3}$ are shown, and for $\delta^{15}N_{NO2}$ are smaller than the symbols.

CASCIOTTI (W5860) ELECTRONIC ANNEX 2

EA2. Model Formulation for Enzyme Reversibility

Reaction reversibility would be expected to affect the time course of $\delta^{15}N_{NO2}$ and $\delta^{15}N_{NO3}$ values and their relationship to $[NO_2^-]$ and $[NO_3^-]$. As noted in the main text, increased reaction reversibility would be expected to lead to the expression of lower (more inverse) ${}^{15}\alpha_{k,NXR}$ values. Therefore, it is important to know whether nitrate reduction is occurring simultaneously with nitrite oxidation in order to best interpret the results of the current study. To first order, reversibility can be evaluated in the nitrite oxidation/nitrate reduction system through mass and isotope balance calculations. Here, a simple finite difference model was used to examine the effects of reaction reversibility on observed [NO₂⁻], [NO₃⁻], $\delta^{15}N_{NO2}$, and $\delta^{15}N_{NO3}$. The results from the model are then compared with the experimental results to evaluate evidence for reaction reversibility in the current experiments.

The following equations comprise the model formulation:

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$$[{}^{14}NO_{2}^{-}]_{t} = [{}^{14}NO_{2}^{-}]_{t-1} + {}^{14}k_{NAR} \cdot [{}^{14}NO_{3}^{-}]_{t-1} \cdot dt - {}^{14}k_{NXR} \cdot [{}^{14}NO_{2}^{-}]_{t-1} \cdot dt$$
(EA2.1)

$$[{}^{14}NO_{3}^{-}]_{t} = [{}^{14}NO_{3}^{-}]_{t-1} - {}^{14}k_{NAR} \cdot [{}^{14}NO_{3}^{-}]_{t-1} \cdot dt + {}^{14}k_{NXR} \cdot [{}^{14}NO_{2}^{-}]_{t-1} \cdot dt$$
(EA2.2)

$$[{}^{15}NO_{2}^{-}]_{t} = [{}^{15}NO_{2}^{-}]_{t-1} + {}^{15}k_{NAR} \cdot [{}^{15}NO_{3}^{-}]_{t-1} \cdot dt - {}^{15}k_{NXR} \cdot [{}^{15}NO_{2}^{-}]_{t-1} \cdot dt$$
(EA2.3)

$$[{}^{15}NO_{3}^{-}]_{t} = [{}^{15}NO_{3}^{-}]_{t-1} - {}^{15}k_{NAR} \cdot [{}^{15}NO_{3}^{-}]_{t-1} \cdot dt + {}^{15}k_{NXR} \cdot [{}^{15}NO_{2}^{-}]_{t-1} \cdot dt$$
(EA2.4)

$$^{15}\alpha_{k,NXR} = {}^{14}k_{NXR} / {}^{15}k_{NXR}$$
 (EA2.5)

$${}^{15}\alpha_{k,\text{NAR}} = {}^{14}k_{\text{NAR}} / {}^{15}k_{\text{NAR}}$$
(EA2.6)

$$^{14}k_{NAR} = x \cdot {}^{14}k_{NXR}$$
 (EA2.7)

$${}^{15}k_{NAR} = x \cdot {}^{15}k_{NXR} \cdot {}^{15}\alpha_{k,NXR} / {}^{15}\alpha_{k,NAR}$$
(EA2.8)

$$\delta^{15} N_{NO2} (\%) = (([^{15} NO_2^{-}]_t / [^{14} NO_2^{-}]_t \div {}^{15} R_{AIR}) - 1) \cdot 1000$$
(EA2.9)

$$\delta^{15} N_{NO3} (\%) = (([^{15} NO_3^-]_t / [^{14} NO_3^-]_t \div {}^{15} R_{AIR}) - 1) \cdot 1000$$
(EA2.10)

In these equations, the molar concentrations (moles L⁻¹) of NO₂⁻ and NO₃⁻ containing ¹⁴N and ¹⁵N are given in brackets, t represents time, dt represents the time step for integration, k's represent first order rate constants (with the dimension t⁻¹) for nitrite oxidation (subscript 'NXR') and nitrate reduction (subscript 'NAR') for ¹⁴N-containing molecules (superscript '14') and ¹⁵N-containing molecules (superscript '15'). The factor 'x' (dimensionless) represents the ratio of reverse and forward rate constants for ¹⁴N-containing molecules. ¹⁵ $\alpha_{k,NXR}$ is taken to be 0.983 and ¹⁵ $\alpha_{k,NAR}$ is taken to be 1.040 so that ¹⁵ $\alpha_{eq,NOX}$ (=¹⁵ $\alpha_{k,NXR}$ /¹⁵ $\alpha_{k,NAR}$) is 0.945, as discussed in the main text. Initial conditions were taken to be [¹⁴NO₂⁻]₀ = 99.6337 x 10⁻⁶ moles L⁻¹, [¹⁵NO₂⁻]₀ = 0.3663 x 10⁻⁶ moles L⁻¹ and [¹⁴NO₃⁻]₀ = [¹⁵NO₃⁻]₀ = 0 x 10⁻⁶ moles L⁻¹.

Four time course 'experiments' were run in MS Excel over 70 time steps (with dt = 1), for x equal to 0, 0.2, 0.5 or 1 (Figure EA2.1). In each case with non-zero reversibility (x>0), initial rates of net nitrite oxidation decreased over time, as the reaction came to equilibrium and the mass fluxes through the forward and reverse reaction became balanced (Figure EA2.1A). Increasing reversibility also lowers the initial net rate of nitrite oxidation and the apparent extent of nitrite oxidation, with correspondingly less nitrate accumulation. The only scenario in which nitrite is completely consumed is with no back reaction (x=0).

In each model run, $\delta^{15}N_{NO2}$ decreased over time as expected with the inverse isotope effect (Figure EA2.1B). Increasing amounts of nitrate reduction caused $\delta^{15}N_{NO2}$ to decrease more quickly for a given amount of nitrite oxidation. From the perspective of $\delta^{15}N_{NO2}$, increasing reversibility lead to an increasingly inverse apparent isotope effect (Figure EA2.1C), and the extent of back reaction was difficult to diagnose from $\delta^{15}N_{NO2}$ alone. However, reaction reversibility also led to unique trends in $\delta^{15}N_{NO3}$ over time that should allow us to distinguish between a large irreversible isotope effect and a contribution from enzyme reversibility.

At low values of x, $\delta^{15}N_{NO3}$ decreased over time as it was dominated by the inverse isotope effect for nitrite oxidation (Figure EA2.1B, x = 0, x = 0.2). As the value of x was increased to 0.5, $\delta^{15}N_{NO3}$ remained relatively constant over the time course (Figure EA2.1B), and in the most extreme case (x=1), $\delta^{15}N_{NO3}$ actually increased over time as it approached equilibrium with NO₂⁻ (Figure EA2.1B). When x was greater than 0.5, the isotope effect for nitrite oxidation based on $\delta^{15}N_{NO3}$ switched from inverse to apparently normal (Figure EA2.1C) while $\delta^{15}N_{NO2}$ fractionation continued to become increasingly inverse. The differential behavior of $\delta^{15}N_{NO2}$ and $\delta^{15}N_{NO3}$ under different levels of reaction reversibility present a diagnostic pattern for reversibility in the $\delta^{15}N_{NO2}/\delta^{15}N_{NO3}$ system, with apparent isotope effects based on $\delta^{15}N_{NO2}$ and $\delta^{15}N_{NO3}$ diverging as the amount of back reaction increases (Figure EA2.1C).

In experiments with *Nitrococcus mobilis*, discussed in the main text, nitrite was completely oxidized to nitrate over the time course experiments (Figure 1). Both nitrite and nitrate become depleted in ¹⁵N over time (Figure 1), with $\delta^{15}N_{NO3}$ approaching the initial $\delta^{15}N_{NO2}$ after all of the nitrite had been oxidized to nitrate. The Rayleigh plots were also consistent with a mass balance between NO₂⁻ and NO₃⁻ and a nearly irreversible reaction of NO₂⁻ to NO₃⁻ under our experimental conditions (Figure 2). Therefore, both the mass and isotope balance observed in *N. mobilis* experiments suggest that reaction reversibility was insignificant under the conditions employed.





FIGURE EA2.1. Results from nitrite oxidoreductase reversibility model. (A) Modeled nitrite (red) and nitrate (blue) concentrations versus time for x ($=^{14}k_{NAR}/^{14}k_{NXR}$) values of 0, 0.2, 0.5, and 1.0. (B) Modeled $\delta^{15}N_{NO2}$ (red) and $\delta^{15}N_{NO3}$ (blue) versus time for the same values of x. (C) Modeled $\delta^{15}N_{NO2}$ (red) and $\delta^{15}N_{NO3}$ (blue) versus $-\ln(f)$ and $-f \cdot \ln(f)/(1-f)$, respectively. In each panel, x = 0 is indicated by the solid line, x = 0.2 by the dot-dashed line, x = 0.5 by the dashed line, and x = 1 by the dotted line.